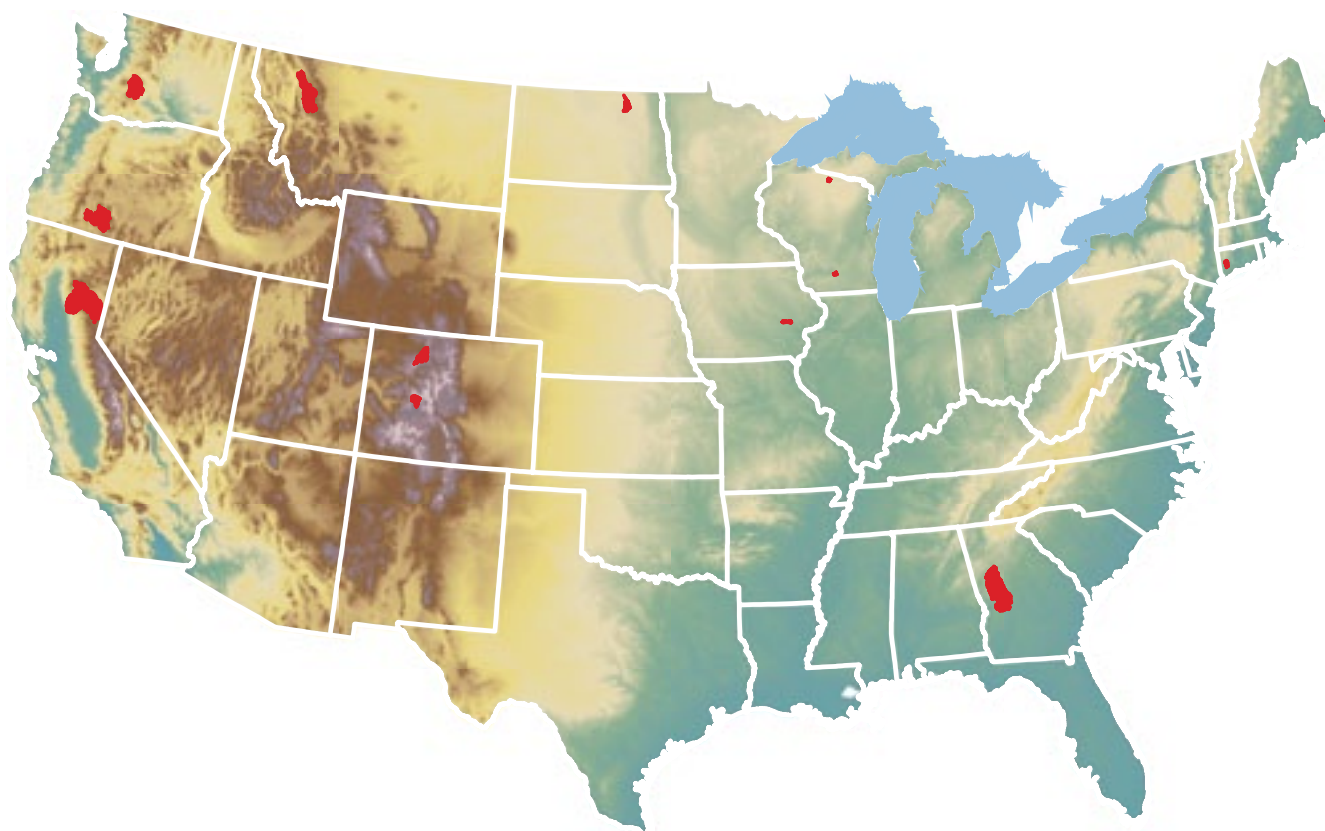


Integrated Watershed-Scale Response to Climate Change for Selected Basins Across the United States



Scientific Investigations Report 2011–5077

Cover. Output from the General Circulation Model used to simulate the response of the Earth's atmosphere to all important driving forces, especially greenhouse gases, and are used for projecting changes in climate globally.

Integrated Watershed-Scale Response to Climate Change for Selected Basins Across the United States

By Steven L. Markstrom, Lauren E. Hay, Christian D. Ward-Garrison, John C. Riskey, William A. Battaglin, David M. Bjerklie, Katherine J. Chase, Daniel E. Christiansen, Robert W. Dudley, Randall J. Hunt, Kathryn M. Koczot, Mark C. Mastin, R. Steven Regan, Roland J. Viger, Kevin C. Vining, and John F. Walker

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Conversion Factors

SI to Inch/Pound

Multiply	By	To obtain
Length		
millimeter (mm)	0.03937	inch (in.)
meter (m)	3.281	foot (ft)
kilometer (km)	0.6214	mile (mi)
Area		
square kilometer (km ²)	247.1	acre
square meter (m ²)	10.76	square foot (ft ²)
square kilometer (km ²)	0.3861	square mile (mi ²)
Volume		
liter (L)	0.2642	gallon (gal)
cubic meter (m ³)	264.2	gallon (gal)
cubic meter (m ³)	35.31	cubic foot (ft ³)
Flow rate		
cubic meter per second (m ³ /s)	70.07	acre-foot per day (acre-ft/d)
meter per second (m/s)	3.281	foot per second (ft/s)
meter per day (m/d)	3.281	foot per day (ft/d)
cubic meter per second (m ³ /s)	35.31	cubic foot per second (ft ³ /s)
cubic meter per day (m ³ /d)	35.31	cubic foot per day (ft ³ /d)
cubic meter per day (m ³ /d)	264.2	gallon per day (gal/d)
Mass		
gram (g)	0.03527	ounce, avoirdupois (oz)
kilogram (kg)	2.205	pound avoirdupois (lb)
Pressure		
kilopascal (kPa)	0.2961	inch of mercury at 60°F (in Hg)
kilopascal (kPa)	0.1450	pound per square inch (lb/ft ²)
Density		
gram per cubic centimeter (g/cm ³)	62.4220	pound per cubic foot (lb/ft ³)
Hydraulic conductivity		
meter per day (m/d)	3.281	foot per day (ft/d)
millimeter per day (mm/d)	0.03937	inch per day (in/d)
Hydraulic gradient		
meter per kilometer (m/km)	5.27983	foot per mile (ft/mi)
Transmissivity*		
meter squared per day (m ² /d)	10.76	foot squared per day (ft ² /d)

Temperature in degrees Celsius (°C) may be converted to degrees Fahrenheit (°F) as follows:

$$^{\circ}\text{F}=(1.8\times^{\circ}\text{C})+32$$

Temperature in degrees Fahrenheit (°F) may be converted to degrees Celsius (°C) as follows:

$$^{\circ}\text{C}=(^{\circ}\text{F}-32)/1.8$$

Vertical coordinate information is referenced to North American Vertical Datum of 1988 (NAVD 88).

Altitude, as used in this report, refers to distance above the vertical datum.

*Transmissivity: The standard unit for transmissivity is cubic foot per day per square foot times foot of aquifer thickness [(ft³/d)/ft²]ft. In this report, the mathematically reduced form, foot squared per day (ft²/d), is used for convenience.

Abbreviations and Acronyms

20C3M	20th century climate emission scenario
A1B	21st century A1B climate emission scenario
A2	21st century A2 climate emission scenario
R ² _a	Adjusted R ² (see Helsel and Hirsh, 2002)
AET	Actual Evapotranspiration
B1	21st century B1 climate emission scenario
CMIP3	World Climate Research Programme's Coupled Model Intercomparison Project phase 3
COOP	Cooperative Observer Program
GCM	General Circulation Models or Global Climate Model
GSFLOW	Groundwater and Surface-Water Flow Model
HRU	Hydrological Response Unit
IPCC	Intergovernmental Panel on Climate Change
NOAA	National Oceanic and Atmospheric Administration
NRCS	Natural Resources Conservation Service
NWIS	National Water Information System
PET	Potential Evapotranspiration
PRMS	Precipitation Runoff Modeling System
SNOTEL	Snow Telemetry
SR	Solar Radiation
SRES	IPCC Fourth Assessment Special Report on Emission Scenarios
USGS	U.S. Geological Survey
WY	Water Year

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Abstract

A study by the U.S. Geological Survey (USGS) evaluated the hydrologic response to different projected carbon emission scenarios of the 21st century using a hydrologic simulation model. This study involved five major steps: (1) setup, calibrate and evaluated the Precipitation Runoff Modeling System (PRMS) model in 14 basins across the United States by local USGS personnel; (2) acquire selected simulated carbon emission scenarios from the World Climate Research Programme's Coupled Model Intercomparison Project; (3) statistical downscaling of these scenarios to create PRMS input files which reflect the future climatic conditions of these scenarios; (4) generate PRMS projections for the carbon emission scenarios for the 14 basins; and (5) analyze the modeled hydrologic response. This report presents an overview of this study, details of the methodology, results from the 14 basin simulations, and interpretation of these results.

A key finding is that the hydrological response of the different geographical regions of the United States to potential climate change may be different, depending on the dominant physical processes of that particular region. Also considered is the tremendous amount of uncertainty present in the carbon emission scenarios and how this uncertainty propagates through the hydrologic simulations.

Introduction

To evaluate the effects of long-term climate change on the freshwater resources of the United States, the U.S. Geological Survey (USGS) Global Change Research and Development Program initiated a study in 2008 (Markstrom and Hay, 2009). A long-term goal of the study was to provide a foundation to analyze the effects of climate change on hydrology across the nation. This work was a collaborative effort between the USGS National Research Program, eight USGS Water Science Centers, and a team of USGS scientists from

the hydrology, hydroclimatology, biology, geography, and natural hazards disciplines.

In this study (fig. 1), 14 basins from different hydroclimatic regions of the United States were selected as systems to be modeled. These 14 basins were all the subject of previous watershed modeling studies using the USGS Precipitation Runoff Modeling System (PRMS, Leavesley and others, 1983). Carbon emission scenarios (Intergovernmental Panel on Climate Change, 2007), hereafter referred to as the scenarios, were simulated by several different General Circulation Models or Global Climate Models (GCMs). The results of the GCM simulations of the scenarios are hereafter referred to as the GCM projections. The GCM projections contain a wide range of climatic conditions, each of which is used by PRMS to simulate hydrologic response for each of the 14 basins. The results of the PRMS simulations of the GCM projections are hereafter referred to as the PRMS projections. The PRMS projections are used to assess the hydrologic effects of climate change on each of the 14 basins.

This report presents an overview of this study, details of the methodology, results from the 14 basin simulations, and interpretation of the results.

Study Sites

The 14 basins (fig. 1) selected as modeling sites in this study have all been modeled previously (table 1) with PRMS by the USGS; these are hereafter referred to as the PRMS applications. The output from a specific run of a PRMS application is hereafter referred to as a PRMS simulation. Once a PRMS application has been successfully calibrated, a PRMS simulation is made using the available historical climate data. This PRMS simulation is hereafter referred to as the baseline condition. These PRMS applications were developed for different reasons, such as evaluation of water supply or ecological habitat. Whenever possible, the original developers of the PRMS applications were engaged in this study to provide local expertise and insight for validation and interpretation of the PRMS projections.

2 Integrated Watershed-Scale Response to Climate Change for Selected Basins Across the United States

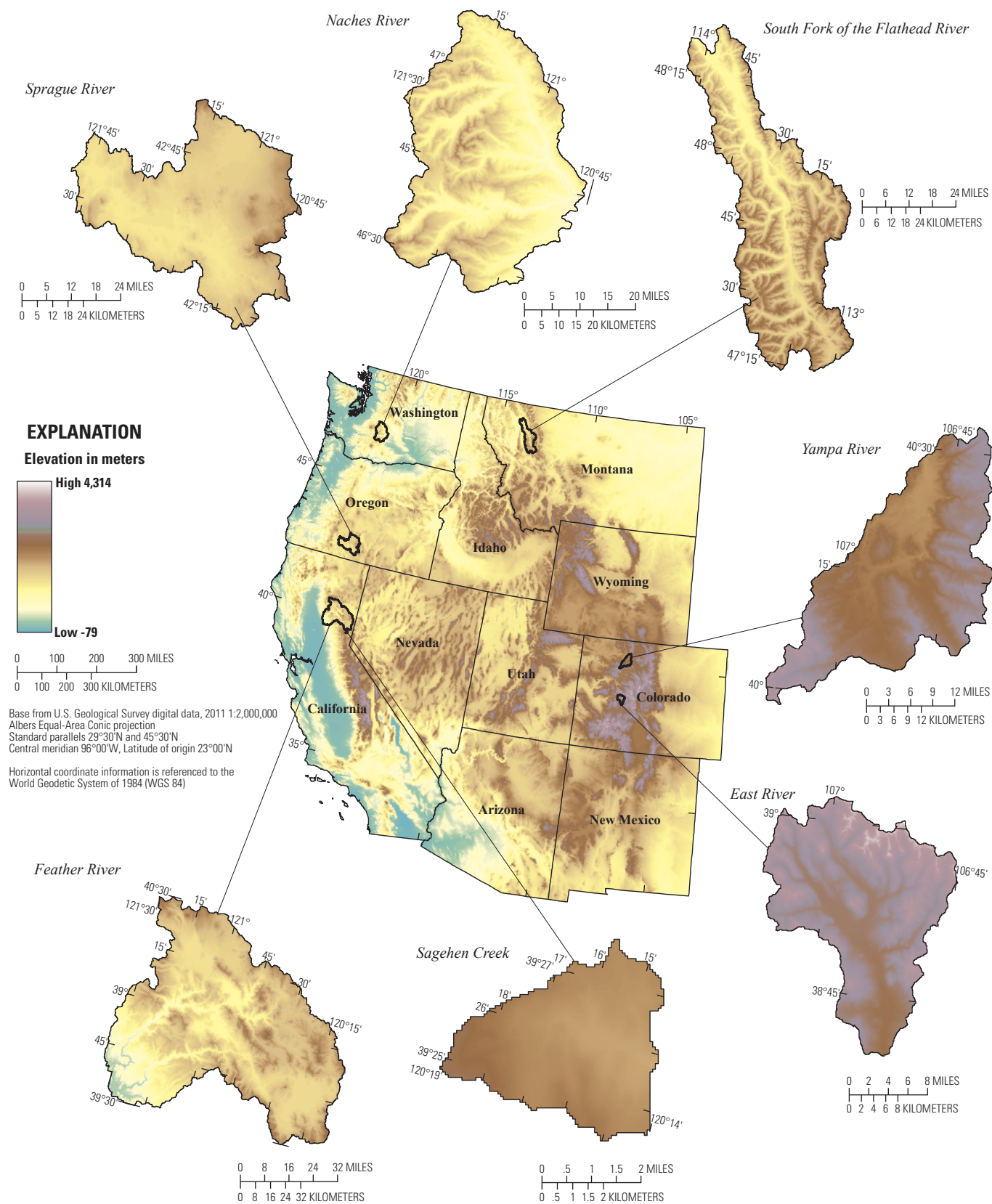


Figure 1. Precipitation Runoff Modeling System application sites for selected basins across the United States. Streamflow-gaging station locations used for model calibration.

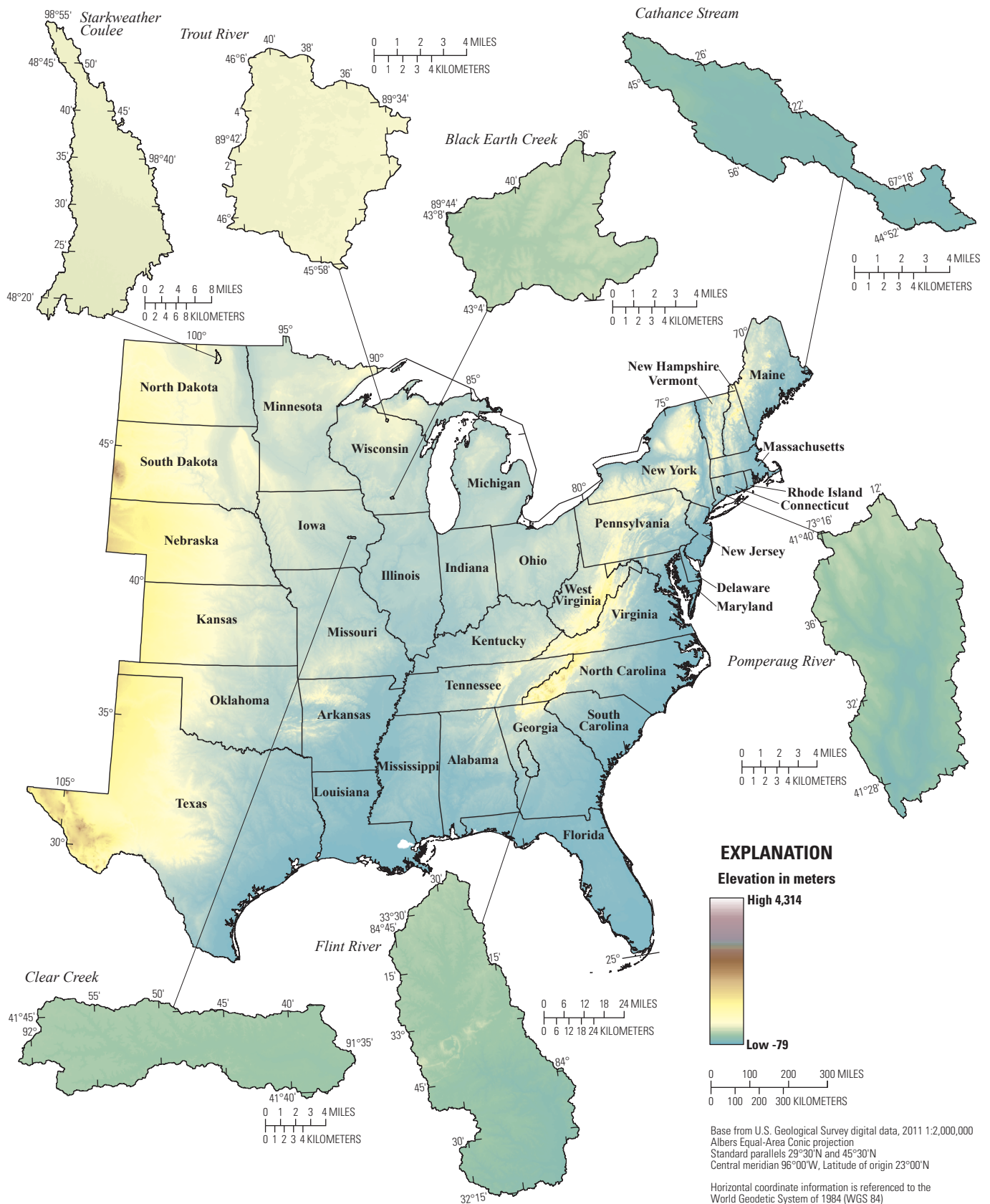


Figure 1. Precipitation Runoff Modeling System application sites for selected basins across the United States. Streamflow-gaging station locations used for model calibration.—Continued

4 Integrated Watershed-Scale Response to Climate Change for Selected Basins Across the United States

Table 1. Drainage area, number of hydrologic response units, and elevation ranges for selected basins across the United States.

[USGS, U.S. Geological Survey; NAVD88, North American Vertical Datum of 1988]

Basin name	Basin abbreviation	USGS streamflow-gaging station identification number	Drainage area upstream from streamflow-gaging station (square kilometers)	Number of hydrologic response units	Elevation range of hydrologic response units (meters above NAVD88)	Published reports
Black Earth Creek at Black Earth, Wisc.	Black Earth	05406500	118	780	718–1,551	4, 5, 8, 9
Cathance Stream at Edmunds, Maine	Cathance	01021230	85	54	46–174	Dudley (2008) 4, 5, 6, 7, 8, 9
Clear Creek near Coralville, Iowa	Clear	05454300	254	75	205–261	4, 5, 7, 8, 9
East River at Almont, Colo.	East	09112500	748	71	2,583–3,796	McCabe and Hay (1994, 1995) 3, 4, 5, 6, 7, 8, 9
Feather River, Calif.	Feather	Not applicable	9,324	324	325–2,212	Koczot and others (2005, 2011) 4, 5, 6, 7, 8, 9
Middle Fork of the Feather River near Merrimac, Calif. ¹	Not applicable	11394500	2,751	58	940–1,962	Koczot and others (2005, 2011) ⁵
South Fork of the Flathead River, Mont. ²	Flathead	12362500	4,307	106	1,045–2,078	Chase (2011) 4, 5, 6, 7, 8, 9
Flint River at Montezuma, Ga.	Flint	02349500	7,511	128	90–286	Viger and others (2009, 2011) 4, 5, 7, 8, 9
Naches River below Tieton River, Wash.	Naches	12494000	2,437	363	562–1,650	Mastin and Vaccaro (2002) 4, 5, 6, 8, 9
Pomperaug River at Southbury, Conn.	Pomperaug	01204000	194	55	63–332	Bjerklie and others (2009, 2011) 4, 5, 7, 8, 9
Sagehen Creek near Truckee, Calif.	Sagehen	10343500	27	128	1,941–2,589	Markstrom and others (2008) 4, 5, 6, 8, 9
Sprague River near Chiloquin, Oreg.	Sprague	11501000	4,053	102	1,316–2,203	Hay and others (2006a, 2009), 4, 5, 6, 7, 8, 9
Starkweather Coulee near Webster, N. Dak.	Starkweather	05056239	543	50	446–491	Vining (2002) 4, 5, 7, 8, 9
Trout River at Trout Lake, Wisc.	Trout Lake	05357245	120	146	491–522	4, 5, 8, 9
Yampa River at Steamboat Springs, Colo.	Yampa	09239500	1,471	68	2,124–3,504	Hay and others (2006b, 2006c, 2011c) 3, 4, 5, 6, 7, 8, 9

¹ Middle Fork is a gaged interior basin of the Feather River Basin.

² For the South Fork of the Flathead River, natural daily mean streamflow values calculated by the Bureau of Reclamation were used for calibration.

³ Included in analysis by Battaglin and others (2011).

⁴ Included in analysis by Christiansen and others (2011).

⁵ Included in analysis by Hay and others (2011).

⁶ Included in analysis by Mastin and others (2011).

⁷ Included in analysis by Milly and Dunne (2011).

⁸ Included in analysis by Risley and others (2011).

⁹ Included in analysis by Walker and others (2011).

The 14 basins can be grouped into 6 geographical regions of the United States; these regions are:

- The Naches and Sprague River Basins in the Cascade Mountains of the Pacific Northwest region of the United States.
- The Feather River and Sagehen Creek Basins in the Sierra Nevada region of the United States.
- The South Fork of the Flathead, East, and Yampa River Basins in the Rocky Mountain region of the United States.
- The Starkweather Coulee, Clear Creek, Black Earth Creek, and Trout Lake Basins in the Midwestern region of the United States.
- The Cathance Stream and Pomperaug River Basins in the Northeastern region of the United States.
- The Flint River Basin in the Southeastern region of the United States.

Description of the Simulation Models Used in this Study

This study uses two different types of simulation models. The first type is the GCM models, which simulate the response of the Earth's atmosphere to all the important driving forces, especially greenhouse gases, and are used for projecting changes in climate globally. These models have been run by members of the Intergovernmental Panel on Climate Change (IPCC) and their output was made available for download, as described below in section "Step 1: Select and Acquire General Circulation Model Projections." The second type of model is the watershed hydrology model PRMS, hereafter referred to as the PRMS model. The PRMS model is applied at a much finer spatial resolution and can generate detailed hydrologic output which typically is not produced by the GCMs. Specifically, the average resolution used in the 14 PRMS applications is less than 14 square kilometers (km²), whereas the resolution of grid cells used by the GCM models are larger than one degree

of latitude and longitude (table 2). For mid-latitude basins, this could represent resolutions as large as 18,000 km² per grid cell. The PRMS model was coupled to the GCMs through a downscaling procedure (fig. 2). The GCM models, PRMS model, and downscaling procedure, as applied in this study, are described below in the section "Methods."

General Circulation Models

GCMs are numerical models that represent all components of the Earth system including the atmosphere, oceans, cryosphere, and land surface. The climate-change signal from individual GCMs (that is, the difference between simulated future climate and simulated historical climate) can vary in direction and magnitude because of the uncertainties associated with each GCM parameterization, initial forcings, emission scenarios, and representation of the Earth system components (Fealy and Sweeney, 2008). The same forcings can produce different GCM responses because of differences in the method for modeling climate system and feedbacks and the initial conditions used for the climate simulations. Given this parametric and structural uncertainty in climate modeling, it is desirable to use output from more than one GCM to obtain a range of potential future climatic conditions. Ensembles of multiple GCM outputs provide an envelope of possible regional changes and their accompanying uncertainties (Murphy and others, 2004; Boorman and Sefton, 1997; Intergovernmental Panel on Climate Change, 2007).

Representing GCM uncertainties using multiple GCM outputs is now feasible because of the large data archives available at the GCM modeling centers. The World Climate Research Programme's Coupled Model Intercomparison Project phase 3 (CMIP3) multimodel dataset archive, which was used in the IPCC Fourth Assessment Special Report on Emission Scenarios (SRES) (Intergovernmental Panel on Climate Change, 2007) contains 28 archived GCMs with multiple emission scenarios. The emission scenarios describe how greenhouse gas emissions could evolve between 2000 and 2100, depending on various hypotheses on what the state of the future world may be. The range in results from the various GCMs can then be compared by emission scenario.

Table 2. General Circulation Model projections used in this study.

GCM	Description	Cell size (latitude x longitude)
BCC-BCM2.0	Bjerknes Centre for Climate Research, Norway	2.79° x 2.81°
CSIRO-Mk3.0	Commonwealth Scientific and Industrial Research Organization, Australia	1.865° x 1.875°
CSIRO-Mk3.5	Commonwealth Scientific and Industrial Research Organization, Australia	1.865° x 1.875°
INM-CM3.0	Institute for Numerical Mathematics, Russia	4.0° x 5.0°
MIROC3.2	National Institute for Environmental Studies, Japan	2.79° x 2.81°

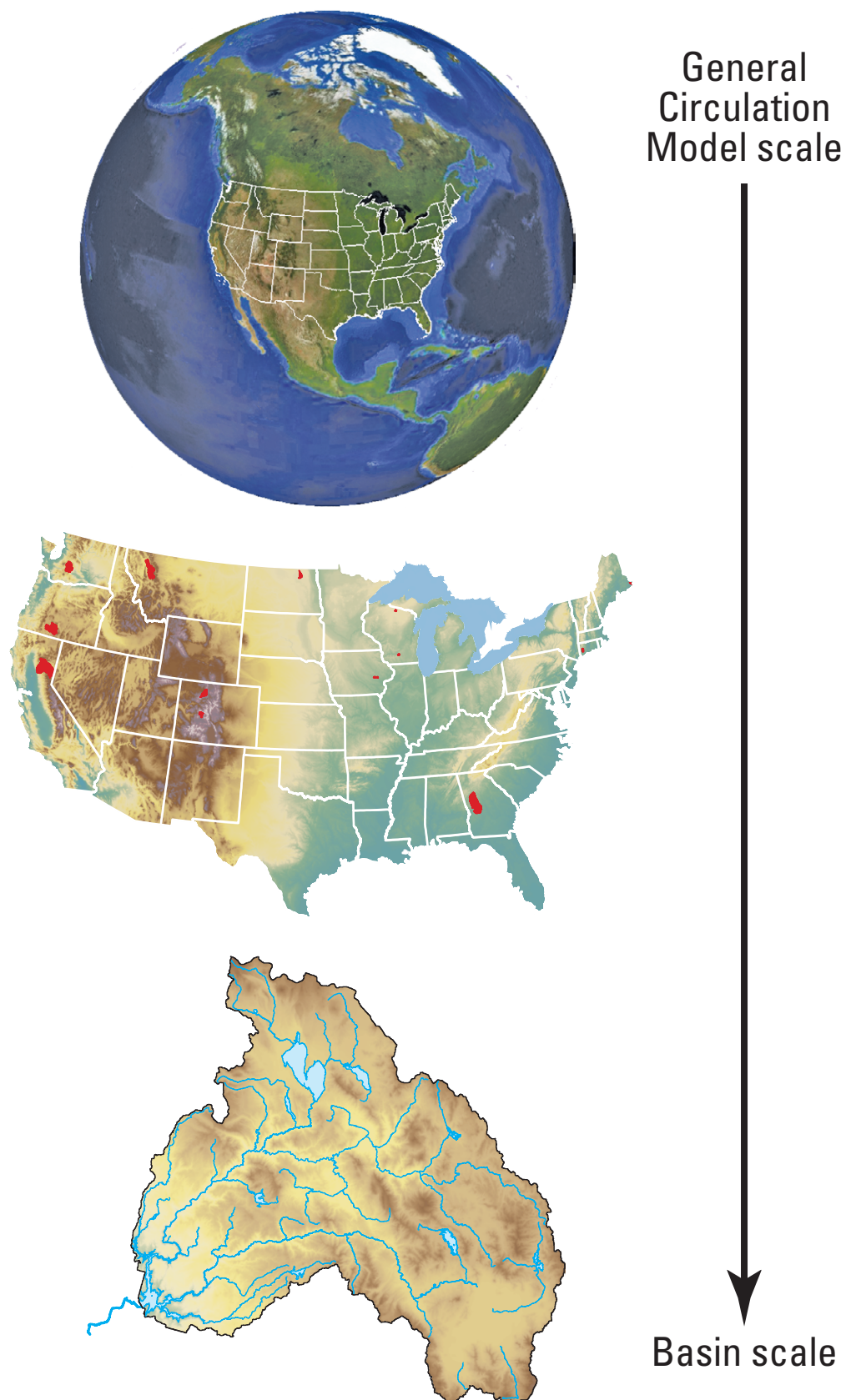


Figure 2. Modeling sequence: output from coarse General Circulation Models used as input to fine scale Precipitation Runoff Modeling System applications across the United States.

Precipitation Runoff Modeling System

The PRMS model is a deterministic, distributed-parameter, process-based watershed model. Each hydrologic component used for generation of streamflow is represented within the PRMS model by a process algorithm that is based on a physical law or an empirical relation with measured or estimated characteristics (see Leavesley and others, 1983; and Markstrom and others, 2008). PRMS is used to simulate and evaluate the effects of various combinations of climate (precipitation, air temperature, and solar radiation) and land use on basin hydrologic response. Response to normal and extreme rainfall and snowmelt can be simulated to evaluate changes in water-balance relations, streamflow regimes, soil-water relations, and groundwater recharge.

Distributed-parameter capabilities are provided by partitioning a basin into hydrologically similar units, using characteristics such as slope, aspect, elevation, vegetation type, soil type, and precipitation distribution. These units are called hydrologic response units (HRUs) and are assumed to be homogeneous with respect to hydrologic response. Water and

energy balances are computed daily for each HRU. Although more sophisticated streamflow routing methods are available in PRMS, in this study, the sum of the responses of all HRUs, weighted on a unit-area basis, produced the daily time step basin hydrologic response.

Climate data consisting of daily measurements of precipitation, air temperature, and solar radiation are the driving variables used to compute evaporation, transpiration, sublimation, snowmelt, infiltration, and runoff in a PRMS simulation. The form of precipitation (rain, snow, or mixture of both) is determined from air temperature or can be specified as input. Precipitation can be intercepted by and evaporated from the plant canopy. Throughfall is precipitation that is not intercepted by the plant canopy and is distributed to the land surface. Throughfall can accumulate as part of the snowpack, be stored on impervious areas, infiltrate into the soil zone, be evaporated and transpired, or become surface-water runoff. Water and energy balances are computed for the snowpack to determine snow accumulation, snowmelt, or sublimation (fig. 3).

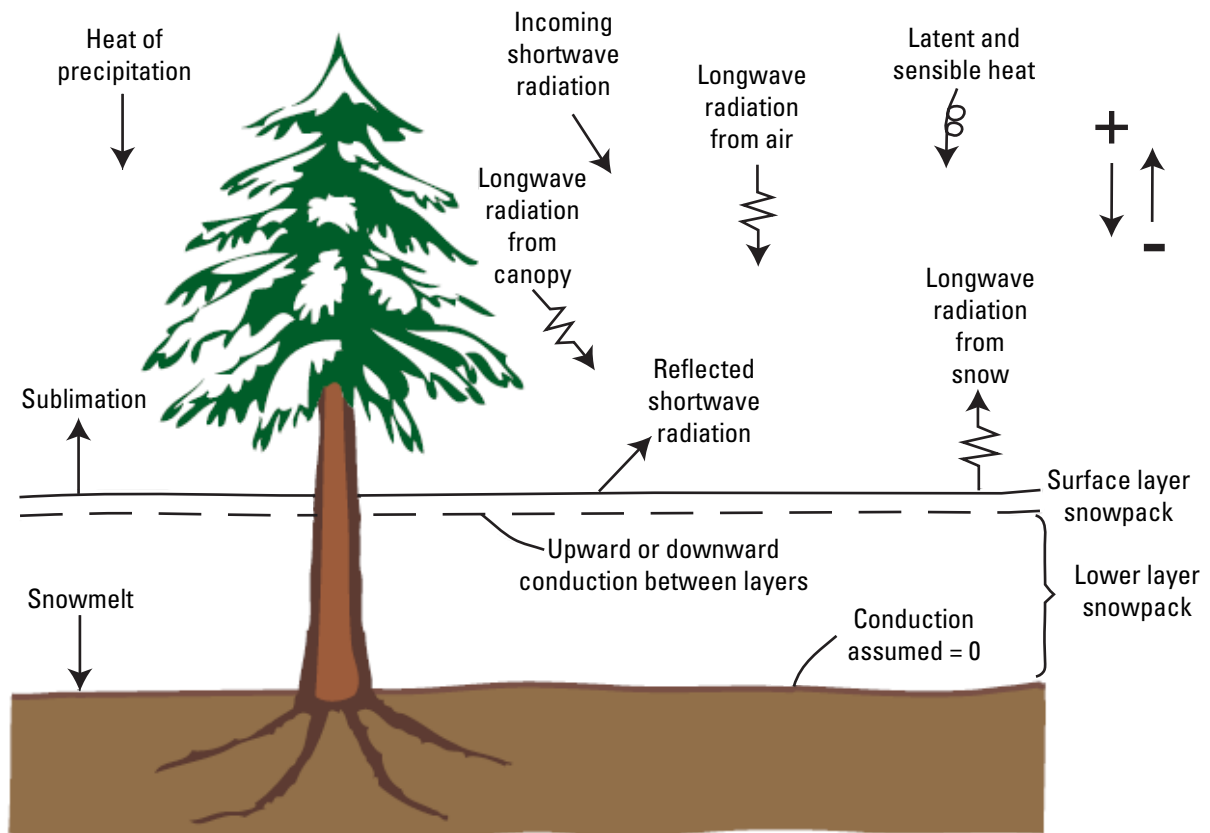


Figure 3. Components of the snowpack energy balance, accumulation, snowmelt, and sublimation (from Kocot and others, 2005).

Precipitation that falls on each HRU is routed through the conceptual storage reservoirs (fig. 4), where water is added and subtracted according to the computational sequence. Each HRU simulates water stored in the plant canopy, snowpack, impervious-zone, capillary, gravity, and groundwater reservoirs. Reservoirs in the HRU are assumed to be homogeneous in hydrologic response and model parameterization. Water entering any HRU reservoir is considered to be instantaneously mixed with water previously in the reservoir.

Several input files are needed for a PRMS simulation (Markstrom and others, 2008, p. 139–150). This study was conducted by modifying two of the PRMS input files. These files are the PRMS Data File and the PRMS Parameter File. The PRMS Data File contains time-series climate data used to drive the PRMS simulation. These input data include daily values of precipitation and maximum and minimum air temperature. The different carbon emission scenarios were simulated using PRMS by creating an input PRMS Data File corresponding to each scenario. The PRMS Parameter File contains

static values which characterize the physical and empirical aspects of the basin. Only the values of the “spring_frost” and “fall_frost” parameters (table A1–1) were changed in this file according to the information developed for each carbon emission scenario. These parameters are described in the description of the “transp_frost_prms” module in Appendix 1.

PRMS has the capability of producing more than 200 output variables which characterize the simulated hydrologic conditions for all HRUs on any day of the simulated time period (some are listed in Markstrom and others, 2008, p. 138, table A1–2). For this study, only 19 of these output variables were considered for analysis (table 3). These variables include streamflow, evapotranspiration, groundwater storage and inflow, soil moisture, and snowpack water equivalent and snowmelt. The results of running PRMS on 14 basins, four carbon emission scenarios from five GCMs for 100 years of the 21st century, and outputting 35 variables at both HRU and basin-aggregated resolutions resulted in almost one terabyte of output data.

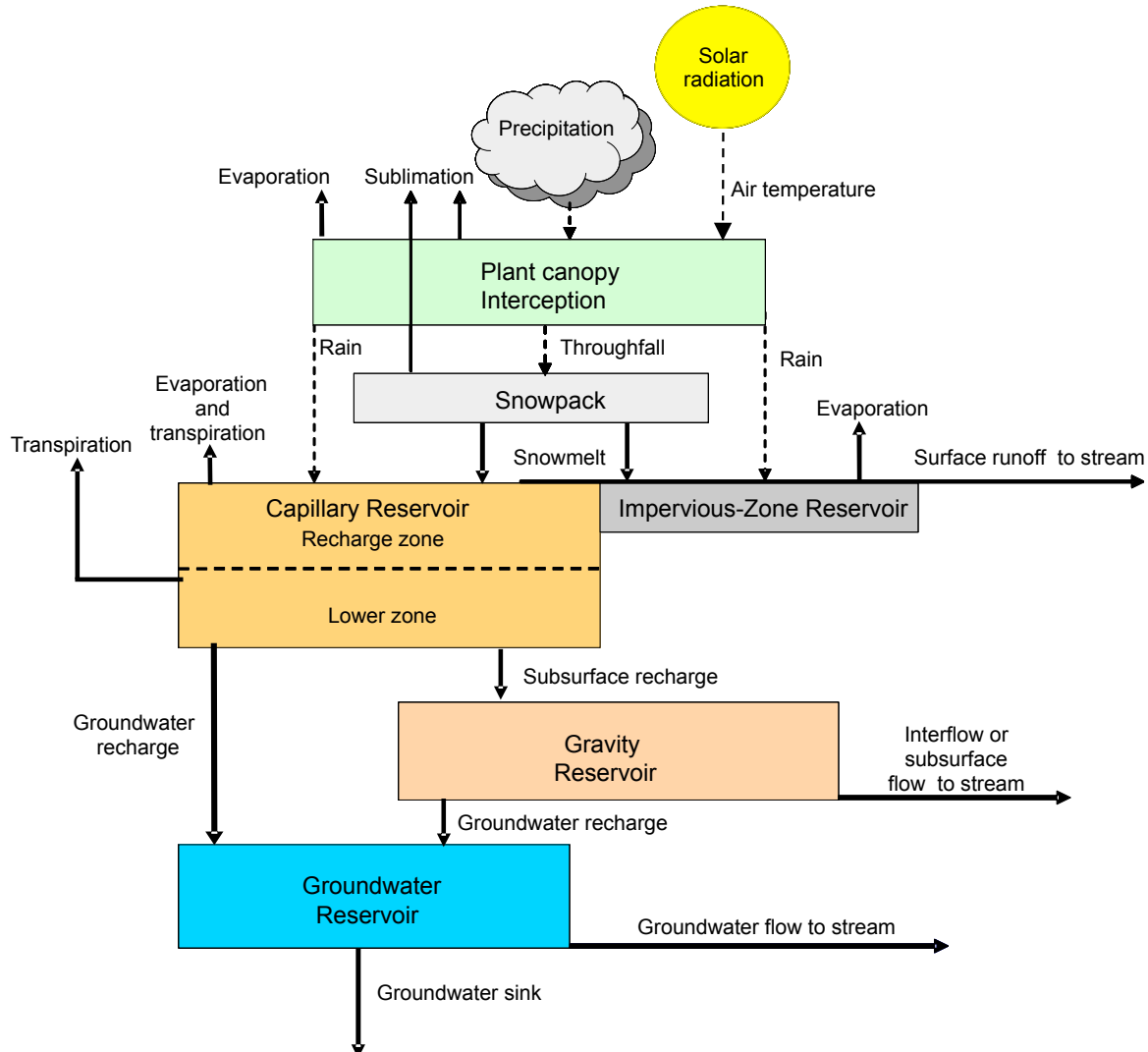


Figure 4. Schematic diagram of a watershed and inputs simulated by the Precipitation Runoff Modeling System model.

Table 3. Description of Precipitation Runoff Modeling System variables analyzed in the climate change study.

[Units of “millimeters” are depth over the basin area per daily timestep]

Variable type	Variable name	Units	Variable description
Input	Precipitation	Millimeters	Area-weighted adjusted average precipitation for basin
	Maximum temperature	Degrees celsius	Basin area-weighted daily maximum temperature
	Minimum temperature	Degrees celsius	Basin area-weighted daily minimum temperature
	Evapotranspiration	Millimeters	Evapotranspiration on basin
	Growing season length	Days	Length of the growing season derived from the temperature data
Streamflow	Streamflow	Cubic meters per second	Streamflow from basin
	Surface runoff	Cubic meters per second	Basin surface runoff
	Subsurface flow	Cubic meters per second	Basin subsurface flow
	Groundwater flow	Cubic meters per second	Basin groundwater flow
Snow	Precipitation that falls as snow	Percent	Percent of precipitation that falls as snow derived from precipitation and snow
	Snow-covered area	Percent	Average snow-covered area for total basin
	Snowpack water equivalent	Millimeters	Average snowpack water equivalent for total basin area
	Snowmelt	Millimeters	Average snowmelt for total basin area
Soil moisture	Infiltration	Millimeters	Basin area-weighted average for infiltration
	Inflow to subsurface reservoirs	Millimeters	Basin weighted average for inflow to subsurface reservoirs
	Inflow to ground-water reservoirs	Millimeters	Basin area-weighted average of inflow to ground-water reservoir (recharge)
	Soil moisture	Millimeters	Basin area-weighted average soil moisture
	Subsurface reservoir storage	Millimeters	Basin weighted average for subsurface reservoir storage
	Groundwater storage	Millimeters	Basin area-weighted average of groundwater storage

Methods

Coupling the GCM projections to the PRMS applications required five major steps: (1) after evaluating a common baseline condition for all of the 14 PRMS applications and recalibrating if necessary; (2) the GCM projections, including precipitation and maximum and minimum temperature, for the baseline period and the period of the scenarios were acquired; (3) the GCM projections were downscaled to the climate stations used for each of the 14 PRMS applications; (4) the PRMS projections were generated from the GCM modified climate station record inputs by running the PRMS model; and (5) the PRMS projections were analyzed for the hydrologic effects of climate change. These steps are described in detail in the following sections.

Step 1: Evaluate Precipitation Runoff Modeling System Applications for Baseline Conditions

The PRMS applications used in this study have all been the subject of previous modeling studies (table 1). These applications were each calibrated for different purposes and for different baseline conditions. The first step in the evaluation of the previous PRMS applications was to find a common baseline period which could be used by all 14 of the applications. A common baseline period was found to start in 1987 because of the availability of measured temperature data from the U.S. Department of Agriculture, Natural Resources Conservation Service, National Water and Climate Center, Snow Survey and Water Supply Forecasting Program (NRCS SNOTEL, <http://www.wcc.nrcs.usda.gov/snow>, accessed July 2010). A common baseline period was found to end in 1999 because many of the 20th century scenarios end in calendar year 1999. These temperature data were used by many of the PRMS applications; consequently, water years (WYs) 1988–99 (October 1, 1987 to September 30, 1999) were chosen as the common baseline period.

Each of the 14 PRMS applications was evaluated for the common baseline period. Climate data from WY 1988 was reserved for initialization of all PRMS simulations. The evaluation was based on four criteria: (1) mean monthly streamflow, where the “mean monthly” is the arithmetic mean of all mean daily values that occurred during a particular month, across all years, resulting in 12 values; (2) monthly mean streamflow, where “monthly mean” is the arithmetic mean of all of the values that occurred in a particular month, resulting in a time series of values, one for each month of each year; (3) mean monthly solar radiation (SR); and (4) mean monthly potential evapotranspiration (PET). All daily values, both measured and simulated, were aggregated to monthly time steps for evaluation.

Measured streamflow data were obtained from the U.S. Geological Survey National Water Information System (USGS NWIS, <http://waterdata.usgs.gov/nwis>, accessed July 2010); however, these data were not always available for all of

the PRMS applications during the common baseline period. In some cases, this was because of incomplete streamflow-gaging station information, or more commonly, because of anthropogenic causes. Natural streamflow (streamflow corrected for withdrawal or return flows or other measurable uses) data records were estimated when possible. If natural streamflow data records could not be estimated, they were not used as evaluation criterion.

Measured monthly SR values are available for 217 stations in the United States (http://tredc.nrel.gov/solar/old_data/nsrdb/redbook/mon2/, accessed July 2010). A multiple linear regression (Helsel and Hirsh, 2002) was developed between these measured monthly SR values and daily climate statistics calculated using data from climate stations collocated with the measured solar radiation. For each month a separate multiple linear regression equation was developed using latitude, longitude, elevation, mean precipitation on days with temperatures greater than 0 °C, mean precipitation, number of rain days, mean maximum temperature, and the difference between mean maximum and mean minimum temperature as possible independent variables. Adjusted R^2 (R^2_a , Helsel and Hirsh, 2002) values ranged from 0.83–0.98. Then, a dataset of mean monthly SR values at each of the climate stations in the National Oceanic and Atmospheric Administration Cooperative Observer Program (NOAA COOP, <http://www.nws.noaa.gov/om/coop>, accessed July 2010) and the NRCS SNOTEL sites was developed using the monthly multiple linear regression equations. The values used as measured mean monthly PET values were derived from the free-water evaporation atlas of Farnsworth and others (1982).

Initial evaluation of the PRMS applications to basin mean monthly values of SR, PET, and streamflow showed varied results. In order to ensure consistency, all PRMS applications were recalibrated to common baseline conditions using Luca (Hay and Umemoto, 2006; Hay and others, 2006b; 2006c), which is a multiple-objective, stepwise, automated procedure that uses the Shuffled Complex Evolution (Duan and others, 1992; 1993; 1994) global search algorithm to calibrate PRMS applications. For this study, SR, PET, and water balance were calibrated for each PRMS application. Luca assures that intermediate model fluxes as well as the water balance are simulated consistently with measured values (Hay and others, 2006c).

PRMS applications that are reliably calibrated for the baseline period are important because any biases are likely to transfer to the future simulations of flow (Prudhomme and Davies, 2009). PRMS simulation of maximum and minimum temperature and precipitation of baseline conditions for all applications are summarized in figure 5. Results from the Luca calibrations of each of the PRMS applications are shown in figure 6. The calibration results shown for the Feather River basin PRMS application are for the Middle Fork of the Feather River basin, the largest sub-basin modeled in the Feather River basin. After calibration, all PRMS applications are judged to produce reasonably accurate mean monthly results for SR, PET, and the water balance (when available) for the baseline period (fig. 6).

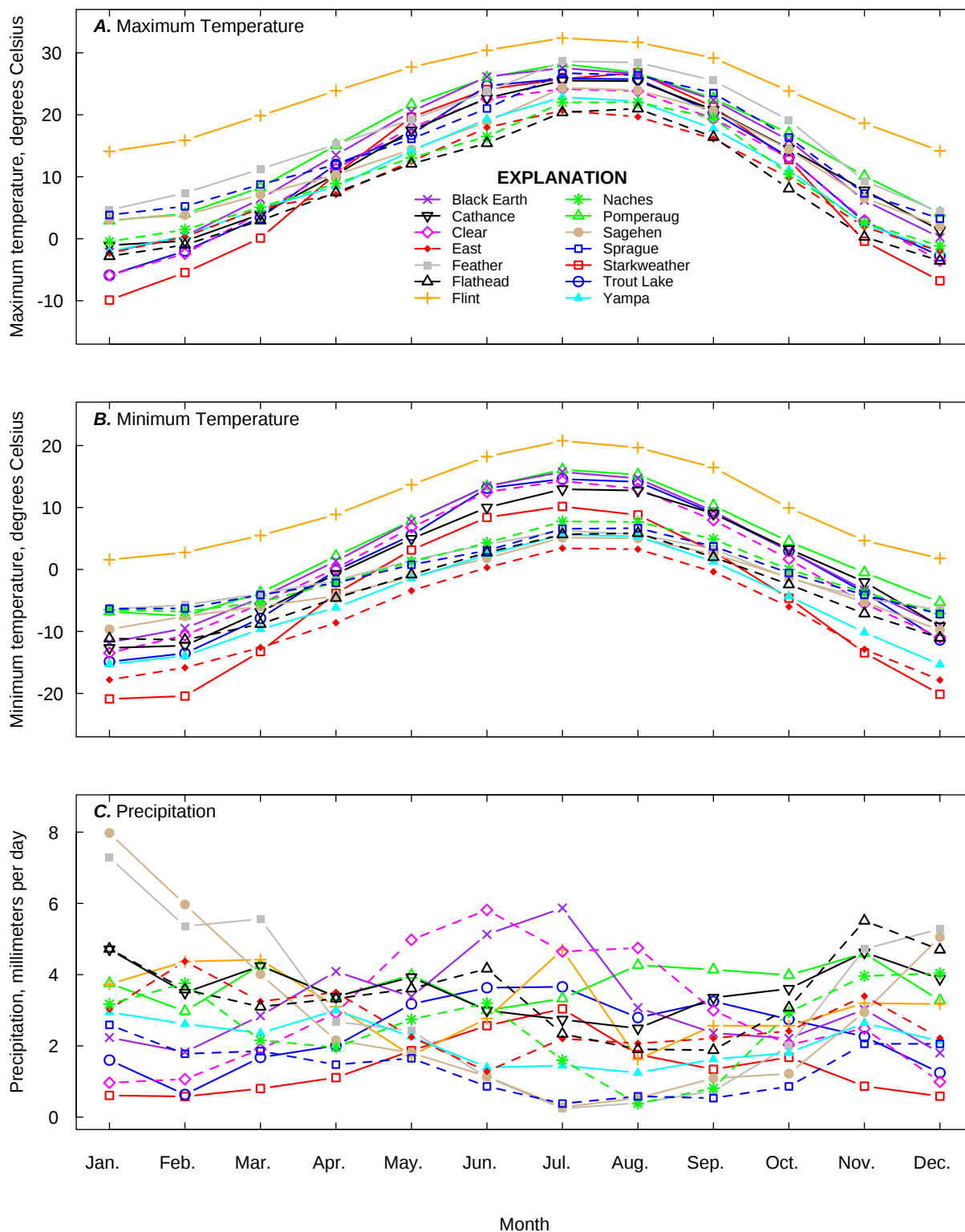


Figure 5. Basin mean monthly (1989–1999): (A) maximum temperature, (B) minimum temperature, and (C) precipitation for the 14 basins.

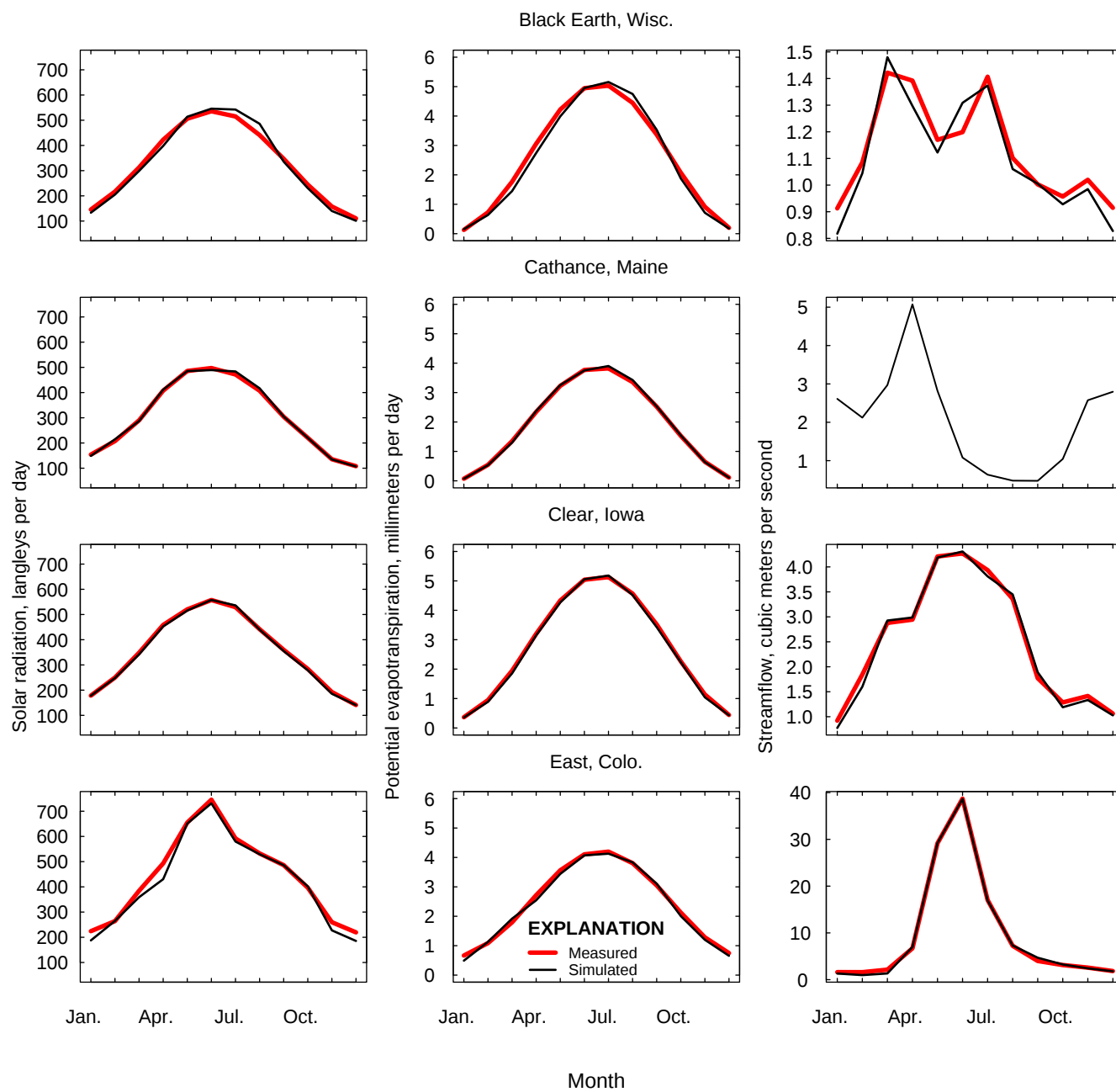


Figure 6. Calibration by basin (1989-1999): mean monthly measured (gray) and Precipitation Runoff Modeling System-simulated (black) solar radiation, potential evapotranspiration, and streamflow.

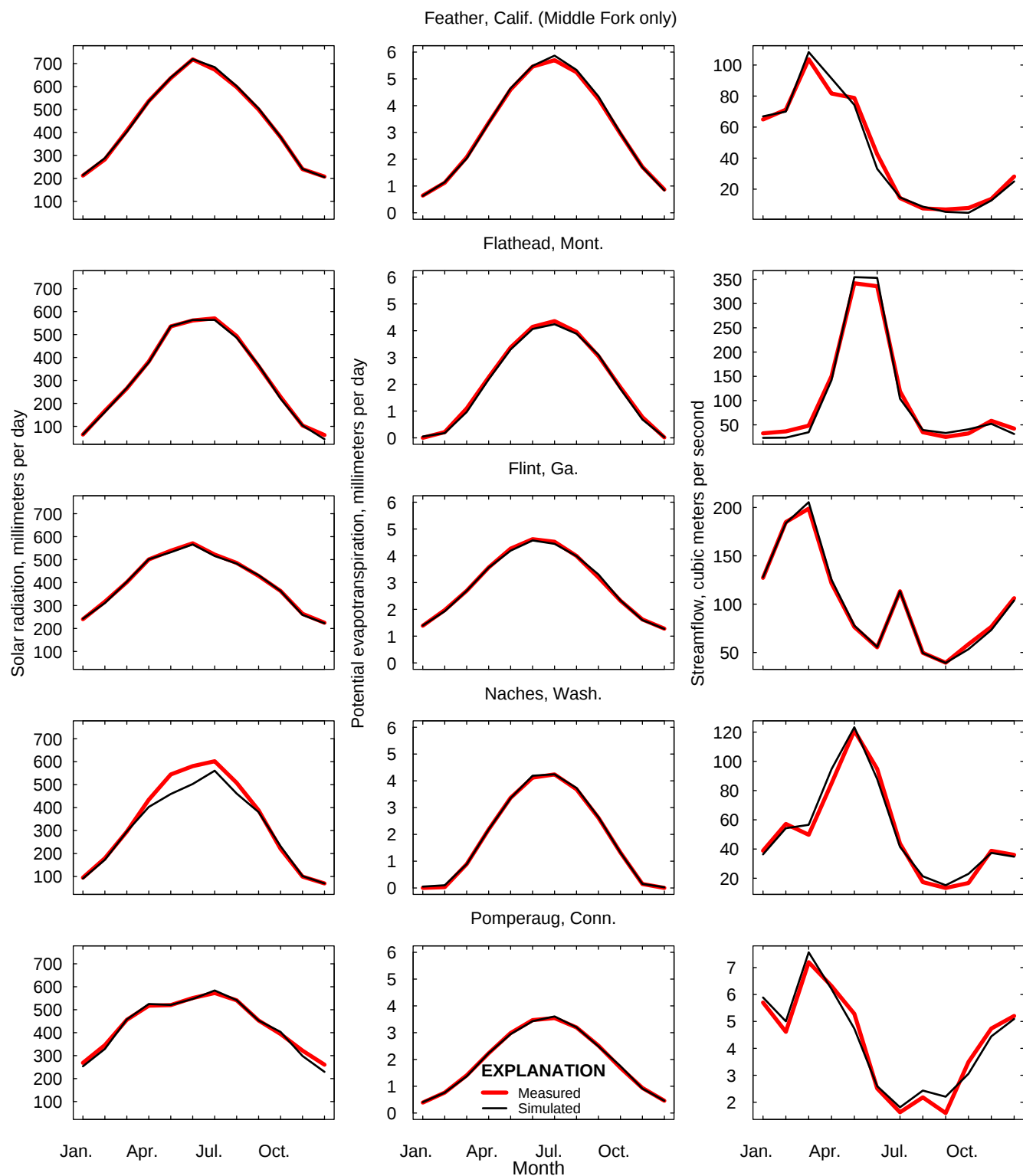


Figure 6. Calibration by basin (1989-1999): mean monthly measured (gray) and Precipitation Runoff Modeling System-simulated (black) solar radiation, potential evapotranspiration, and streamflow.—Continued

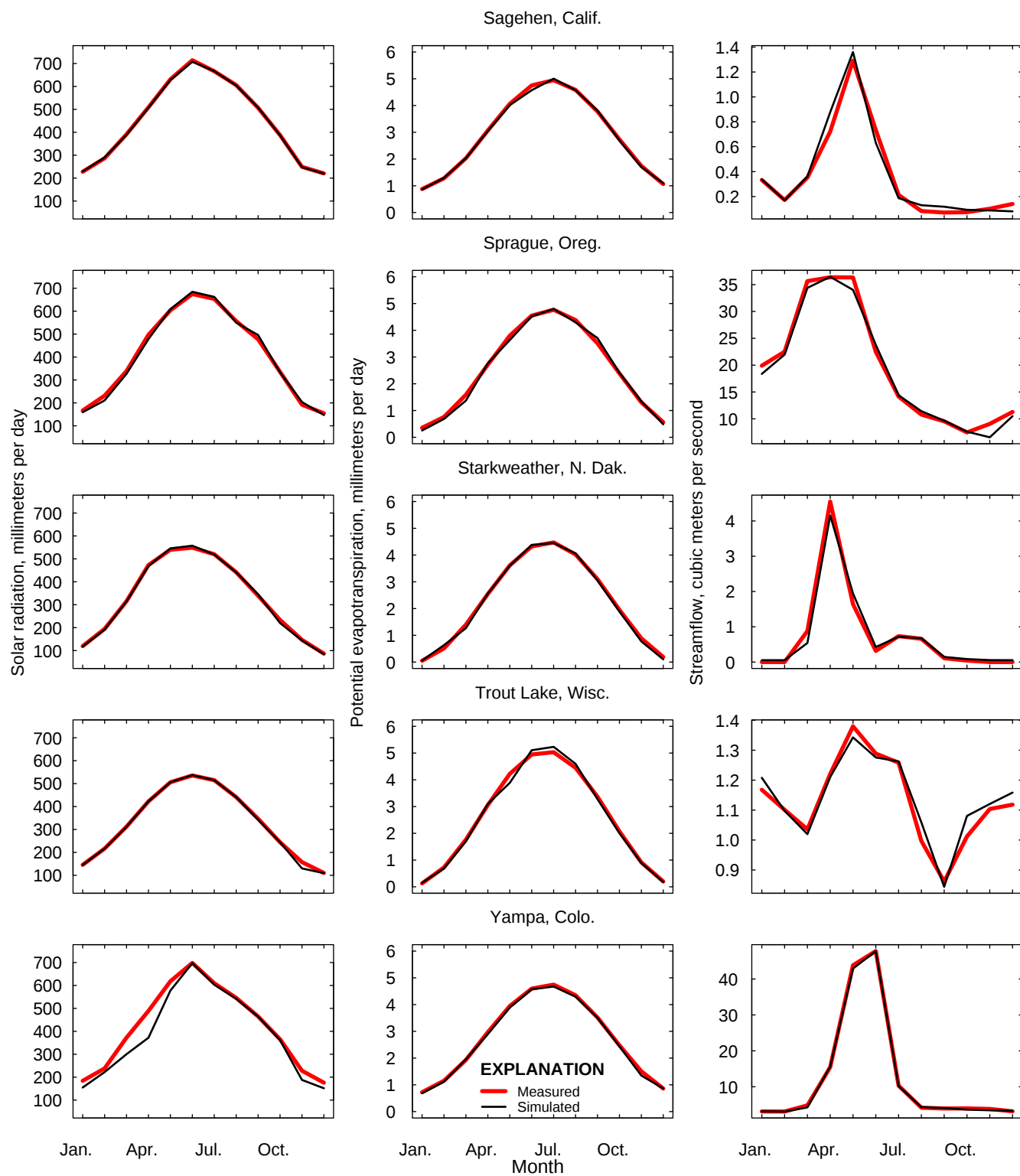


Figure 6. Calibration by basin (1989-1999): mean monthly measured (gray) and Precipitation Runoff Modeling System-simulated (black) solar radiation, potential evapotranspiration, and streamflow.—Continued

Step 2: Select and Acquire General Circulation Model Projections

Four scenarios developed by the IPCC were used in this study (table 4). One of these scenarios (20C3M) corresponds to the latter half of the 20th century. Only the period corresponding to PRMS baseline conditions was considered in this scenario. The other three scenarios represent different levels of carbon emission during the 21st century. The three future emission scenarios lead to different patterns in greenhouse gas emissions and concentrations, from relatively low (B1), to medium (A1B), to high concentrations (A2) during the 21st century (Intergovernmental Panel on Climate Change, 2007). These three scenarios were selected as a representative subset of the other future emission scenarios developed by the IPCC. The total time period covered by these four scenarios was from October 1, 1988 to December 31, 2099.

Because of the uncertainty in any one GCM projection, an ensemble of projections was obtained from multiple GCMs to obtain a range of results. For a GCM projection to be suitable for this study, the World Climate Research Programme's CMIP3 multimodel dataset archive (Intergovernmental Panel on Climate Change, 2007) had to contain monthly projections of precipitation and maximum and minimum temperature for the four scenarios (table 4) during the time period of the study. Only five GCM projections met these criteria (table 2).

Step 3: Downscale General Circulation Model Projections

PRMS projections, at a basin scale, require climatological information at resolutions that are much finer than the grid-size of the highest resolution GCM projections. The process of downscaling takes information produced at coarse spatial resolution and makes it available at finer spatial resolution. In this study, the GCM projections were necessarily downscaled because their resolution is not appropriate to make projections using the 14 PRMS applications. Downscaled GCM projections, at resolutions appropriate for the PRMS applications, have been made available by several research groups (for

example, http://gdo-dcp.ucllnl.org/downscaled_cmip3_projections/dcpInterface.html, accessed July 2010 or <http://www.gisclimatechange.org>, accessed July 2010). These products were considered for this study, but were rejected because the downscaled gridded datasets were not compatible with the previous studies (table 1). All of the PRMS models used in this study distribute precipitation and maximum and minimum temperature data from a set of points (station locations) to HRUs defined for each basin. In order to apply these gridded products, each PRMS model would have to have been reconfigured to use gridded information, changing each model considerably from its original configuration. Therefore, a downscaling procedure that goes from coarse GCM grids to the climate stations used in each PRMS model was desired.

Either statistical or dynamical methods can be used to downscale coarse-resolution GCM projections to the basin scale for the PRMS applications. Statistical downscaling methods (Wilks, 1995; Wilby and others, 1999; Antolik, 2000; Intergovernmental Panel on Climate Change, 2001; Intergovernmental Panel on Climate Change, 2007) use empirical relations between features which are reliably simulated at GCM grid scales, and surface predictands at subgrid scales. Dynamic downscaling methods use finer-scale regional climate models, with initial and lateral boundary conditions determined by the GCM projections, which result in detailed climate simulations over a smaller region of interest (Hay and Clark, 2003; Leung and others, 2003; Giorgi and others, 2001).

Whereas considerable research and development has gone into these downscaling procedures, the most straightforward means of obtaining higher spatial resolution climate projections is to apply the statistical change factor method (Arnell, 2003a; 2003b; Arnell and Reynard, 1996; Hay and others, 2000; Diaz-Nieto and Wilby, 2004; Eckhardt and Ulbrich, 2003; Pilling and Jones, 1999; Prudhomme and others, 2002; Hay and McCabe, 2010). Fowler and others (2007) reviewed downscaling methods for their applicability for use in the evaluation of hydrological effects. They concluded that simple statistical downscaling methods, such as the 'change factor' method, seem to perform as well as more sophisticated methods in reproducing mean climatic characteristics. One limitation of statistical downscaling is that the temporal

Table 4. Carbon emission scenarios simulated by the General Circulation Models in this study.

Emission scenario	Description/Assumptions
20C3M	20th century climate used to determine baseline (1989–1999) conditions
A1B	Rapid economic growth, a global population that peaks in mid-21st century and rapid introduction of new and more efficient technologies with a balanced emphasis on all energy sources
B1	Convergent world, with the same global population as Emission scenario A1B, but with more rapid changes in economic structures toward a service and information economy that is more ecologically friendly
A2	Heterogeneous world with high population growth, slow economic development, and slow technological change

sequencing remains unchanged, limiting this method to studies examining changes in mean climatic conditions. For example, if changes in wet-/dry-spell lengths are important to the effects assessment, then another downscaling choice would be more appropriate (Wilby and others, 2004).

In this study, the change factor method was applied to the coarse-scale GCM projections and the PRMS baseline conditions to generate an ensemble of PRMS Data Files which were

used to simulate the scenarios. The GCM simulations, of the 20C3M scenario, for the PRMS baseline period (WY 1988–1999) were used to compute the change factors (percentage changes in precipitation and degree changes in temperature). Change factors were computed for 12-year moving window periods (2001–2099) for the A2, B1, and A1B scenarios for each month. Figure 7 shows a schematic of the climate change factor method used in this study.

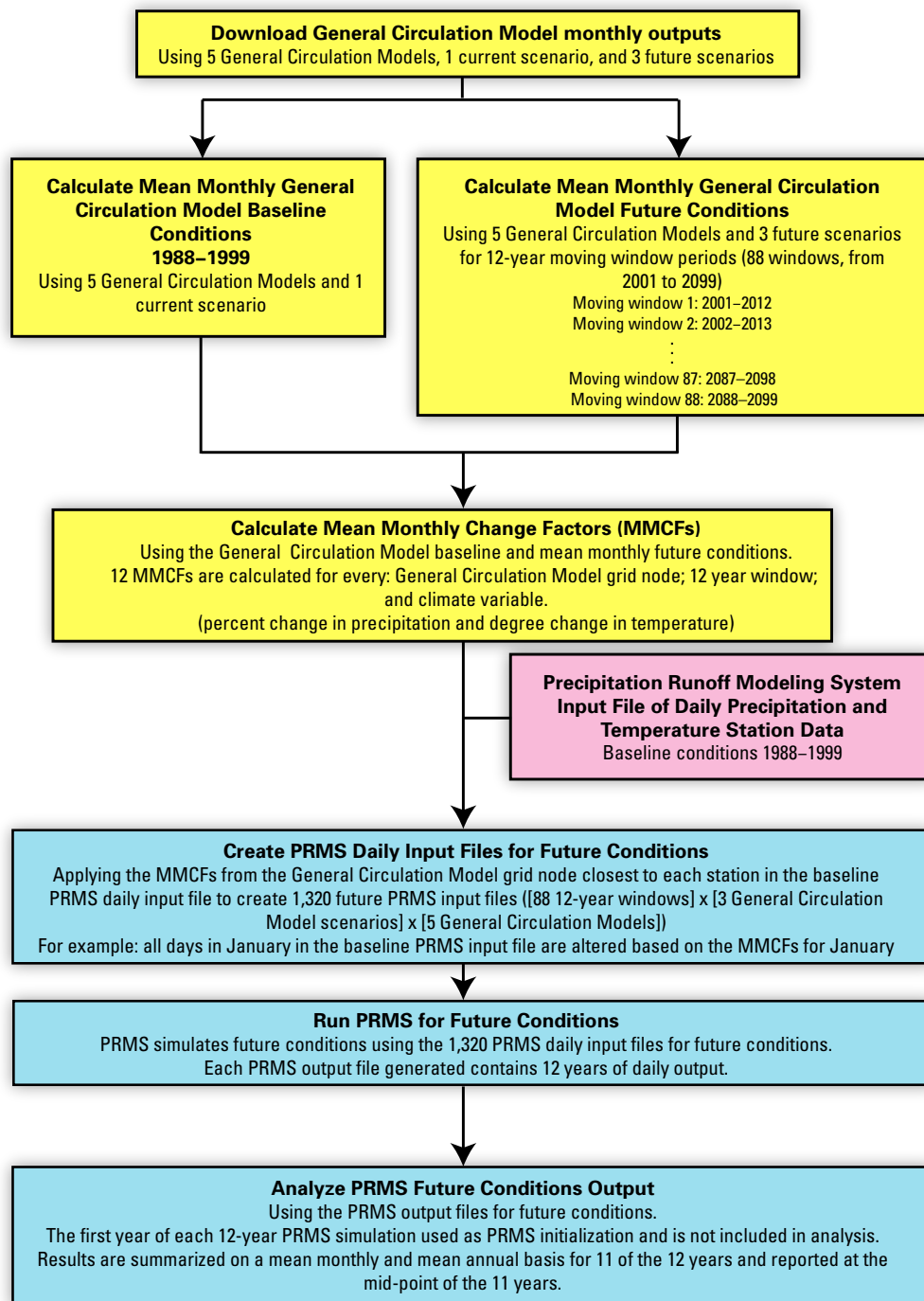


Figure 7. Schematic of the climate change factor method as applied in this study.

A software tool, called Shodan, automates the calculation of change factors and applies them to the PRMS Data Files (fig. 8). The downscaling process, as described above (this section), results in 1,320 PRMS Data Files ([88, 12-year climatologies, 1 per year starting with water years 2001–2012 and ending with water years 2088–2099] x [3 scenarios] x [5 GCMs]). Four steps are required to run Shodan. Each step and the respective user interface screen are described below:

The *Setup Parameters* screen (fig. 8A) allows specification of: (1) the Start Year and End Year of the Baseline

Conditions Period; (2) the Start Year and End Year of the Projection Period; (3) the Moving Window Length, which is the number of years for each projection; (4) the Baseline GCM Scenario; (5) the GCM Projection Scenarios; and (6) the GCM Models.

The *Setup Input* screen (fig. 8B) allows specification of: (1) the Input File Format; (2) the PRMS File name; and (3) the Meta Files, which contain information about the station point locations for the data (fig. 9).

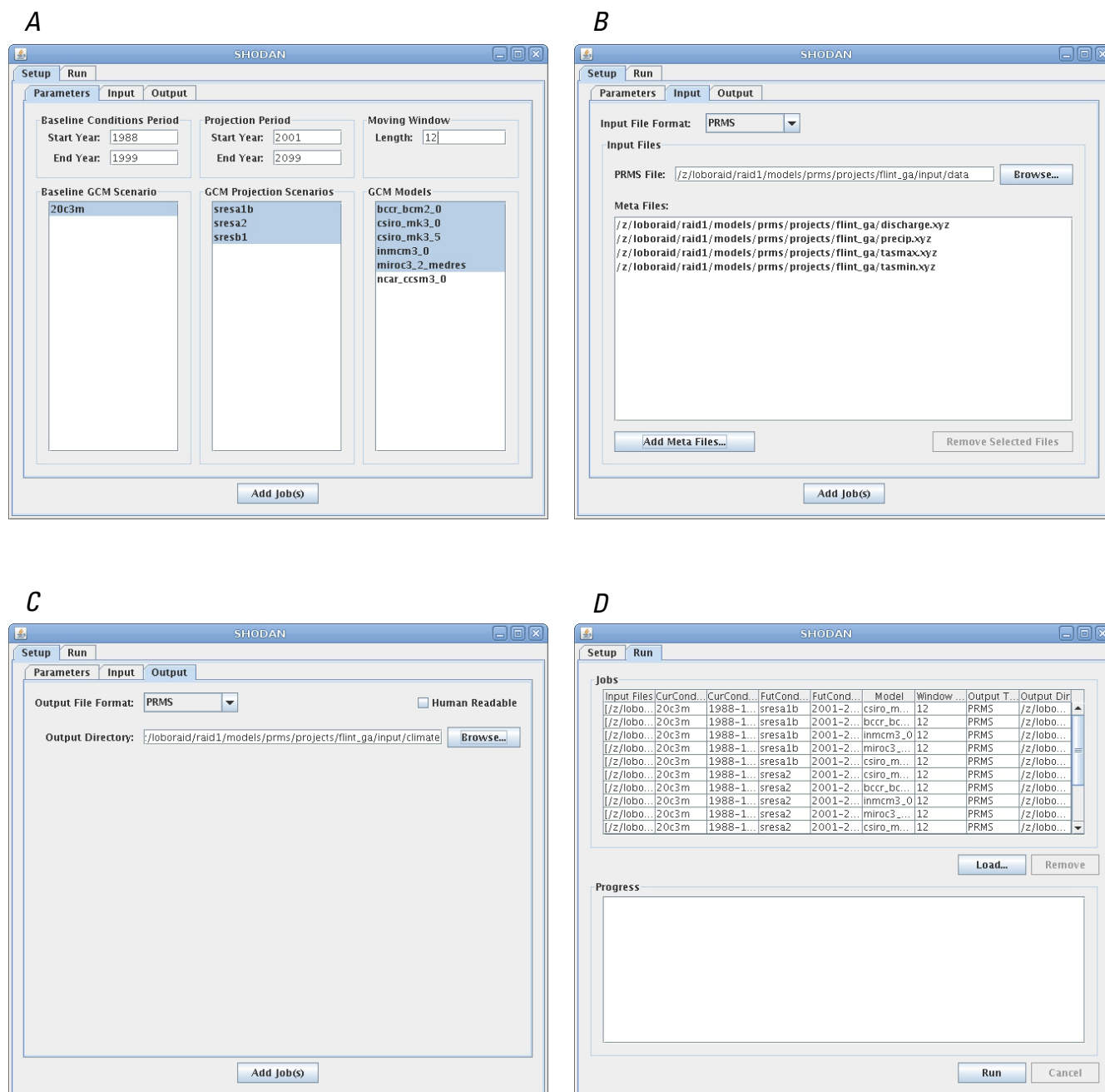


Figure 8. Shodan graphical user interface used to create Precipitation Runoff Modeling System input data files for future climatic conditions: (A) Setup Parameters, (B) Setup Input, (C) Setup Output, and (D) Run windows.

Number of stations:	5				
Variable:	tasmax				
Unit:	fahrenheit				
ID:	172620	174878	172426	176430	179891
Network:	COOP	COOP	COOP	COOP	COOP
Name:	"ELLSWORTH"	"MACHIAS"	"EASTPORT"	"ORONO"	"WOODLAND"
Latitude(degree_N):	44.5333	44.7194	44.9067	44.8992	45.1572
Longitude(degree_E):	-68.4333	-67.4542	-66.9919	-68.6744	-67.4044
Elevation(meter):	6.096	6.096	25.908	35.052	42.672

Figure 9. An example of the tasmax_xyz file of station point locations used by Shodan.

The *Setup Output* screen (fig. 8C) allows specification of: (1) the Output File Format; and (2) the Output Directory where the downscaled PRMS Data Files will be written.

The *Run* screen (fig. 8D) shows: (1) the downscaling Jobs that will be performed by Shodan. Each job consists of the creation of a set of PRMS Data Files, each containing the downscaled climate information; (2) the Run Jobs button which starts the creation of the PRMS Data File; and (3) the status of the jobs in the Progress window.

Step 4: Generate the Precipitation Runoff Modeling System Projections

For each of the 1,320 PRMS Data Files created by Shodan, as described in Step 3: Downscale General Circulation Model Projections of this report, the following procedure was executed:

- Modify the PRMS Control File so that the correct Shodan-produced PRMS Data File will be used.
- Modify the PRMS Control File so that the names of the four PRMS output files will be unique for each scenario. These output files are the PRMS statvar file, the HRU animation file, the Groundwater Reservoir animation file, and the Subsurface Reservoir animation file.
- Modify the PRMS Control File so that the starting and ending time step of the simulation is the appropriate 12-year time window.
- Determine the period of active transpiration for each HRU for the PRMS projection 12-year time window by executing the PRMS model in preprocess mode by specifying the “-preprocess” command line argument.

The period of active transpiration is computed by the new PRMS module “frost_date_prms” which processes the daily HRU temperatures and appends the parameters “fall_frost” and “spring_frost” directly to the PRMS Parameter File. These parameters are used by the new PRMS module “potet_jh_prms.” Both new modules are documented in Appendix 1, “Documentation of New PRMS Modules.”

- Simulate the 12-year time window by executing the PRMS model with the input and output files as specified above in section Step 3: Downscale General Circulation Model Projections.

Step 5: Analyze the Precipitation Runoff Modeling System Projections

This study is limited to “climate change;” therefore, projected changes in the 21st century climate were analyzed on a basin mean monthly basis and on an annual basis for every PRMS model output variable in table 3. The PRMS projections are summarized for each of the basins in the “PRMS Projections by Application” section below.

Assumptions and Limitations

Several assumptions and limitations related to the data and modeling aspects of this study are discussed. These are (1) the spatial and temporal resolution of the data and how it relates to the PRMS model, (2) the historical period used to represent the 20th century, (3) estimation of solar radiation by the PRMS model, (4) PRMS simulation and accounting of snowpacks, (5) PRMS simulation of evapotranspiration, and (6) a general discussion of the different sources of uncertainty.

Spatial and Temporal Resolution

Because the PRMS model has a time step of one day, all flows and storages are mean daily values. In general, flows near land surface occur more rapidly than flows in the sub-surface. Consequently, the daily time step may result in errors because of temporal averaging for near land-surface flows such as surface runoff, infiltration, and interflow. Simulation errors can occur when streamflow changes abruptly over time increments that are less than one day. For this study, simulating the subdaily behavior of these processes is not warranted because all output is analyzed on a mean monthly basis. Basically, it is assumed that the within day errors that might result do not significantly affect the monthly water budget analyses.

The delineation of the HRUs must be appropriate for the PRMS algorithms. In the PRMS model, each HRU is assumed to be homogeneous with respect to both input parameterization and computation of flows and storages. Because the 14 PRMS applications were developed by different modelers for different purposes, HRU delineation methodologies were not always consistent. Some of the 14 basins were set up as medium-to-long-term water-supply studies, which results in large HRUs, while others were set up as the basis for ecological and habitat studies, which results in much finer resolution HRUs. This problem was overcome by analyzing the output on an area-weighted average basis.

Historical Period Used as Baseline Conditions

The historical climate record chosen to represent baseline conditions may affect the PRMS projections. Kocot and others (2011) found when using the change-factor method for downscaling, several baseline conditions may need to be chosen to represent a range in historical climatic conditions. Ideally trends, variability, and extremes in the climate record could be considered. Kocot and others (2011) also found that not all hydrologic variables equally were sensitive to the selection of baseline conditions. Projections of PRMS variables which are not associated directly with long-term water storage in the model, like temperature, tend to be dominated by the climate trends introduced by the scenario. Variables associated with interannual water storage in the model, like content of the soil-zone reservoir, appear to be more sensitive to the selection of baseline conditions. In addition, the geographic location and dominant physical processes (in this case, influences from topography and multidecadal climatic cycles) may have a strong bearing on the sensitivity of PRMS projections to the selection of baseline conditions.

Simulation of Solar Radiation

The PRMS model simulates daily solar radiation for each HRU of the 14 applications in this study. Simulated solar radiation is used for snowpack energy balance calculations, and to simulate PET. These solar radiation values are simulated by

one of two methods, depending on the geographical location of the basin, as described by Markstrom and others (2008, p. 41–43). Both of these methods heavily rely on an assumed positive relation between daily temperature and daily solar radiation. Neither method accounts for variations in solar activity or changes in atmospheric effects. Under conditions of changing climate, these assumptions may not hold. Roderick and others (2009) report that in some locations, despite a positive trend in daily air temperature, there may be a dimming of solar radiation. The PRMS model does not account for this. In future studies, the PRMS model could use solar radiation change factors computed from the GCM projections; however, uncertainty associated with this may not justify the additional effort.

Notwithstanding these issues related to simulation of solar radiation, analyses of the PRMS projections indicate that the simulated changes in solar radiation typically are less than 5 percent over the period of the study. Consequently, this small change in projected solar radiation has a minimal effect on projected basin runoff and other hydrologic variables.

Simulation of Snowmelt and Rainfall on Snowpacks

PRMS applications in cold weather and mountainous basins are sensitive to snowpack simulation. The PRMS model simulates temperature, melting, water content, and temporal and areal extent of the snowpack. It also determines phenological occurrences, such as the initiation of the snowpack and the onset of spring. Simulation of these processes is described in detail by Markstrom and others (2008), and the effects on this study are described below in the “PRMS Projections by Application” section.

Water storage within the snowpack is tracked in two states: ice (solid) and free water (liquid). The amount of free water in the snowpack is determined by freeze-thaw cycles, based on energy balance, and the physical capacity of the snowpack matrix to hold free water. The total amount of water storage in a snowpack is expressed as the volume of water equivalent, neglecting air space, if all the ice is melted to liquid form.

The PRMS model uses the freezing point of water (an all-ice isothermal snowpack at 0 °C) as the condition to initiate melt in the snowpack. When the snowpack energy rises above this state, some ice melts and liquid free water results. The PRMS model reports snowmelt only when the free water in storage exceeds the pore space of the snowpack. This loss of free water from the snowpack represents an output of both water and latent energy.

Complications occur when falling rain enters the snowpack as free water in excess of the pore space. The PRMS model will report this excess water, which has never been frozen, as snowmelt. Another complication occurs when falling rain enters the snowpack as free water in deficit of the pore space and freezes to ice because of an energy deficit. In this

case, the snowpack may contain more ice than the amount of precipitation that fell as snow, eventually causing the PRMS model to report more snowmelt than snowfall. Either of these cases may complicate the interpretation of simulations, particularly when the study basin straddles the seasonal snow line.

Evapotranspiration

A major component in the simulation of the water balance for any basin is the actual evapotranspiration (AET) rate. The PRMS model uses many factors to compute AET, including the simulated PET, vegetation type and cover density, land-use characteristics, soil type, simulated atmospheric conditions (available energy, precipitation, and cloud cover), and soil moisture availability (Markstrom and others, 2008).

An important source of uncertainty in hydrologic simulation modeling is the methodology and available data used to compute PET (Kingston and others, 2009). In the PRMS model, PET is the theoretical absolute upper limit of simulated AET and is not used directly in water balance computation. Measured, or surrogates for measured, PET data can be useful for calibration and evaluation of PRMS applications because they are independent of measured streamflow. Typically, this information is available either as an atlas of stationary mean monthly values (such as Farnsworth and others, 1982) or as time series of pan evaporation rates or corrected lake levels at sites in or nearby the study basin. Research suggests that the use of historical PET information must be considered carefully as rates have demonstratively changed over the latter half of the 20th century (Walter and others, 2004; Roderick and others, 2009; and Szilagyi and others, 2001). The physical process mechanism responsible for non-stationarity in PET is not well understood and PRMS provides no algorithmic simulation capability for it.

This study uses the Jensen-Haise method (Jensen and others, 1970; and Markstrom and others, 2008) to estimate PET. The method uses empirical constants to define a long-term climatic setting for simulating daily PET. The variability of daily PET is determined by the daily variation in solar radiation (as discussed in the “Simulation of Solar Radiation” section of this report) and temperature of the evaporation surfaces.

Two criticisms of the Jensen-Haise method are that it does not explicitly account for daily variation of relative humidity and assumes complete mixing of the atmosphere above the evaporation surface. Consequently, this method does not explicitly account for variation in the near-surface vapor pressure gradient and mass transfer of vapor into the atmosphere. In addition, because the method, as implemented in the PRMS model, relies on near surface air temperature as an index for both sensible heat and solar radiation, under warming conditions, it may overestimate the amount of energy available for PET. Milly and Dunne (2011) report, that for some of the applications in this study, the PRMS model

overestimates PET. Conceptually, overestimation of PET may lead to unrealistically large decreases in projected streamflow; however, because AET, rather than PET is used in the water balance computations, decreases in streamflow are not nearly so severe (Kingston and others, 2009; and Donohue and others, 2010). In fact, all of the basins in this study, except one, are soil moisture limited (Hay and others, 2011).

Some researchers (Hall and others, 2008; Szilagyi, 2001; Walter and others, 2004; and Roderick and others, 2009) have observed that AET rates may be changing differently than PET rates. The “complementary hypothesis” (Szilagyi, 2001) describes how increases in AET may result in decreases in the vapor pressure deficit which consequently results in decreases in pan evaporation (PET). Roderick and others (2009) report that this phenomenon may not occur universally in all locations, as other physical processes, such as dimming of solar radiation and decreased wind speed may be dominate. Regardless, the PRMS model does not algorithmically account for any of these processes. Changes in the simulated AET rates, from the PRMS baseline simulations (Hay and others, 2011), are comparable to measured (or estimates of measured) values presented in the literature (Walter and others, 2004, p. 204; and Szilagyi and others, 2001). This may give an added degree of confidence that the PRMS projections are reasonable.

Under extreme changing climate conditions, the PRMS computational PET assumptions may not be adequate; however, the Jensen-Haise method has been successfully applied in widely varying environments under current and recent past conditions (later half of the 20th century), which suggests its applicability in the near future (Hay and others, 2011). Donohue and others (2010) believe that accurate simulation of PET is dependent on knowledge of net radiation, vapor pressure, wind speed, and air temperature. Ideally, accurate measurements (or projections) of these data and an appropriate PET simulation algorithm would result in improved simulation and reduction of uncertainty.

Uncertainty

Numerous sources of uncertainty have been identified in the steps of this study. Large uncertainties are associated with the representation of the physical processes, model structure, and feedbacks within the climate system as projected by the GCMs (Intergovernmental Panel on Climate Change, 2007). The scenarios chosen for this study (table 4) represent different economic, social, political, and technological development for the future, none of which may be the actual path. In this study, multiple GCM projections (table 2) were used to quantify these uncertainties. While some researchers prefer to weight or select specific GCMs based on their ability to reproduce current climate, others (Murphy and others, 2004; Tebaldi and others, 2005; Stainforth and others, 2007) argue that the entire ensemble of GCM projections can be used to produce a lower bound on the maximum range of uncertainty.

Better quantification of uncertainty in GCM projections may be achieved by combining Bayesian approaches (Draper 1995; Tebaldi and others, 2005; Raftery and others, 2005) with earth models of intermediate complexity or energy balance models (Urban and Keller, 2009). This approach will allow detailed evaluation of the tails of the climate sensitivity distribution and can be used to derive probabilistic projections or posterior model weights, reducing the risk of overconfident projections that can result from truncated estimates of GCM model uncertainty within a finite ensemble of GCM projections.

The spatial scale of the GCM projections is an additional source of uncertainty. The change-factor method was used to downscale the coarse resolution GCM projections to the finer resolution input required by the PRMS model. In this approach, the changes in atmospheric processes at the GCM subgrid scale are not resolved, such as changing interactions between air masses and topography, and these subgrid interactions may be substantial in some cases (Ikeda and others, 2010). The uncertainty from the downscaling procedure may, however, be overwhelmed by the uncertainty of the driving GCM (Fowler and others, 2007; Khan and others, 2006), which has been shown to be consistently greater than uncertainty from the hydrologic model or natural variability (Prudhomme and Davies, 2009). Hydrologic model output was analyzed on a mean-monthly and annual basis. The uncertainty associated with this simple approach to downscaling is not directly addressed, but simple statistical downscaling methods seem to perform as well as more sophisticated methods in reproducing mean characteristics (Fowler and others, 2007).

The uncertainty going from the GCM projection to the hydrologic model projection can propagate in ways that may either be enhanced or compensated for depending on the structure and parameterization of the hydrologic model (Buytaert and others, 2009). There are many areas for hydrologic model improvement. Research towards improved process representation and methods to represent uncertainty in the choice of model structure and model parameters are needed. The approach used in this study ignores the important feedback mechanisms between the hydrology and the climate. The calibration of the hydrologic model assumes that the parameters for baseline conditions will hold under future conditions, but calibration to past data may not always fully represent the future (Blöschl and Montanari, 2010). The use of PRMS applications developed for other purposes, but used in an integrated assessment, leads to a bigger concern of model consistency in terms of formulation, parameterization, and calibration. As noted by Blöschl and Montanari (2010), when two groups examine the same watershed for effects of climate change the

projections can be different. The same can be said for inter-comparison using the same model (PRMS in this case) formulated for different reasons by different sets of researchers.

Precipitation Runoff Modeling System Projections by Application

The description of PRMS projections for each of the 14 application basins starts with a brief overview of the study area, an assessment of baseline conditions, and background on the previous PRMS application. This is followed by four sections: (1) Input: describes annual mean and mean monthly plot for precipitation, maximum temperature, and minimum temperature. Evapotranspiration and length of growing season (annual only) are also shown in this section; (2) Streamflow: describes annual mean and mean monthly plot for streamflow and the components of flow (surface, subsurface, and groundwater); (3) Snow: describes annual mean and mean monthly plot for the percent of precipitation that falls as snow, snow-covered area, snowpack water equivalent, and snowmelt; and (4) Soil Moisture: describes annual mean and mean monthly plot for the infiltration of rainfall and snowmelt, inflow to the capillary and gravity reservoirs, inflow to the groundwater reservoirs (recharge), soil moisture in the capillary reservoir (water stored by soil tension), subsurface storage reservoirs, and groundwater storage reservoirs. A description of all the variables can be found in table 3.

Annual Mean Plots

All annual mean plots (figs. 10–28) show the range in the 11-year moving mean daily values by scenario. The first year of each 12-year window (as described above in the “Step 3: Downscale General Circulation Model Projections” section of this report) was used for initialization of the PRMS projection and is not included in analysis of results. The three solid colored lines indicate the 11-year moving mean daily values (x-axis indicates center of 11-year window) for the 3 scenarios (central tendencies of the 5 GCM projections for A2, A1B, and B1). The mean-annual plots show the range in the 11-year moving mean daily values by scenario. The red, blue, and yellow represent the A2, A1B, and B1 scenarios, respectively. The shaded areas shown for each scenario indicates an envelope (or range) of PRMS projections. Each of these plots is described, basin by basin, in the following sections of this report.

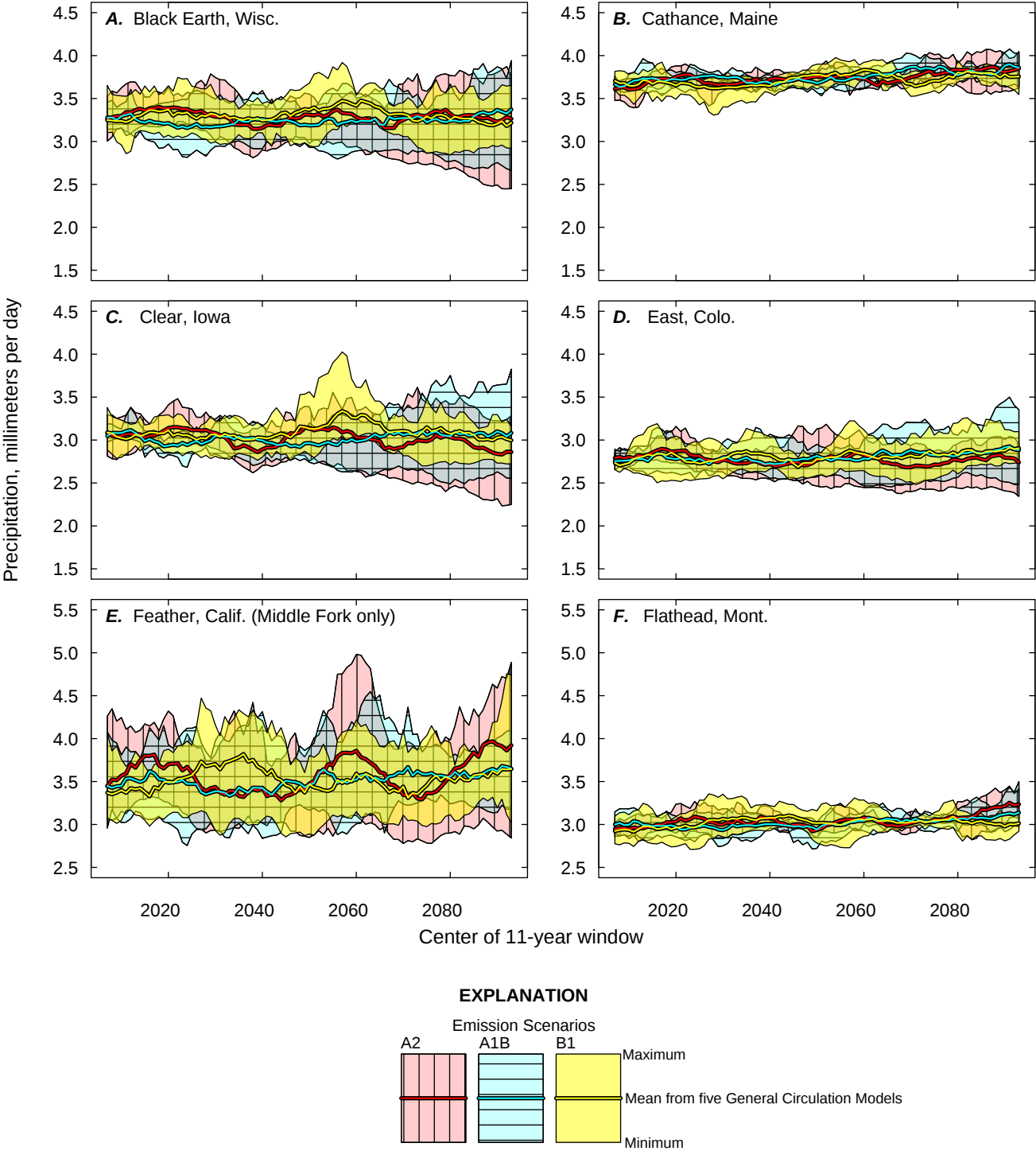


Figure 10. Projected range in 11-year moving mean daily precipitation values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

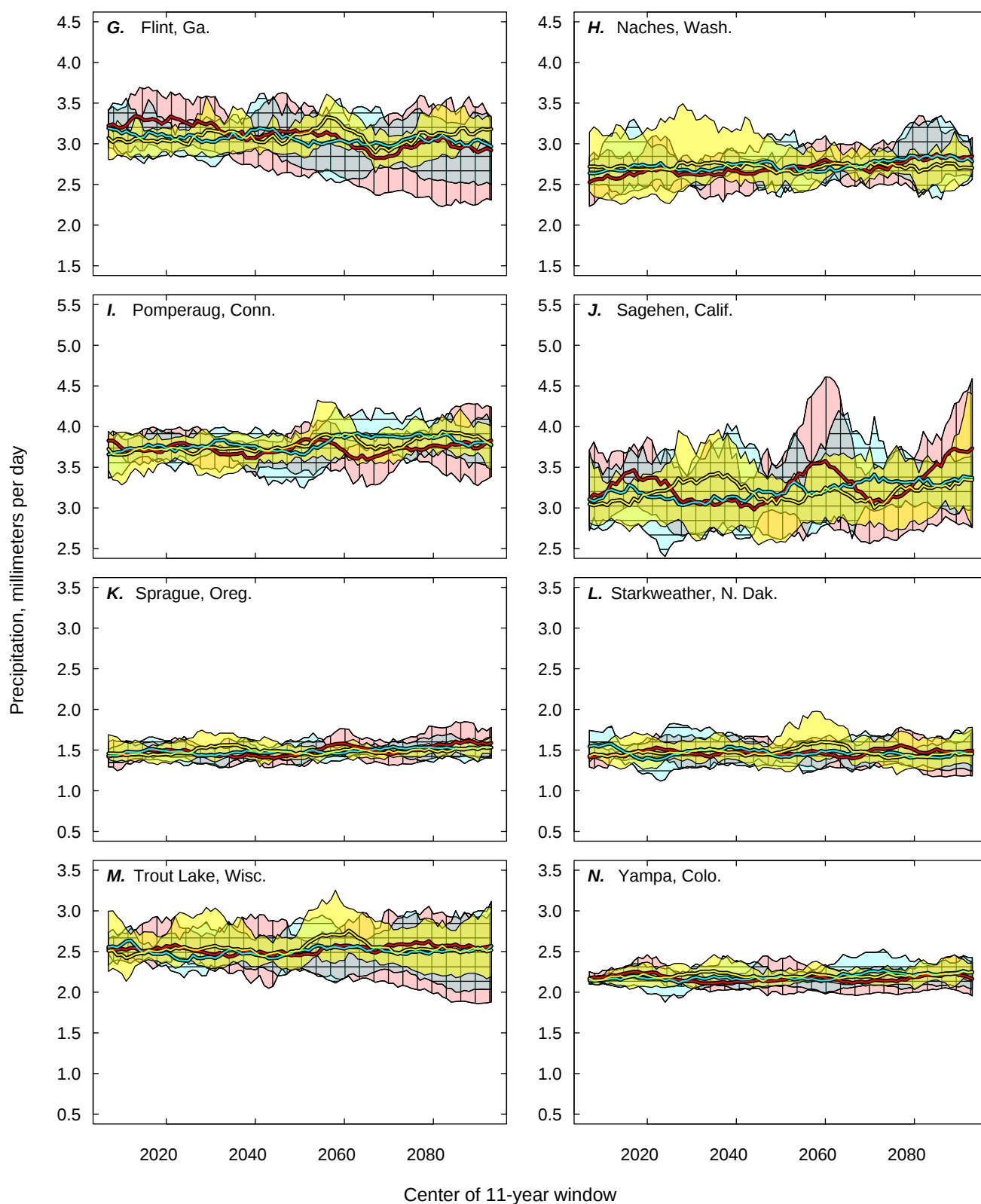


Figure 10. Projected range in 11-year moving mean daily precipitation values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

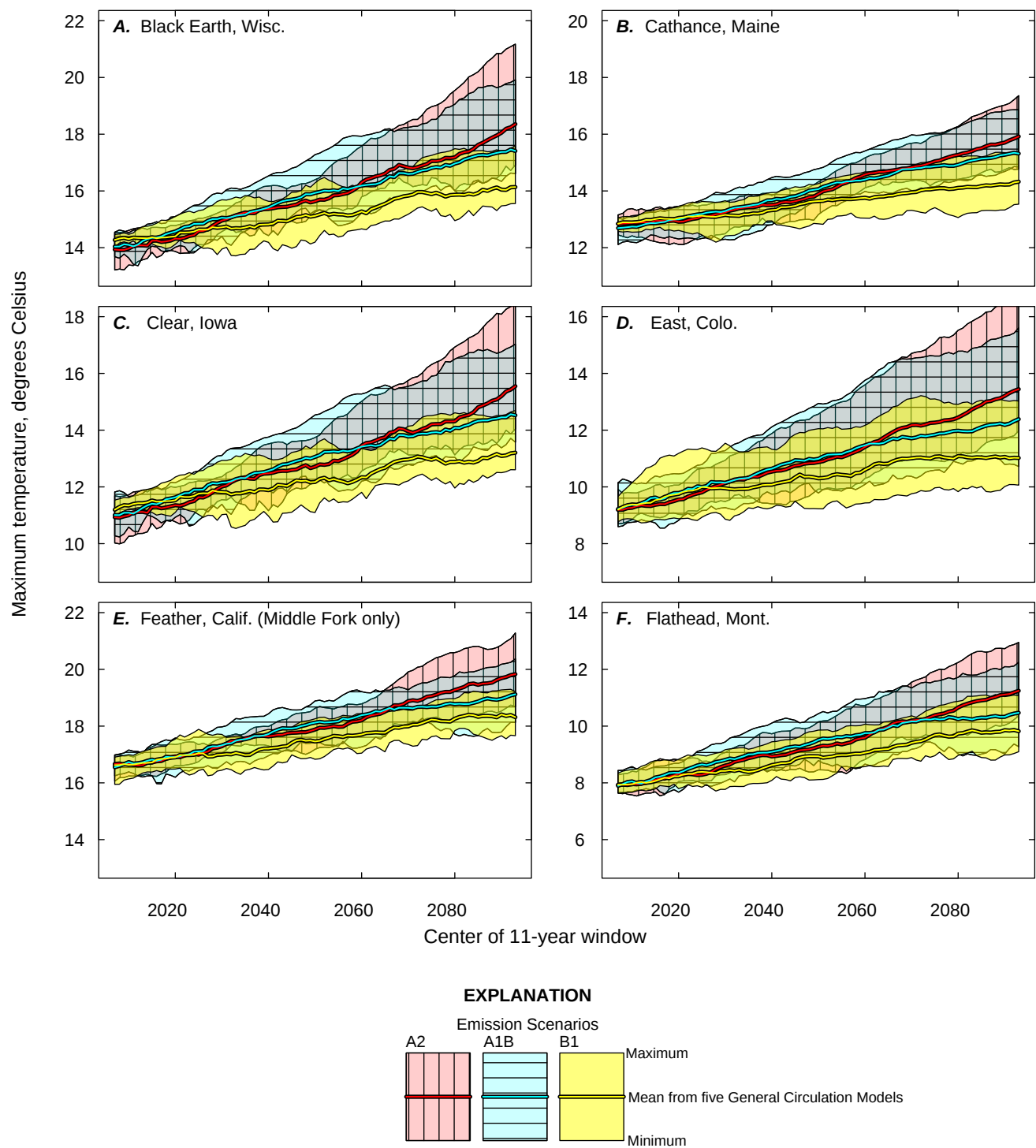


Figure 11. Projected range in 11-year moving mean daily maximum temperature values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

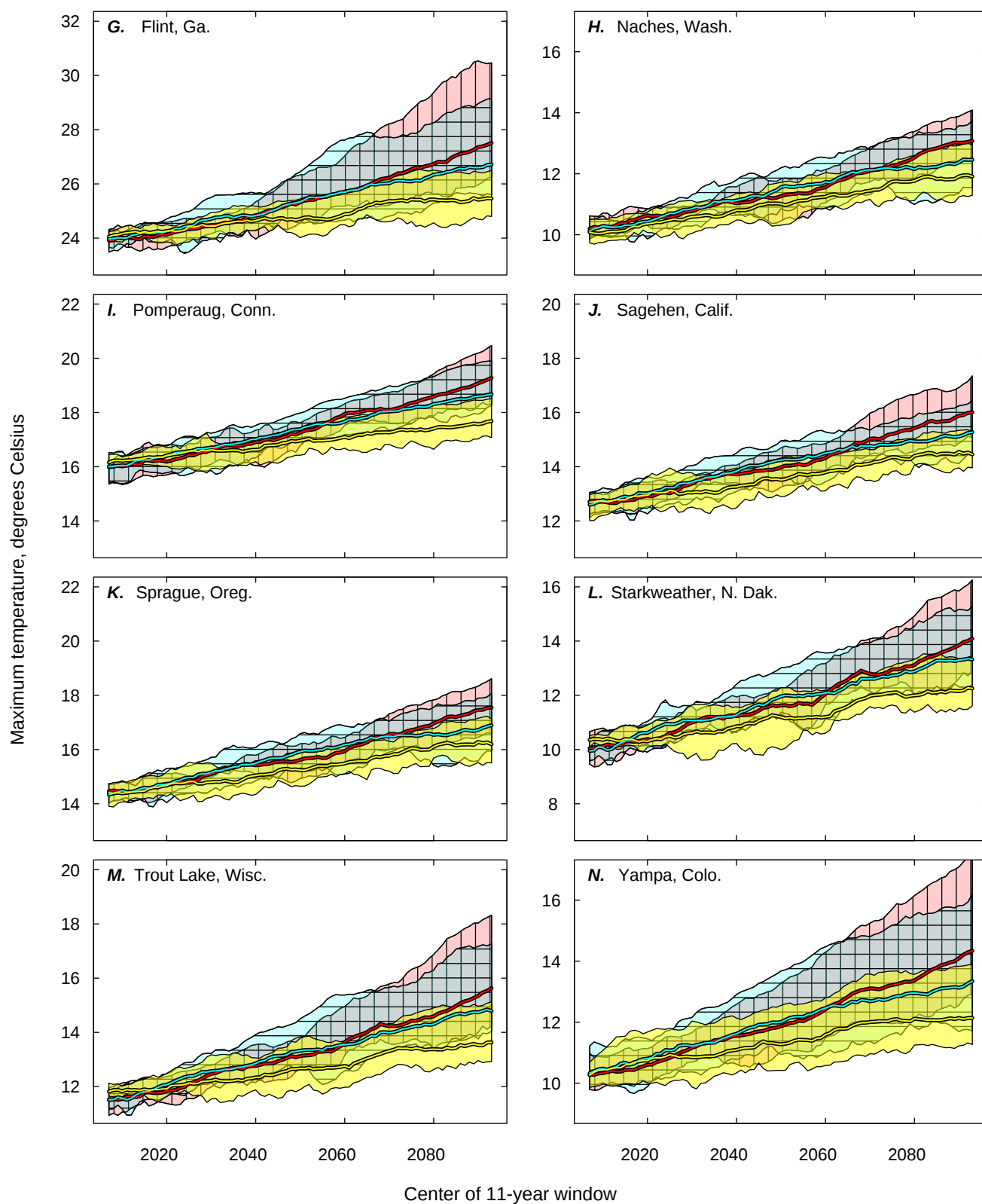


Figure 11. Projected range in 11-year moving mean daily maximum temperature values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

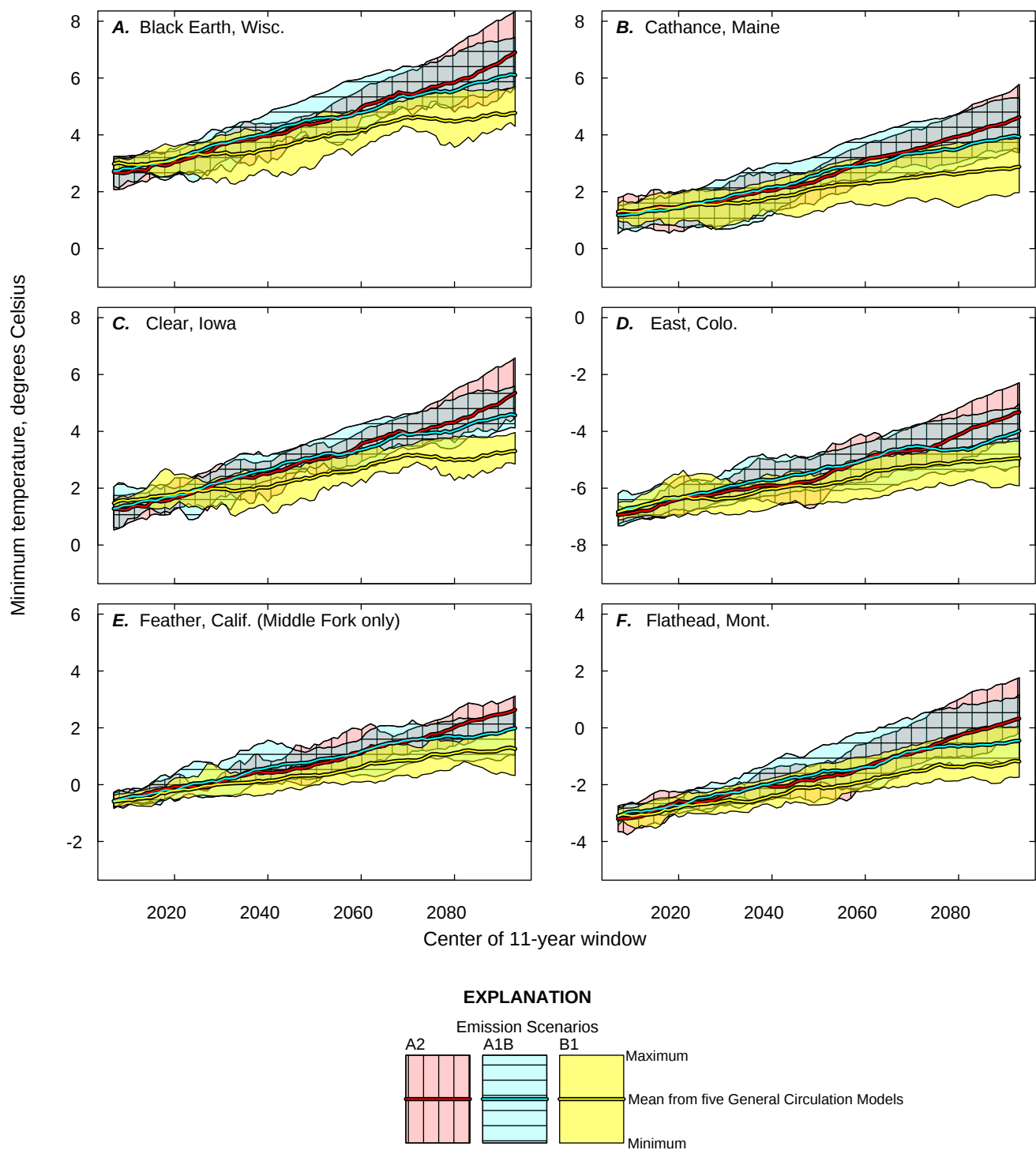


Figure 12. Projected range in 11-year moving mean daily minimum temperature values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

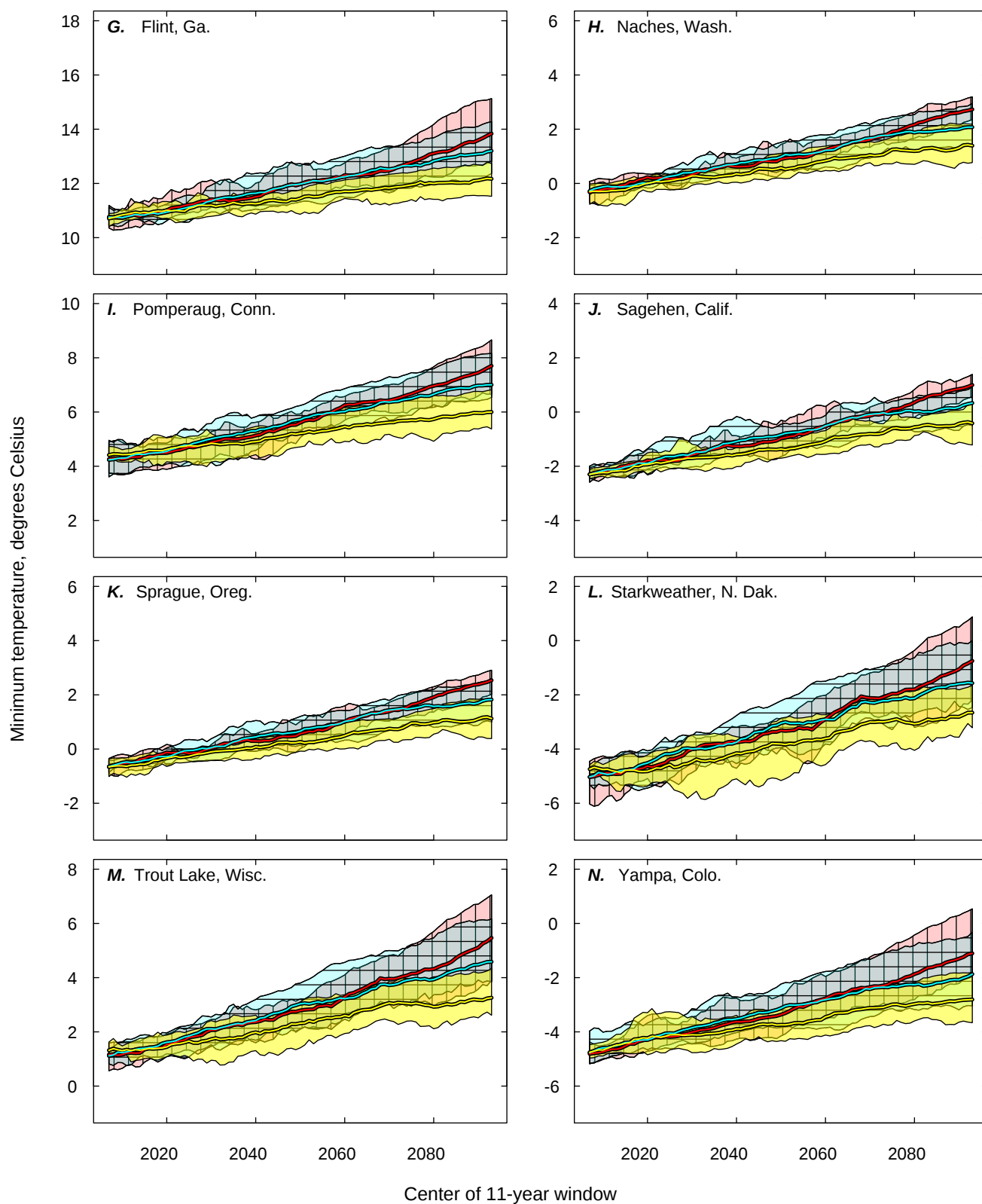


Figure 12. Projected range in 11-year moving mean daily minimum temperature values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

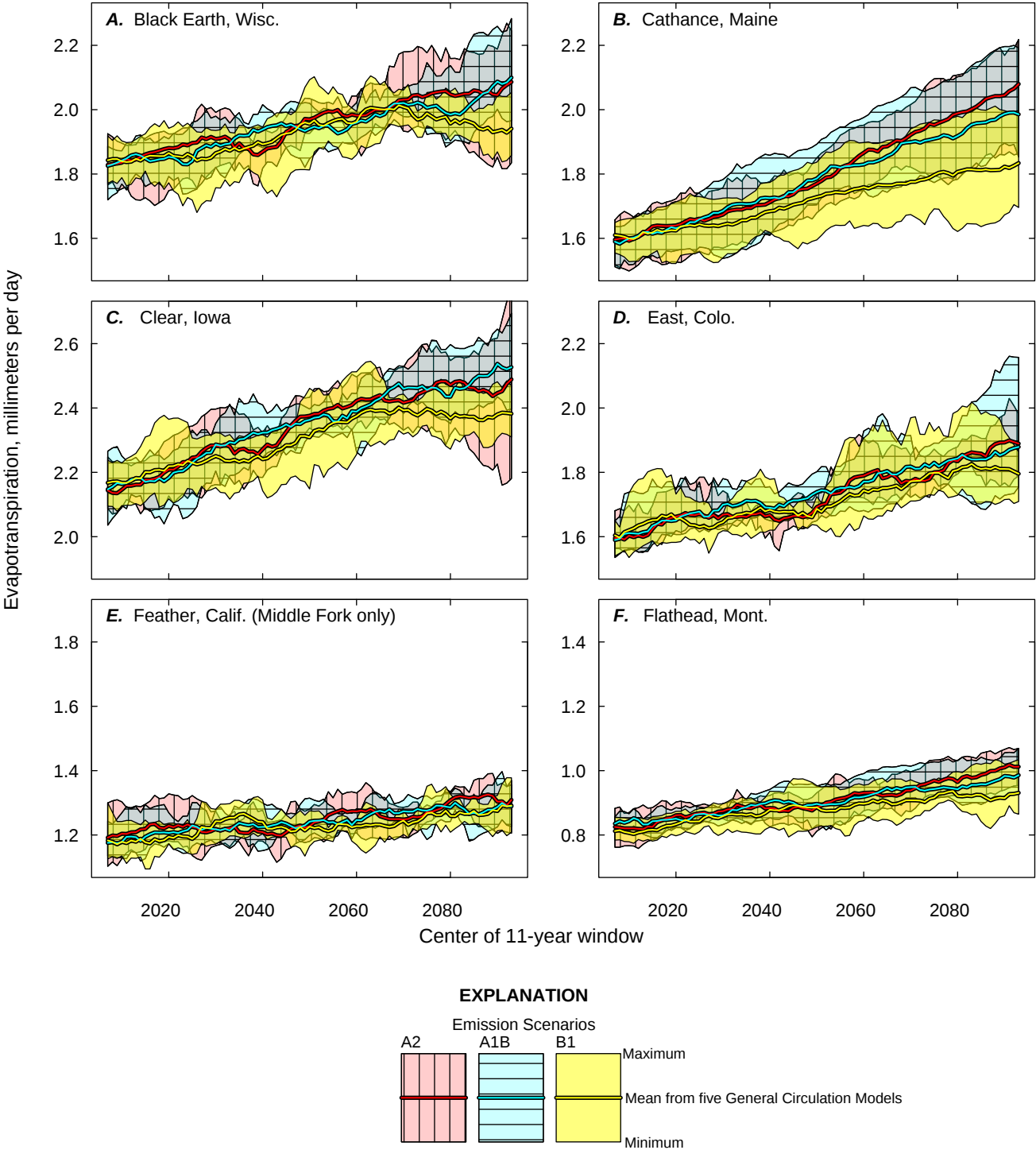


Figure 13. Projected range in 11-year moving mean evapotranspiration by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

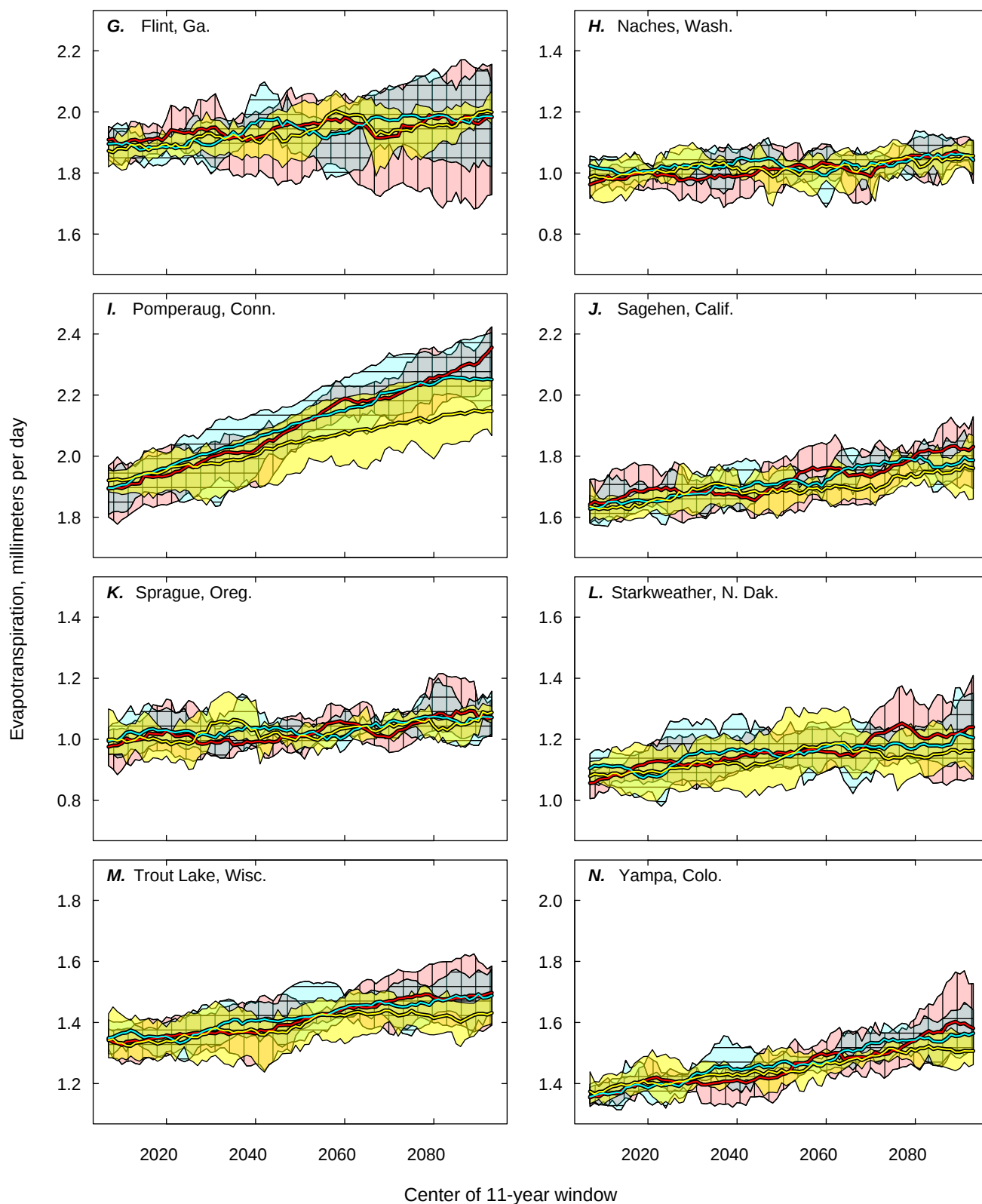


Figure 13. Projected range in 11-year moving mean evapotranspiration by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

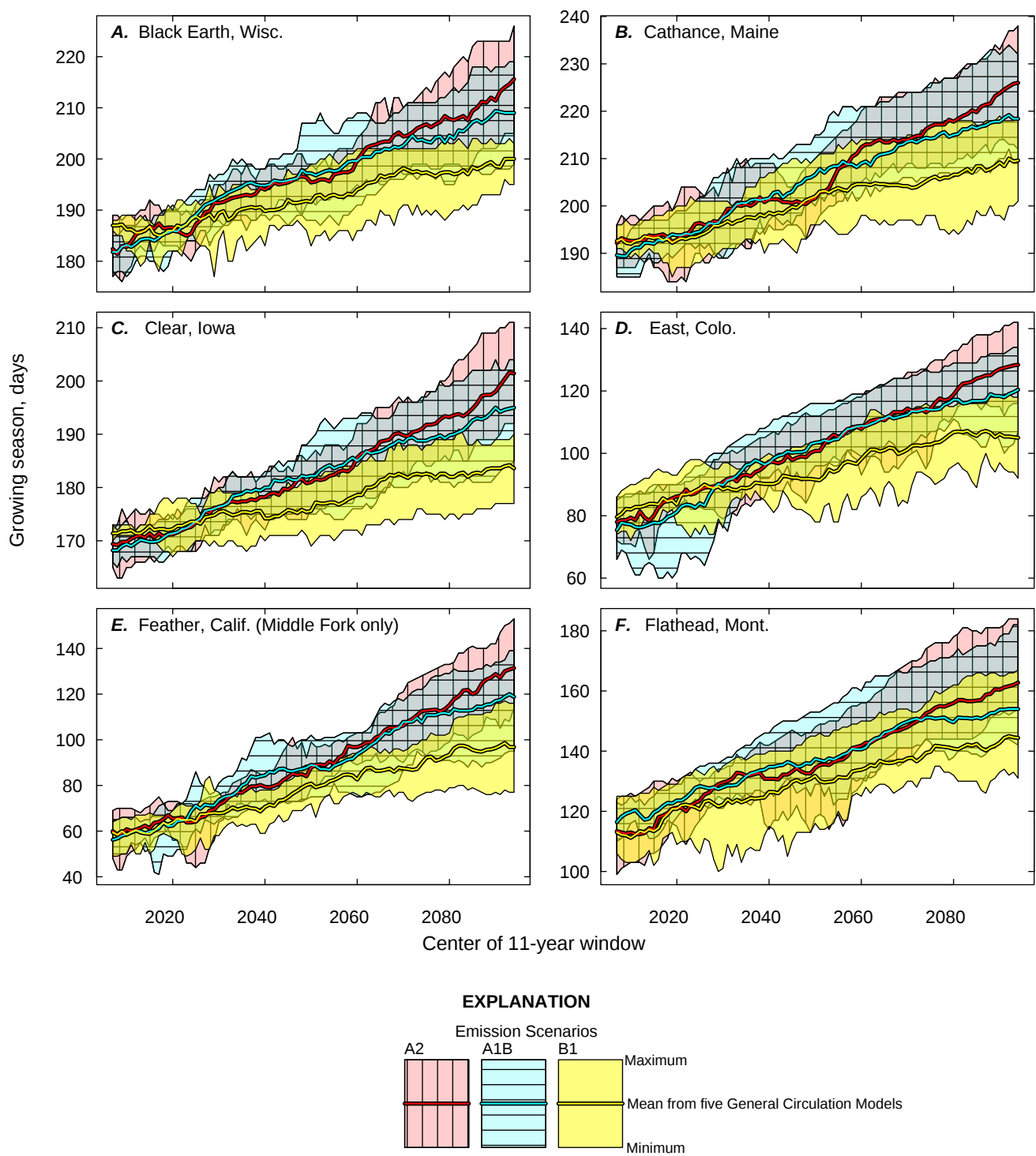


Figure 14. Projected range in 11-year moving mean growing season length by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

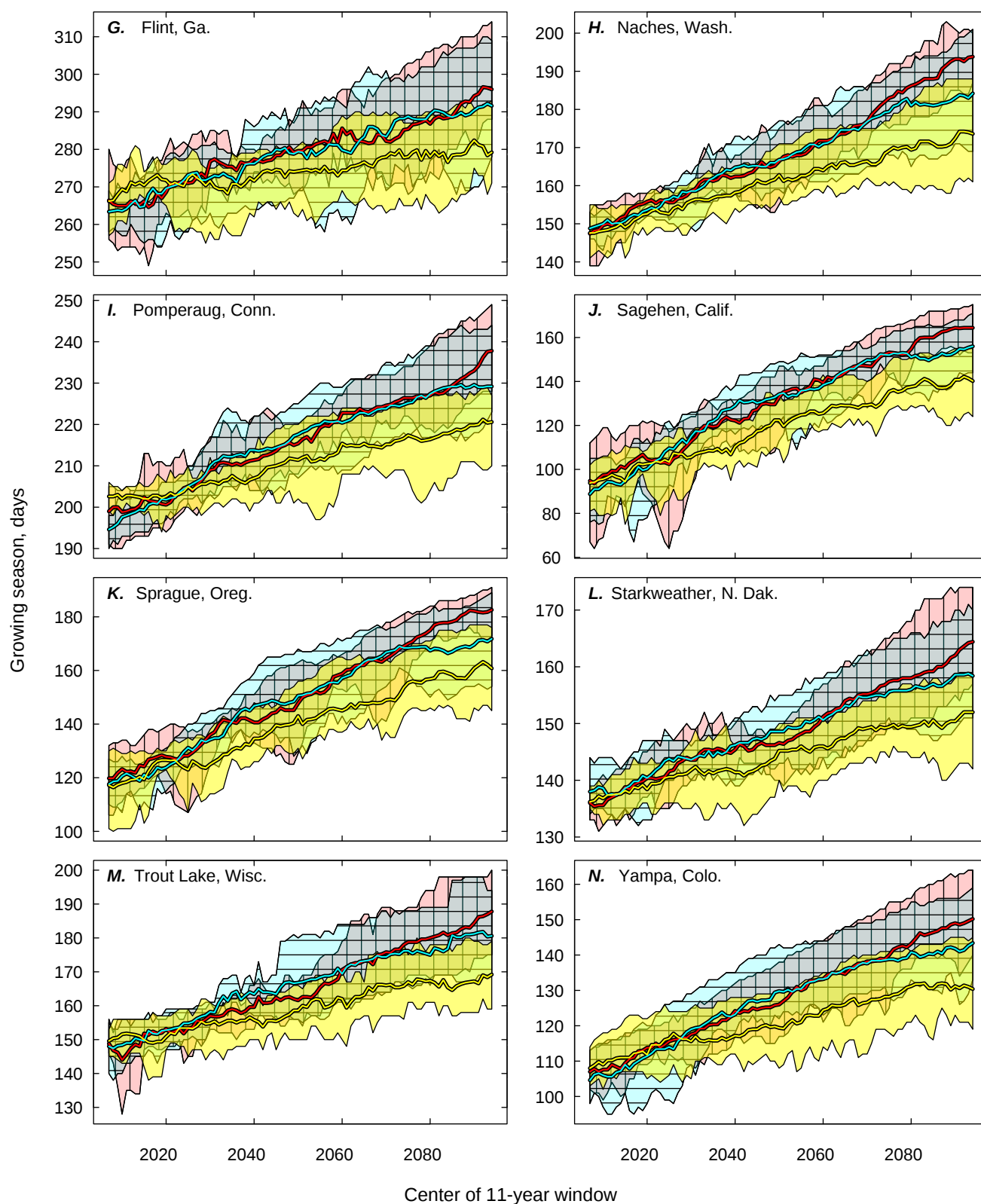


Figure 14. Projected range in 11-year moving mean growing season length by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

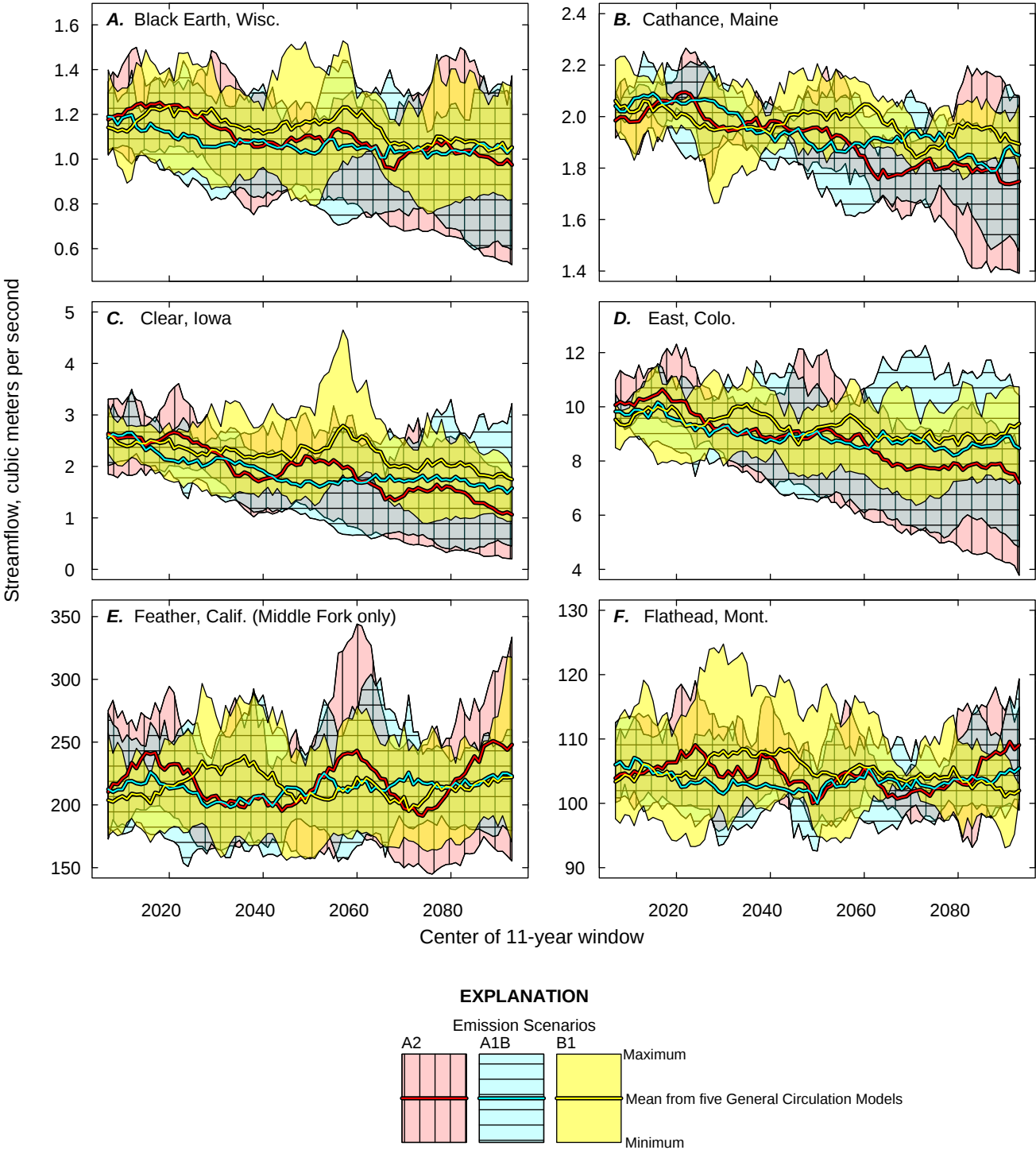


Figure 15. Projected range in 11-year moving mean daily streamflow values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

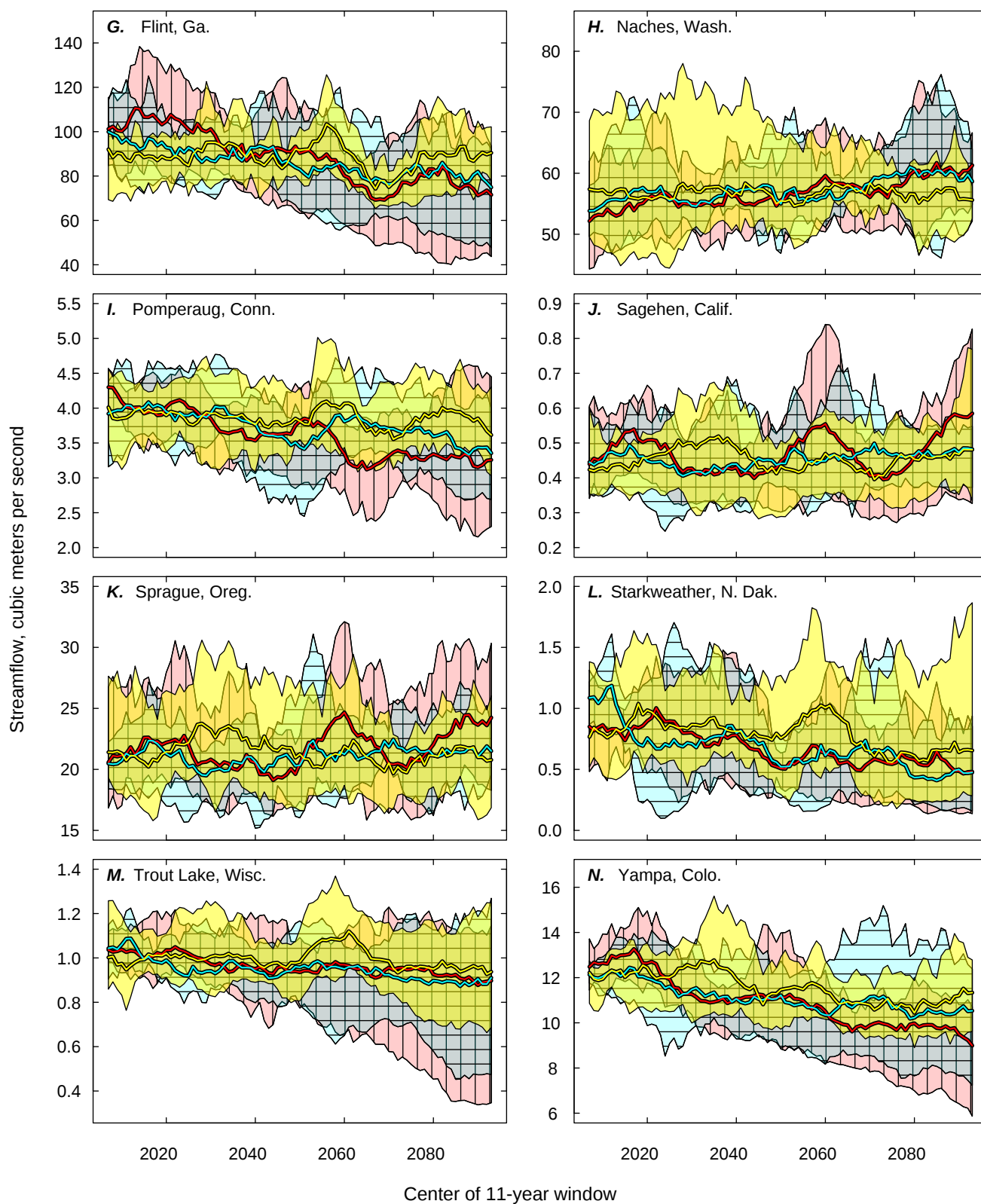


Figure 15. Projected range in 11-year moving mean daily streamflow values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

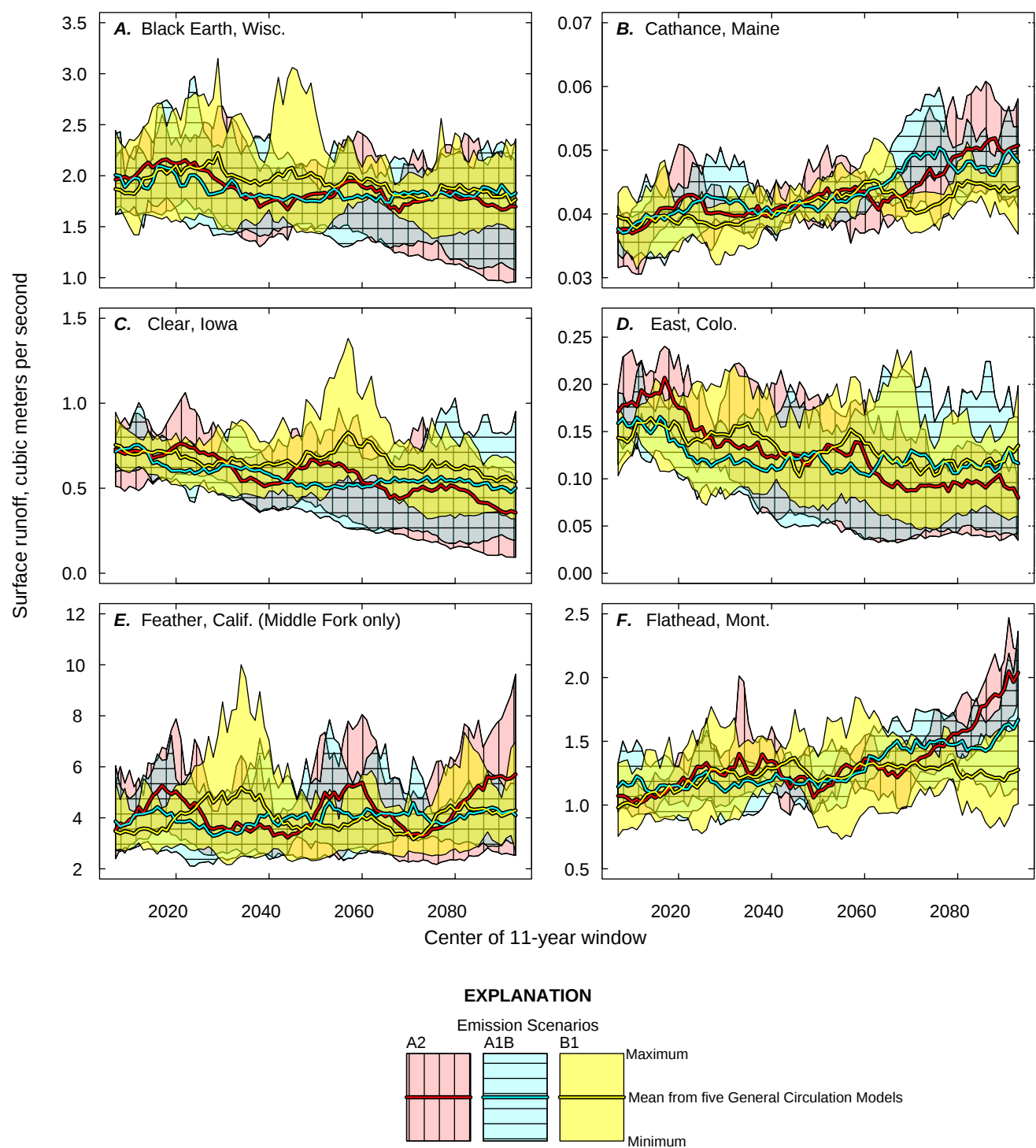


Figure 16. Projected range in 11-year moving mean daily surface runoff values by emission scenarios for the 14 basins.

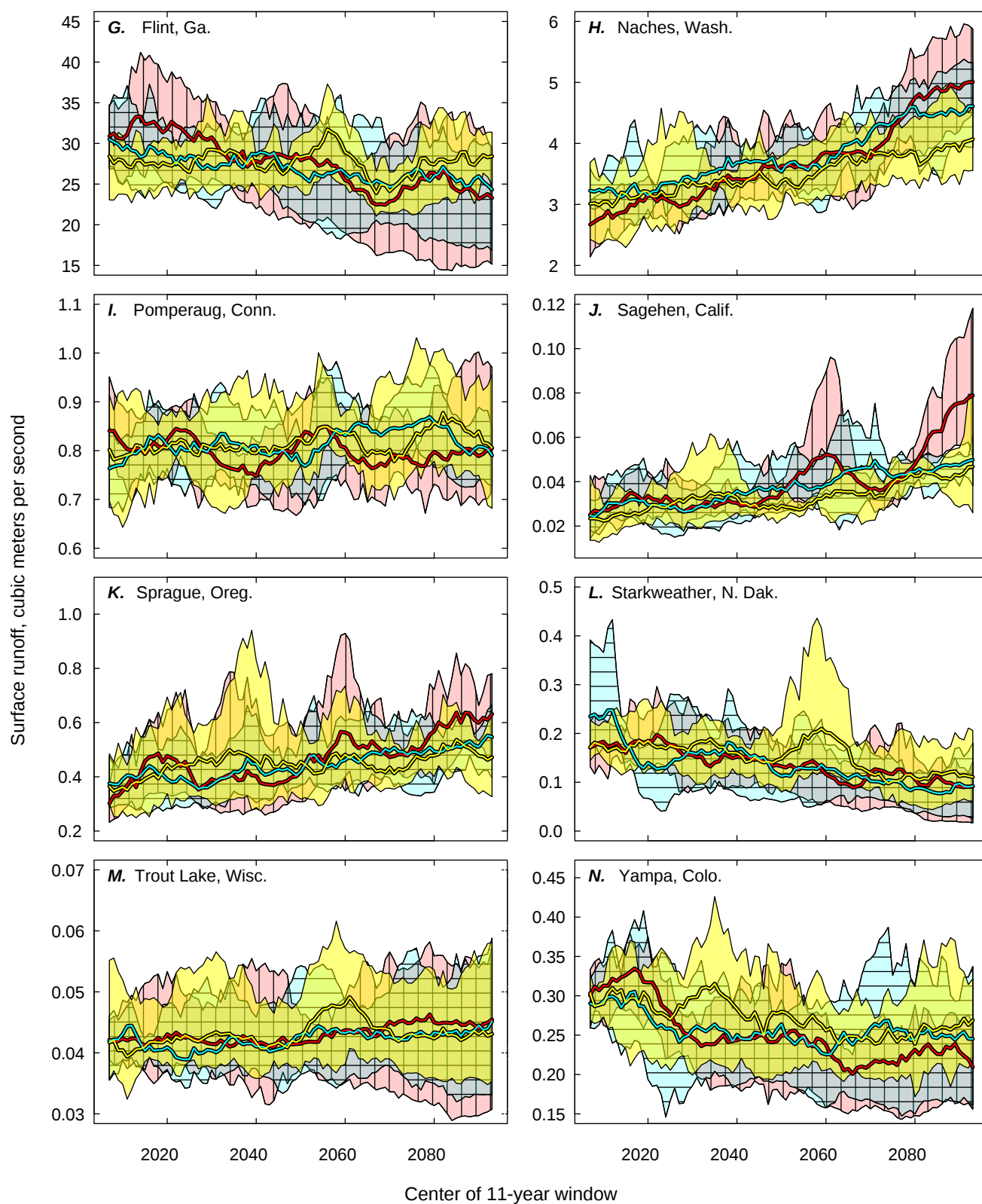


Figure 16. Projected range in 11-year moving mean daily surface runoff values by emission scenarios for the 14 basins.
—Continued

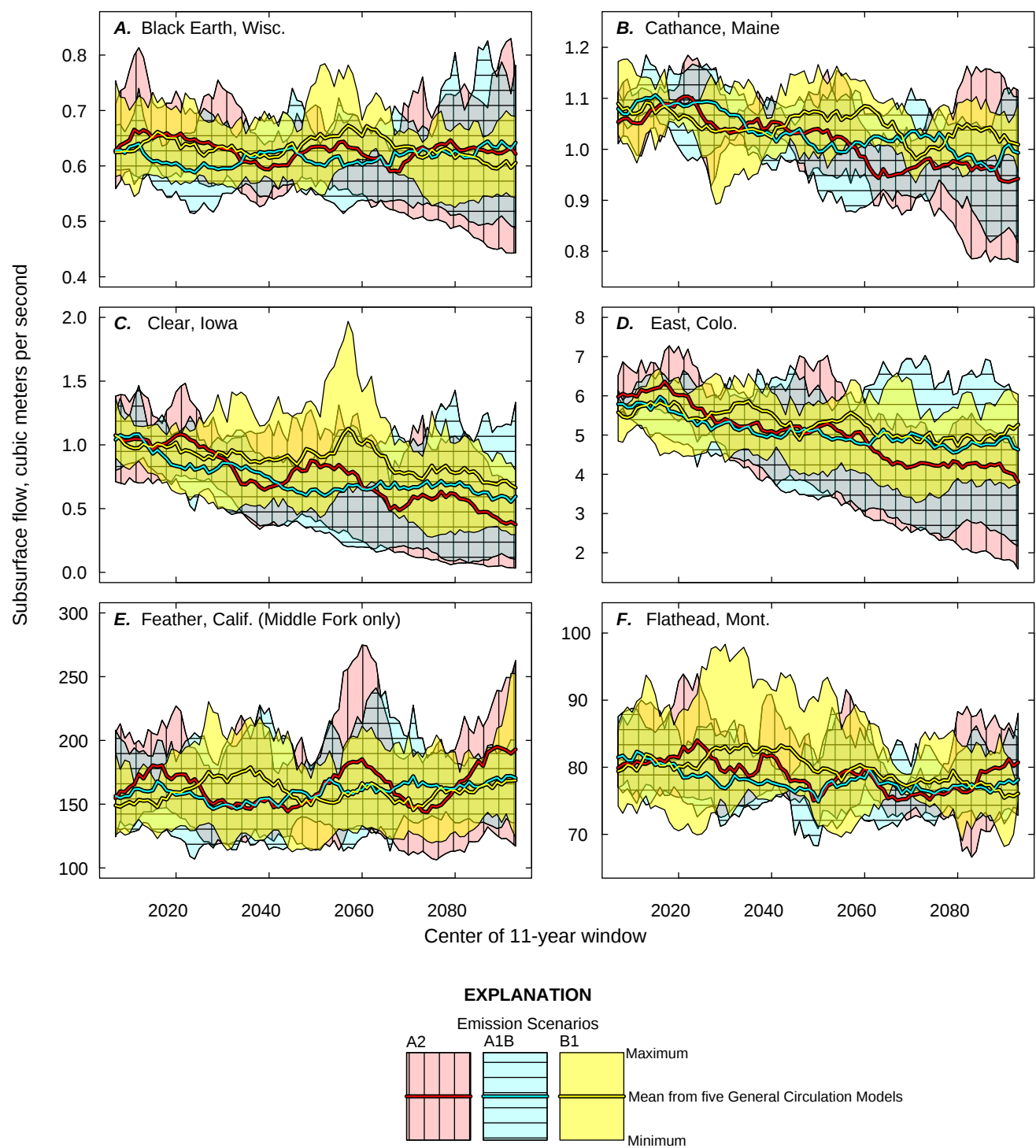


Figure 17. Projected range in 11-year moving mean daily subsurface flow values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

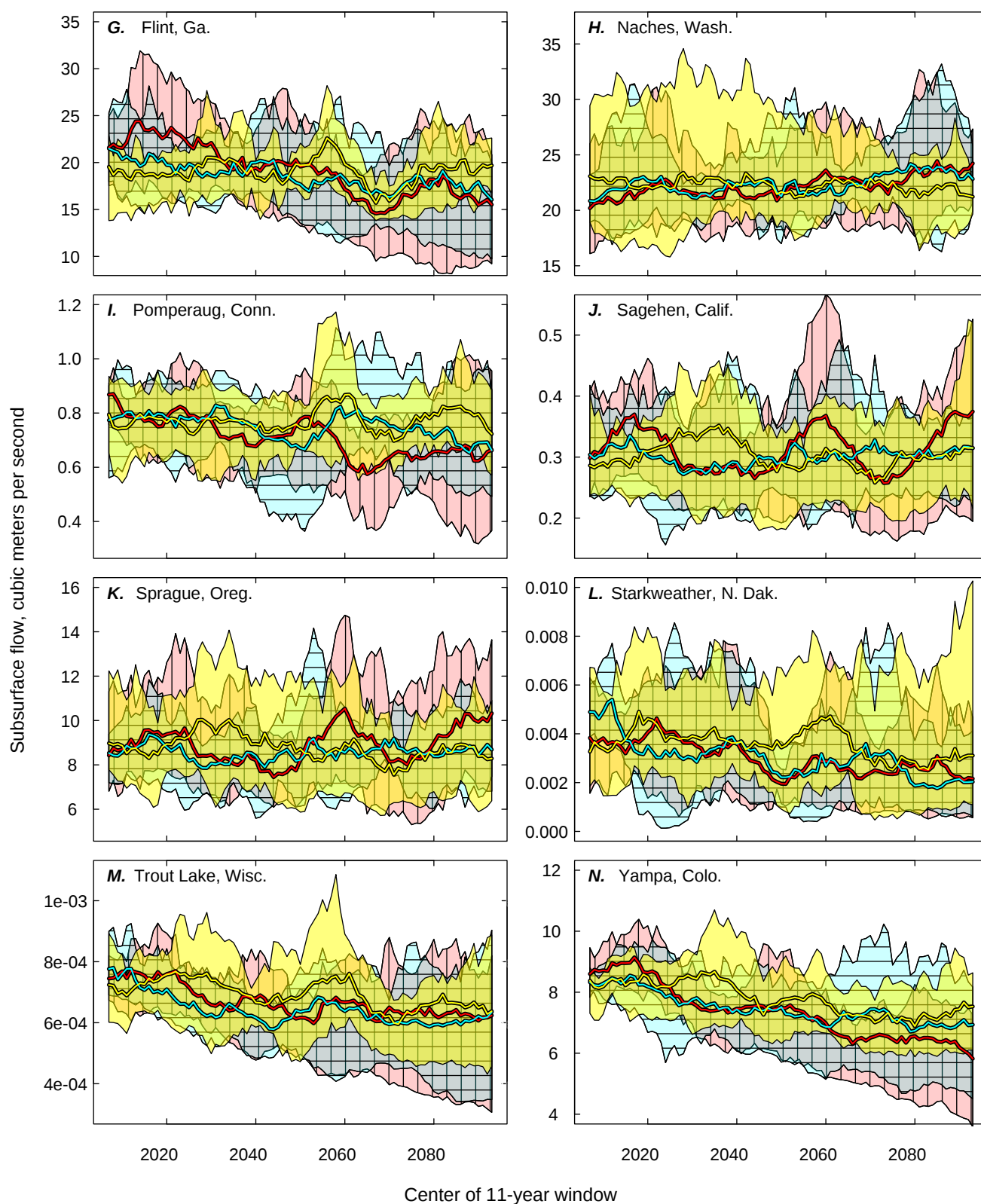


Figure 17. Projected range in 11-year moving mean daily subsurface flow values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

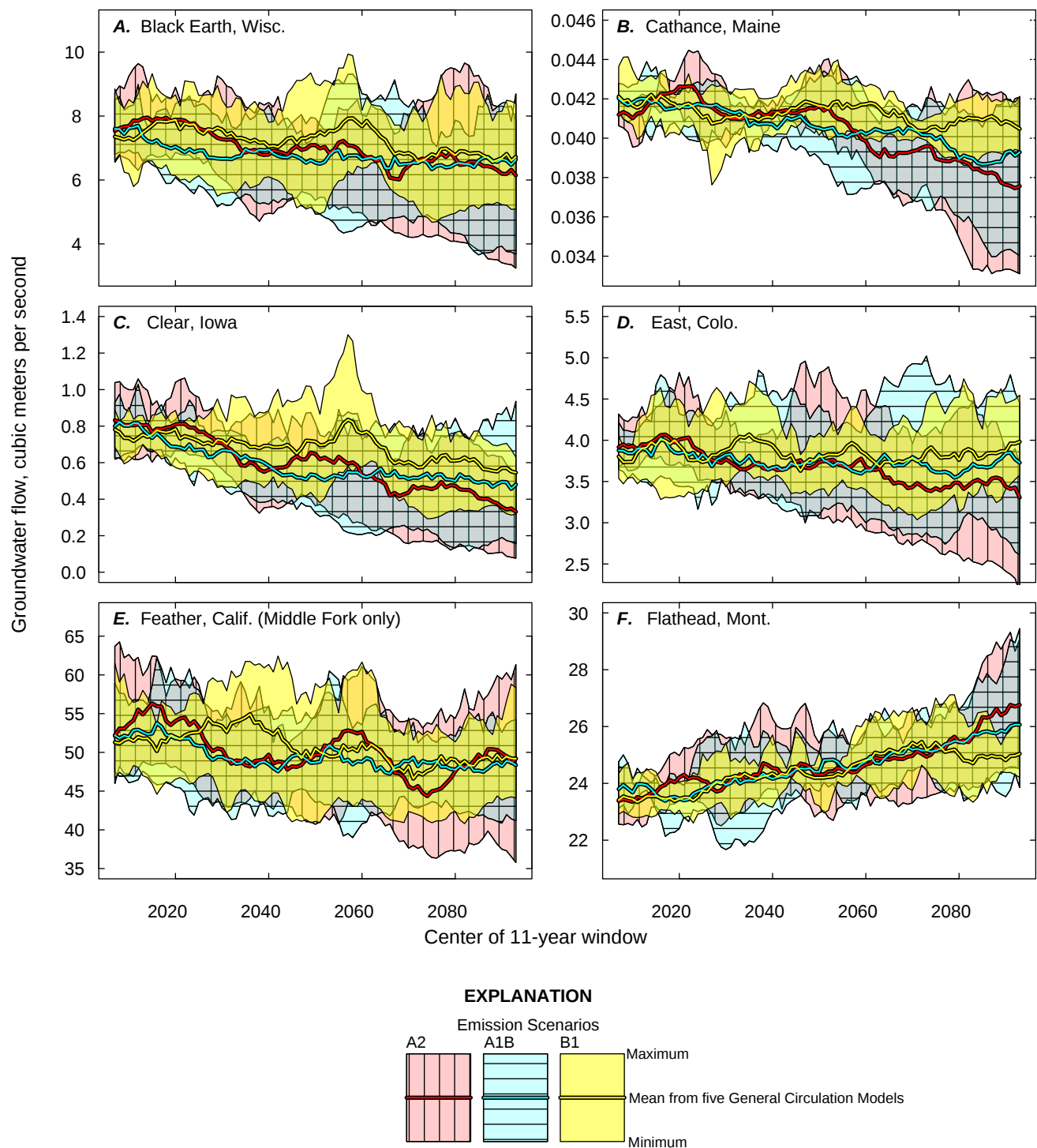


Figure 18. Projected range in 11-year moving mean daily groundwater flow values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

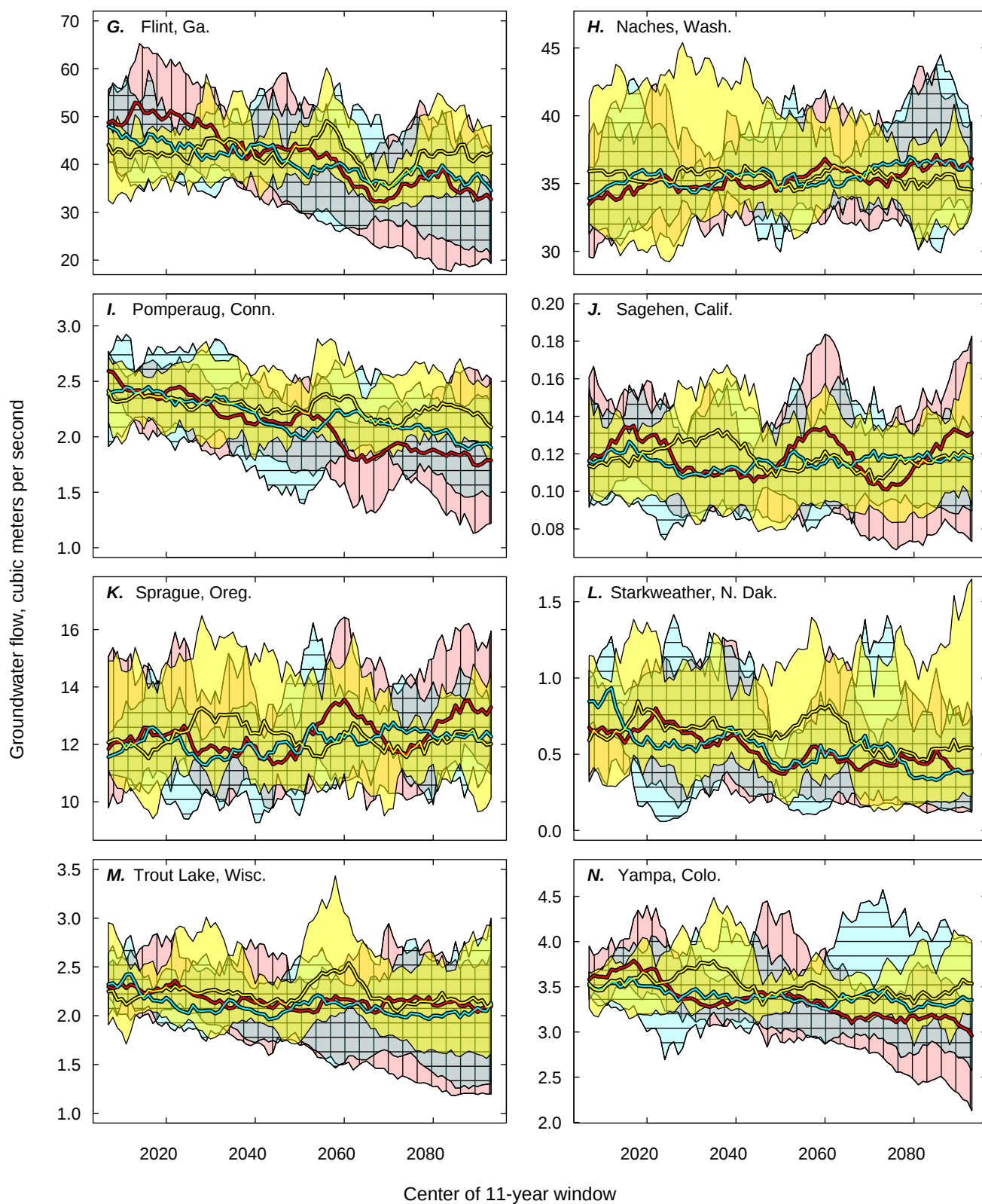


Figure 18. Projected range in 11-year moving mean daily groundwater flow values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

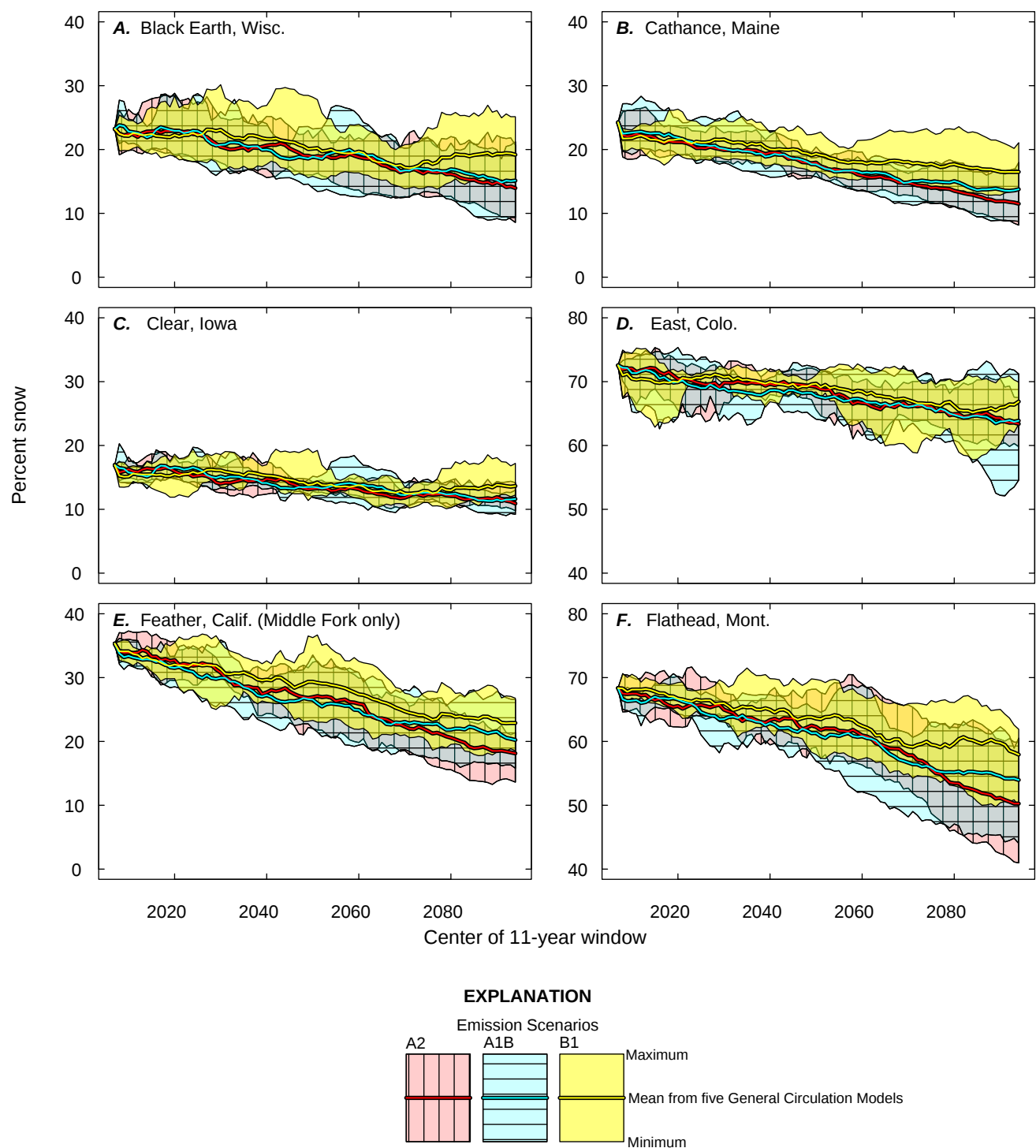


Figure 19. Projected range in 11-year moving mean daily percent of precipitation that falls as snow by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

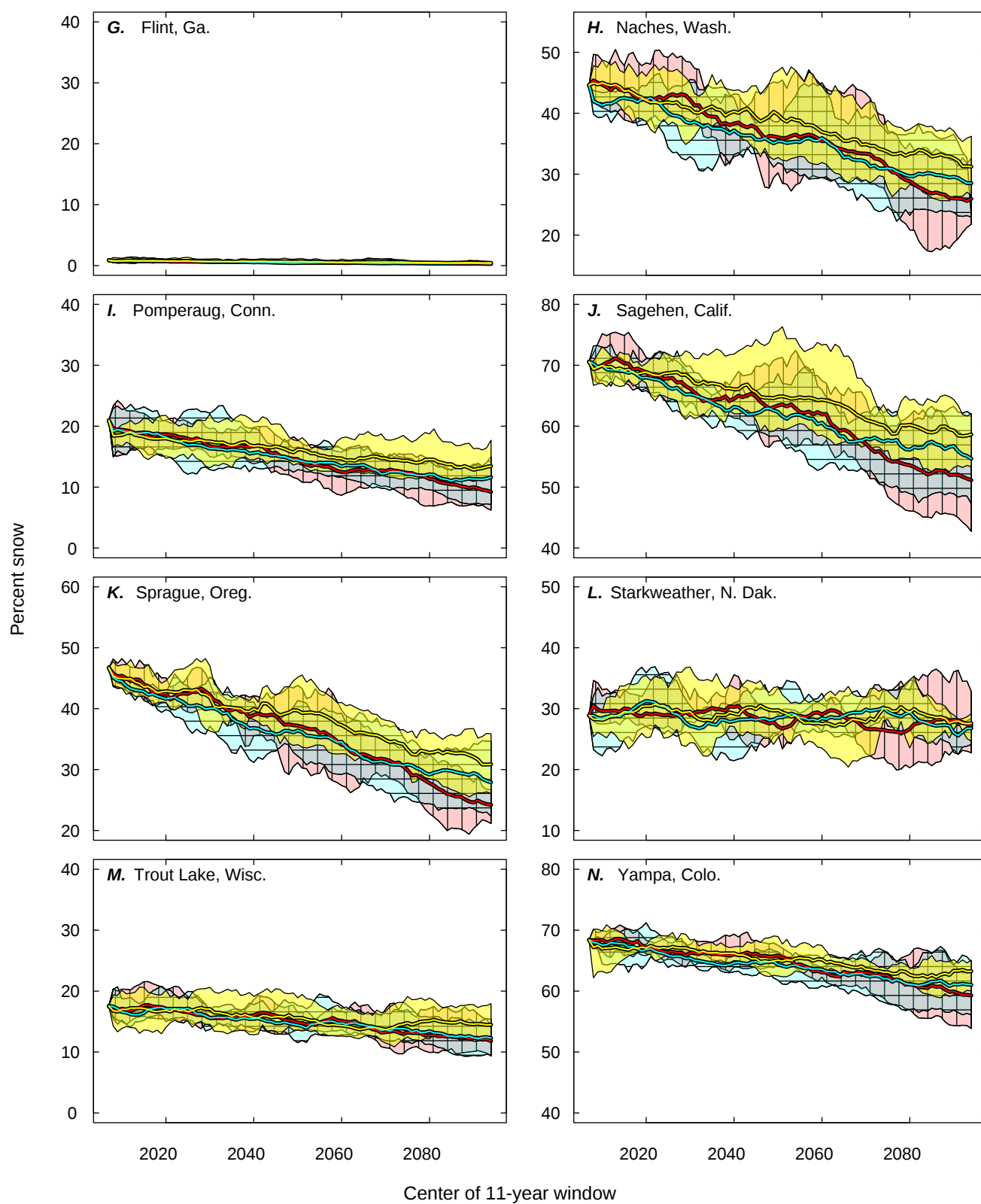


Figure 19. Projected range in 11-year moving mean daily percent of precipitation that falls as snow by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

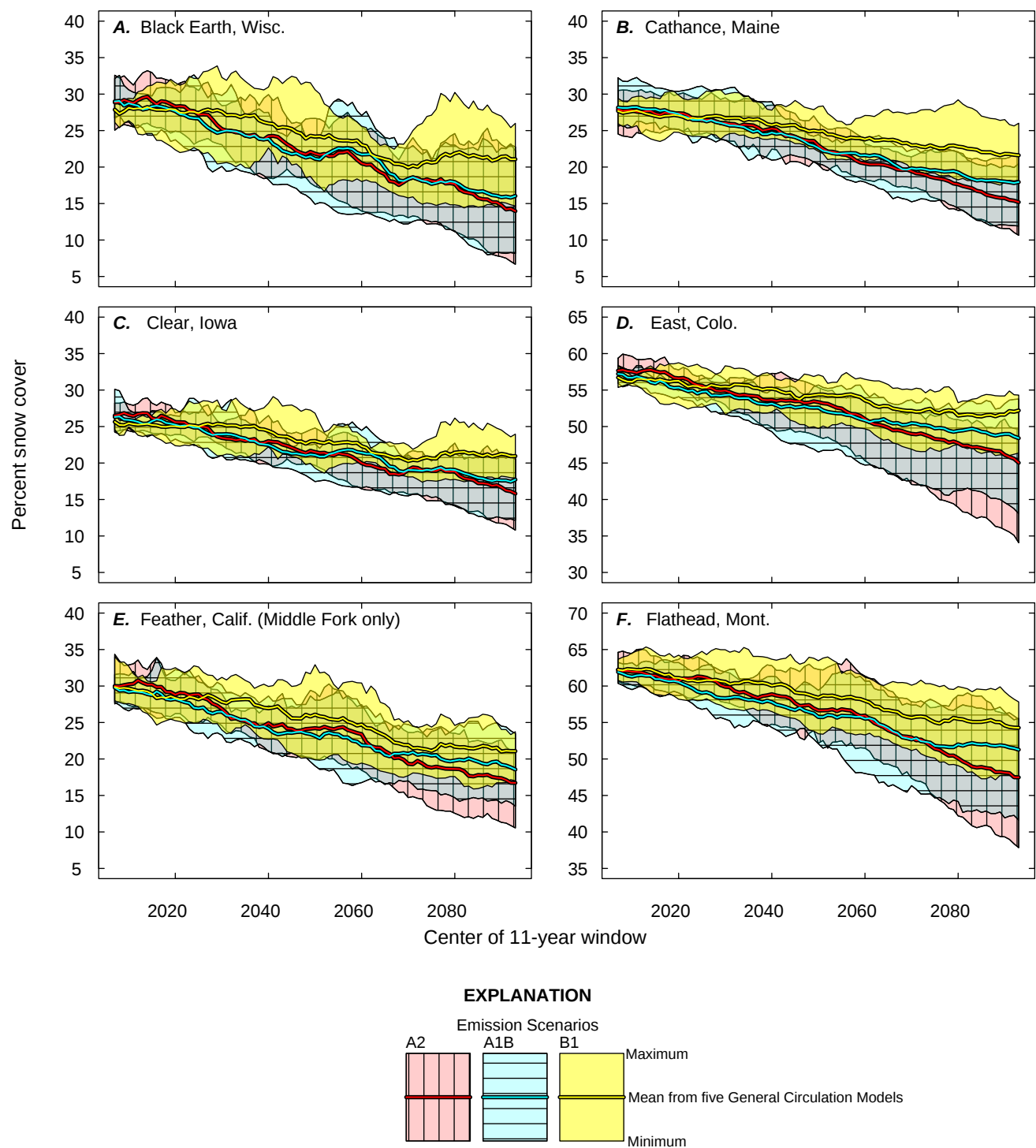


Figure 20. Projected range in 11-year moving mean daily percent of area covered by snow by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

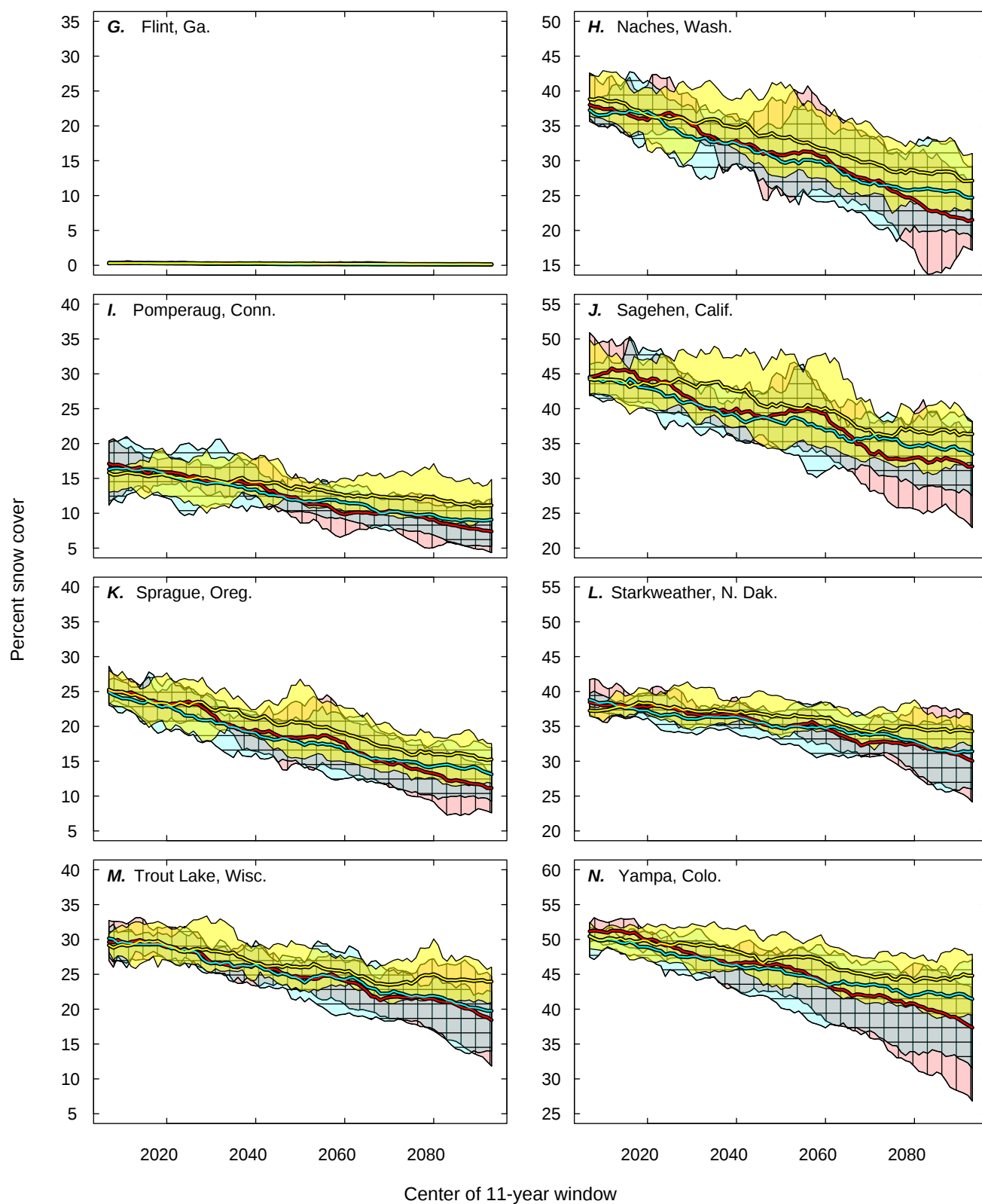


Figure 20. Projected range in 11-year moving mean daily percent of area covered by snow by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

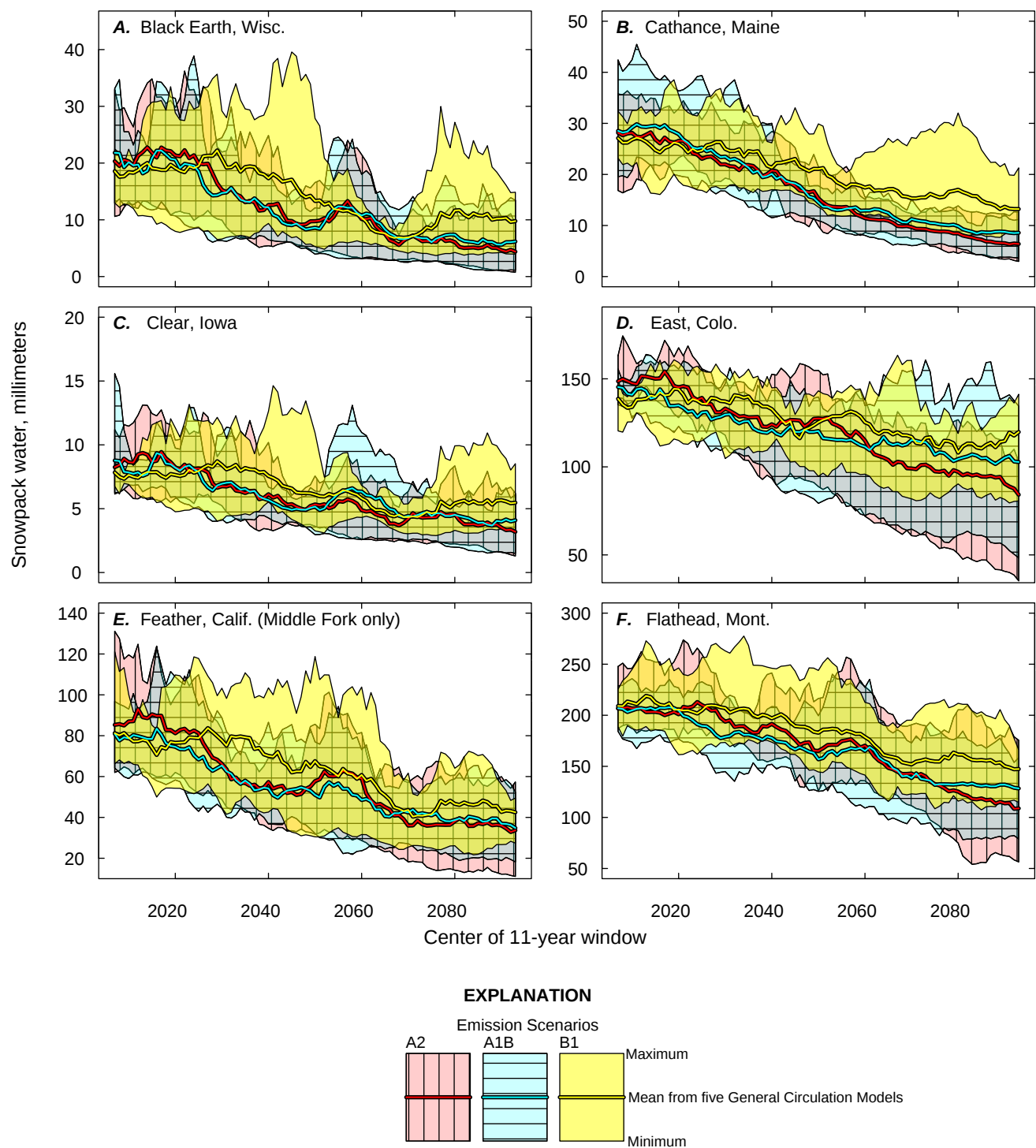


Figure 21. Projected range in 11-year moving mean daily snowpack water equivalent by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

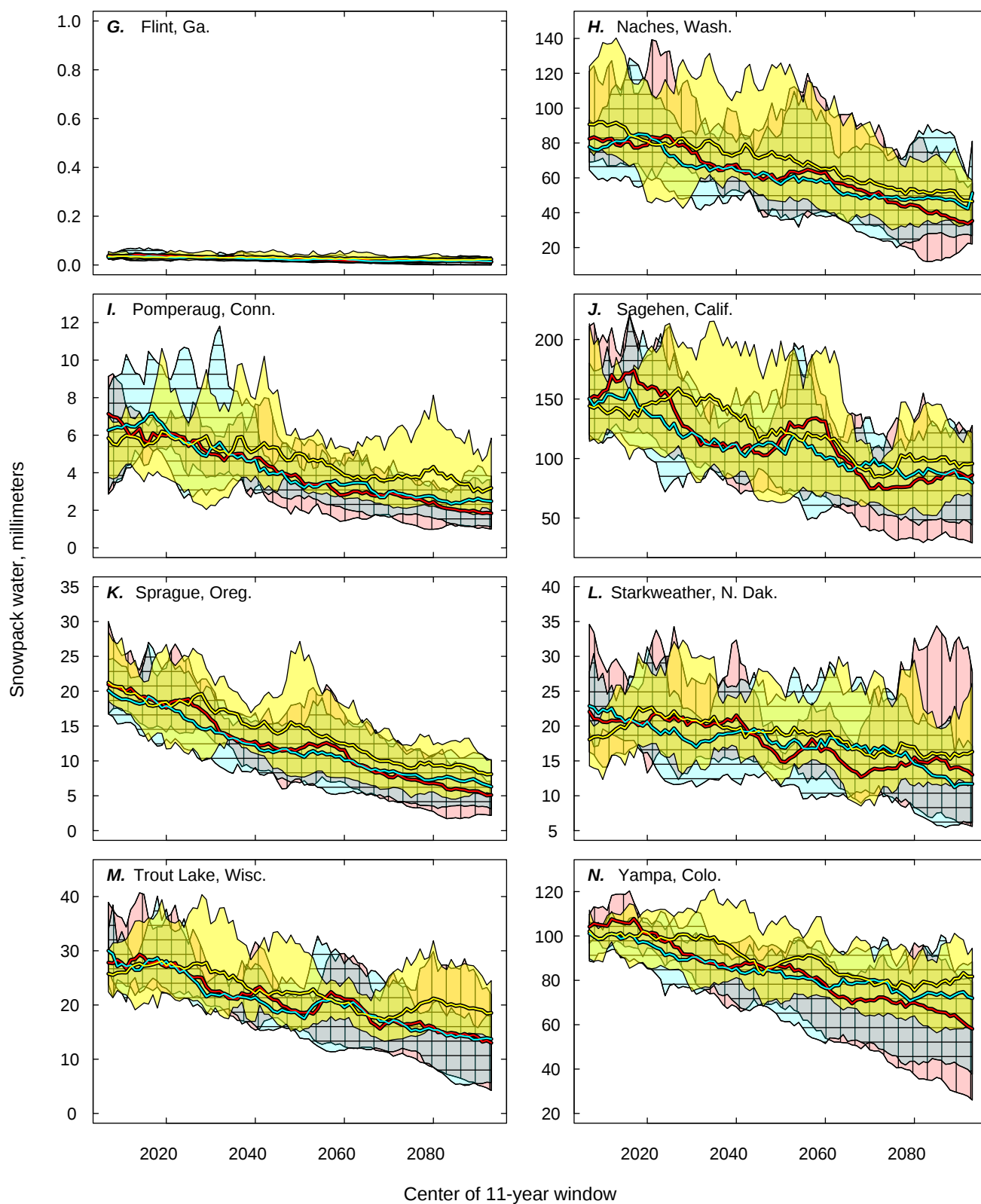


Figure 21. Projected range in 11-year moving mean daily snowpack water equivalent by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

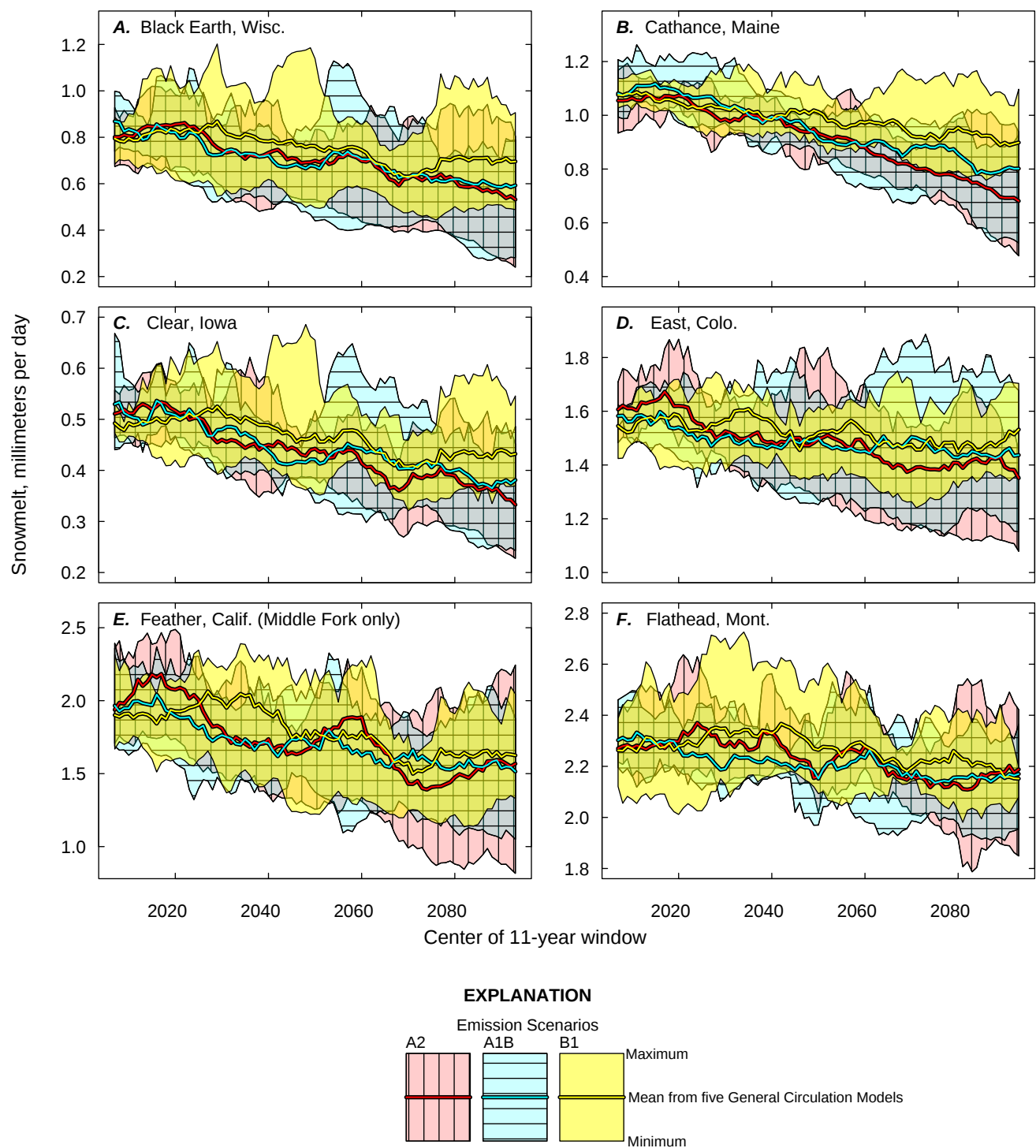


Figure 22. Projected range in 11-year moving mean daily snowmelt values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

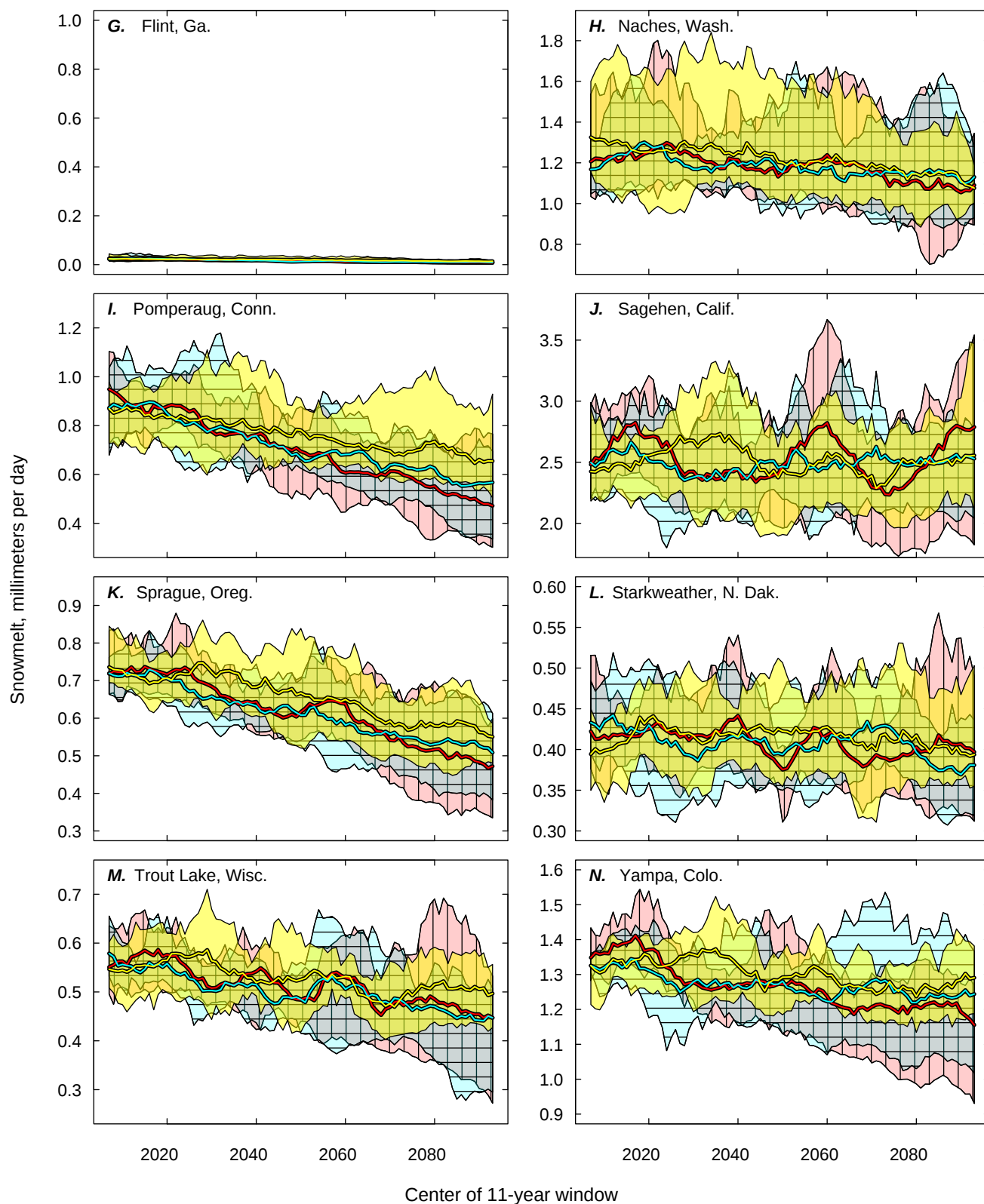


Figure 22. Projected range in 11-year moving mean daily snowmelt values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

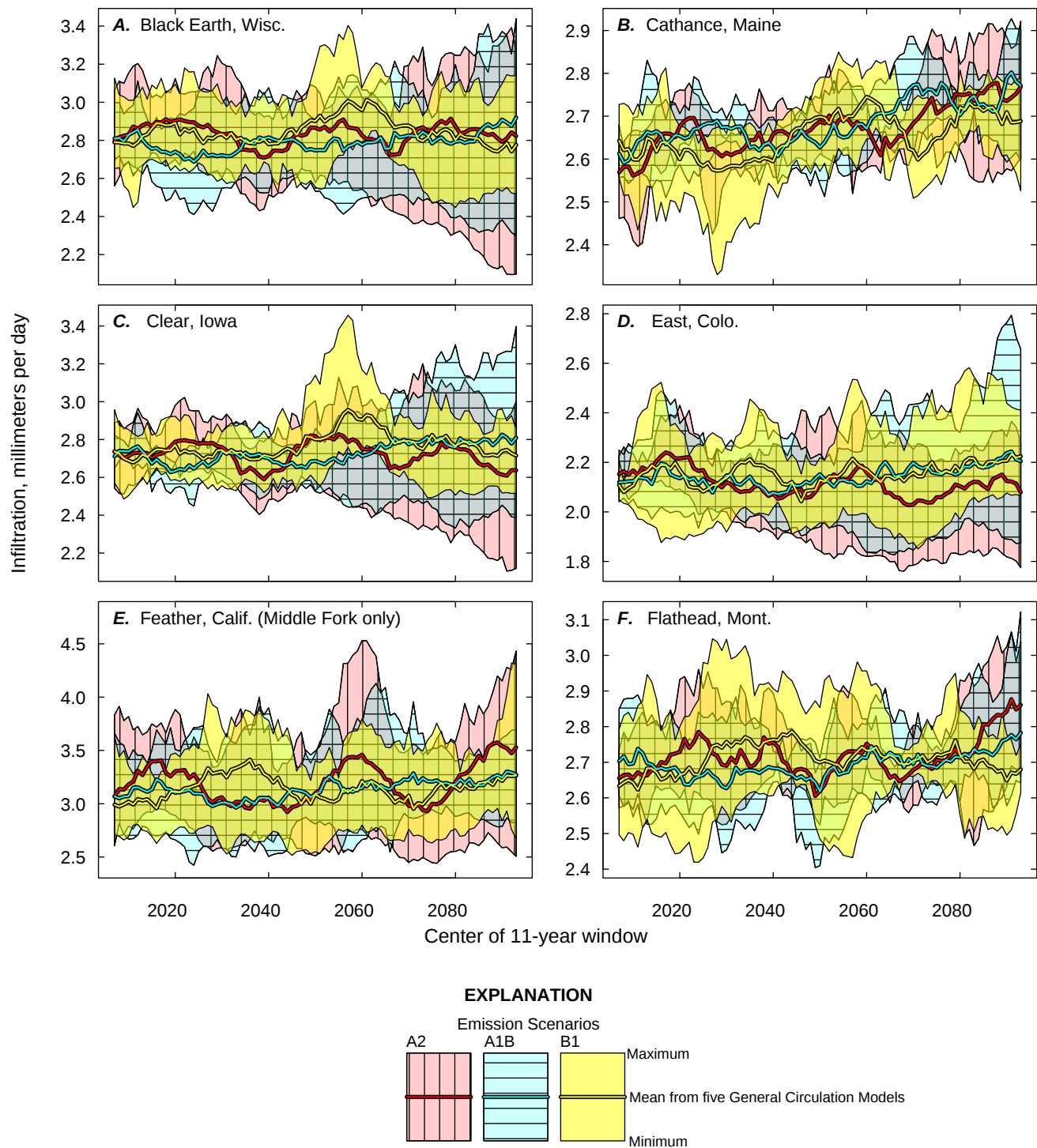


Figure 23. Projected range in 11-year moving mean daily infiltration values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

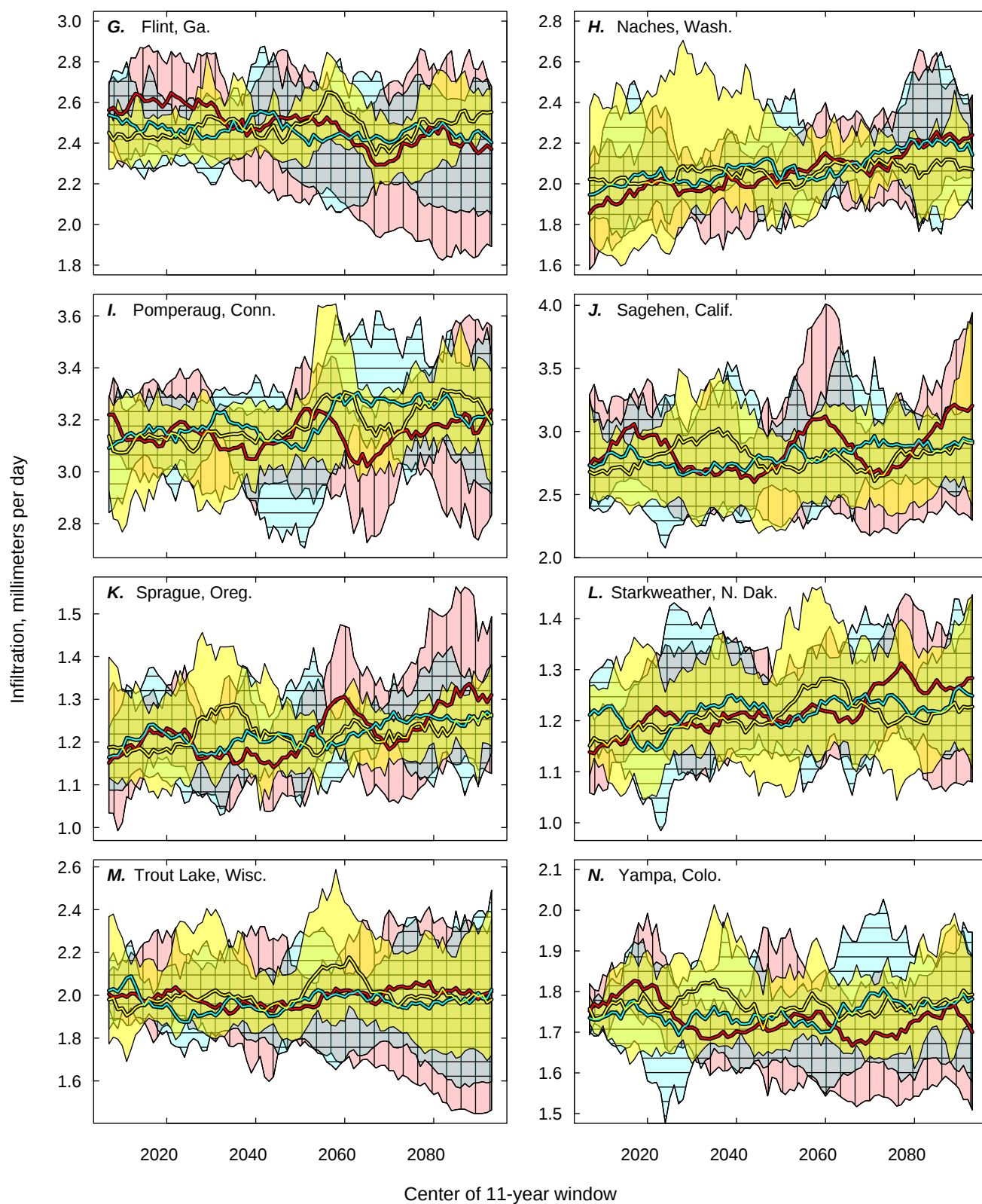


Figure 23. Projected range in 11-year moving mean daily infiltration values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

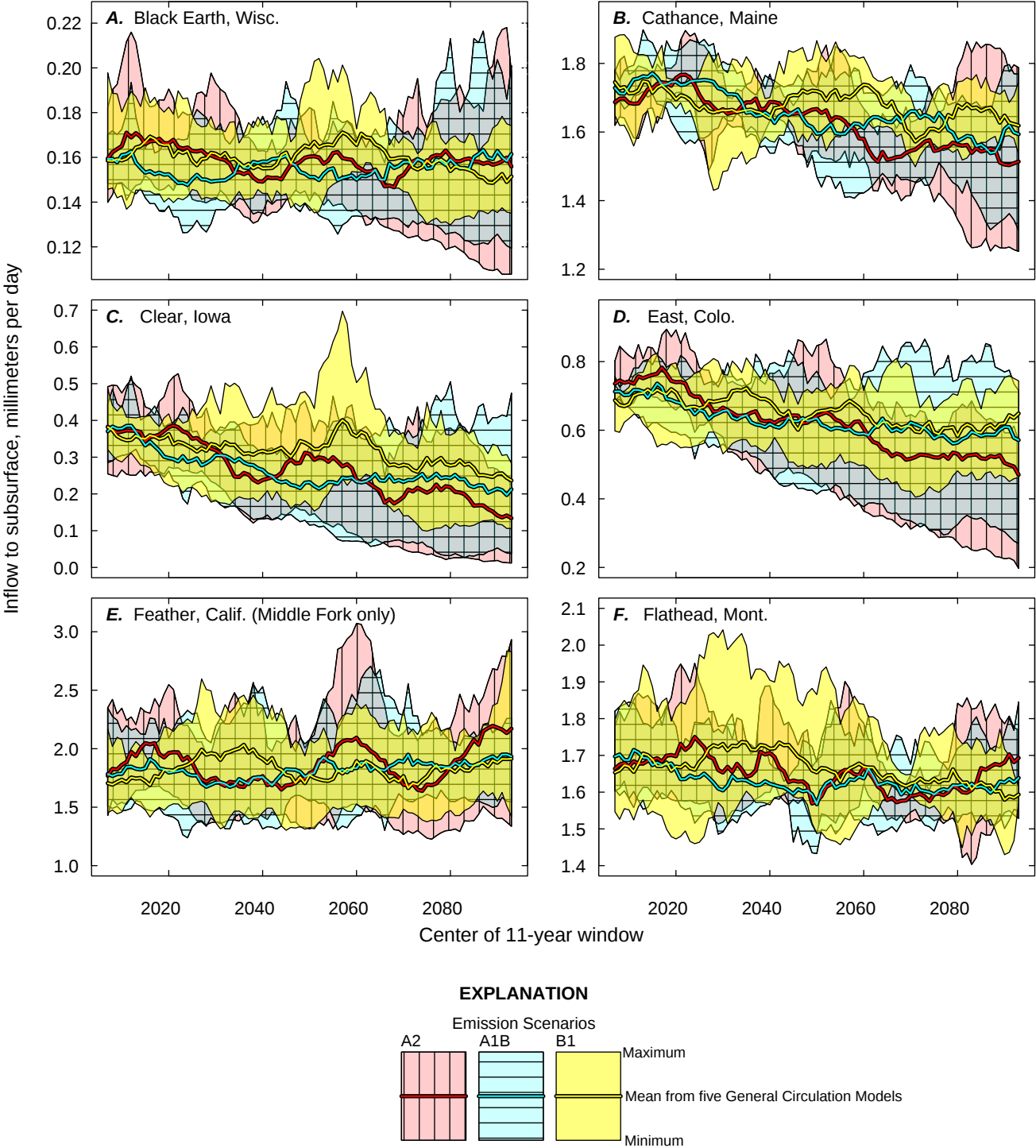


Figure 24. Projected range in 11-year moving mean daily subsurface reservoir inflow values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

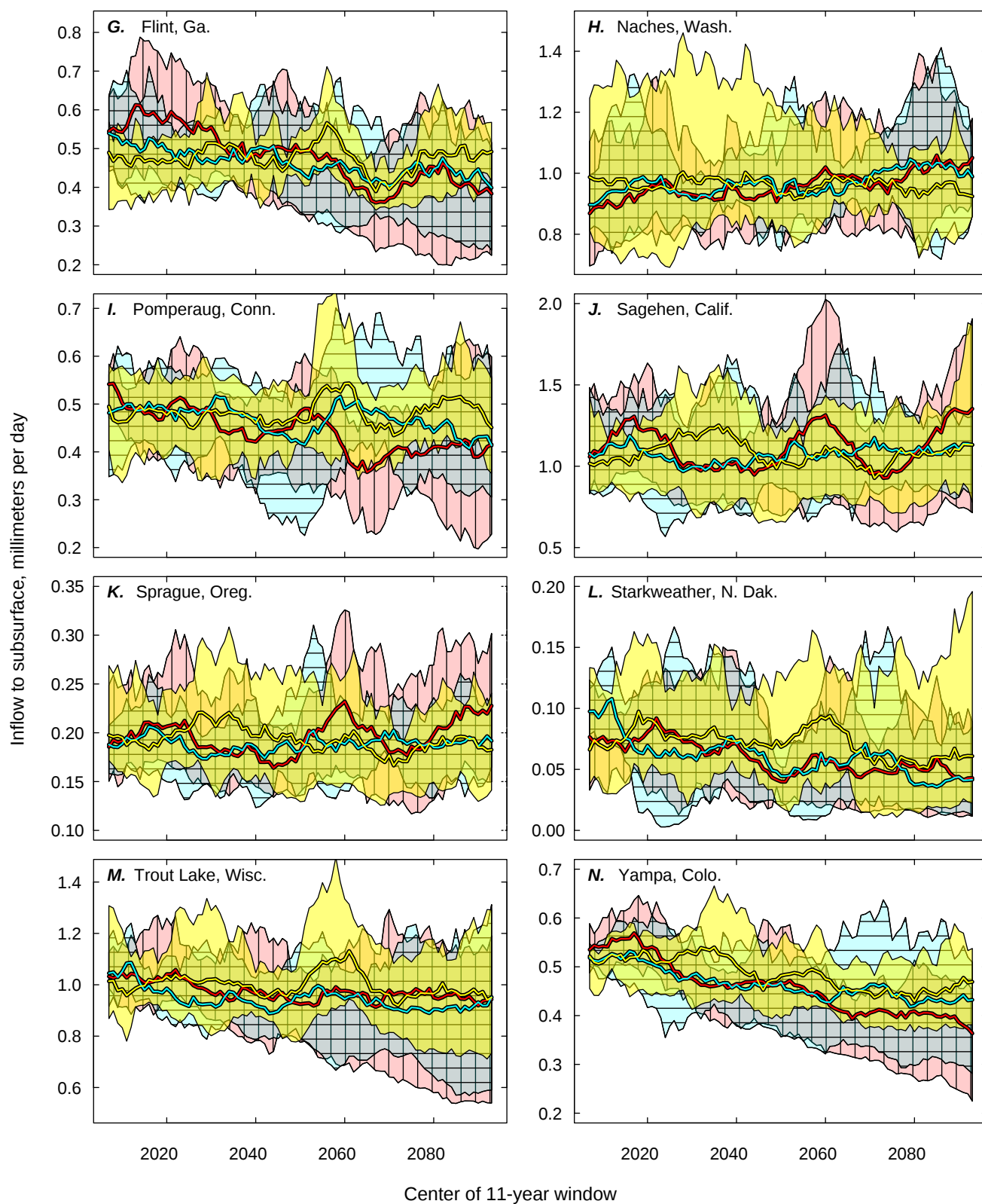


Figure 24. Projected range in 11-year moving mean daily subsurface reservoir inflow values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

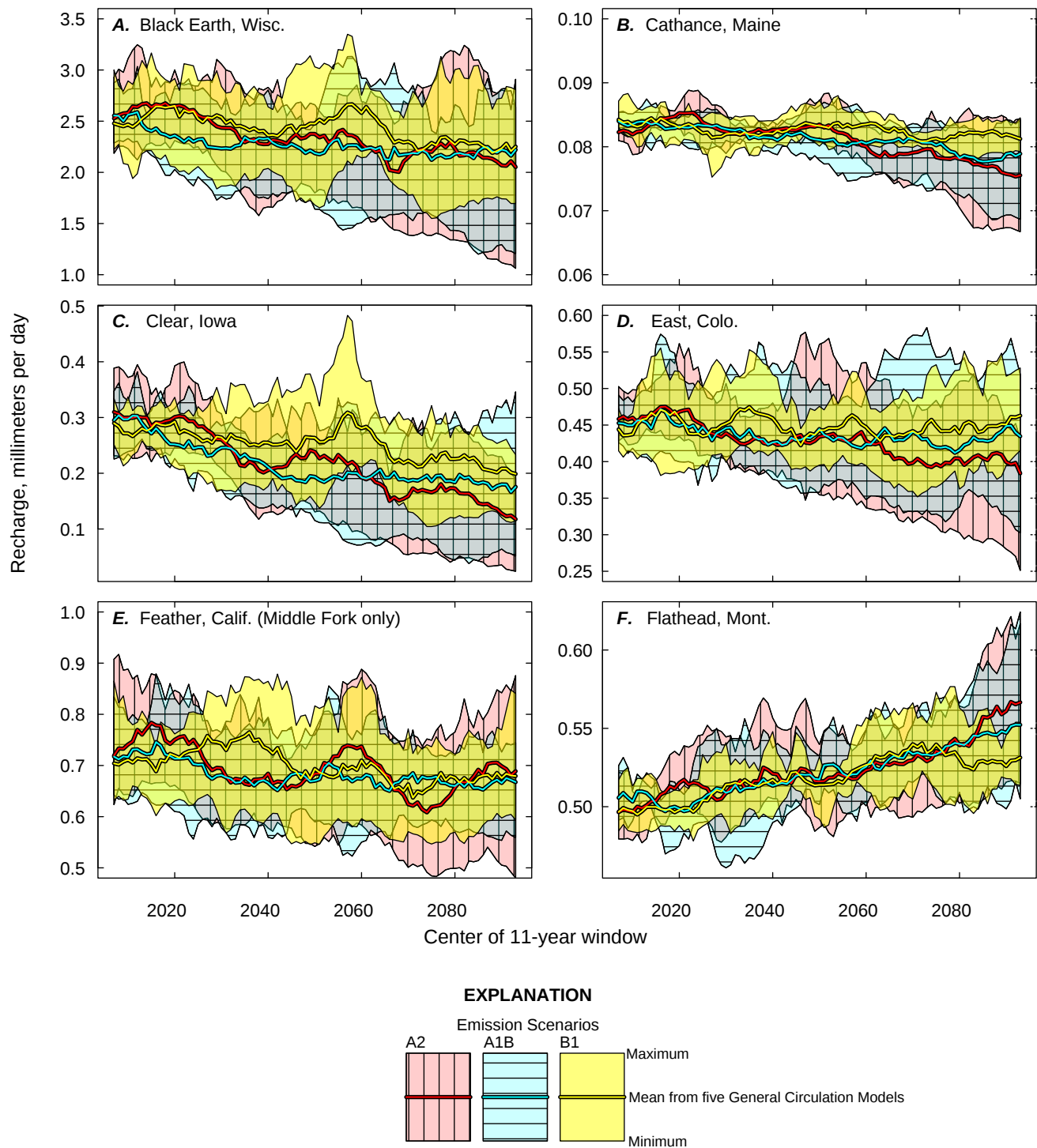


Figure 25. Projected range in 11-year moving mean daily groundwater reservoir recharge values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

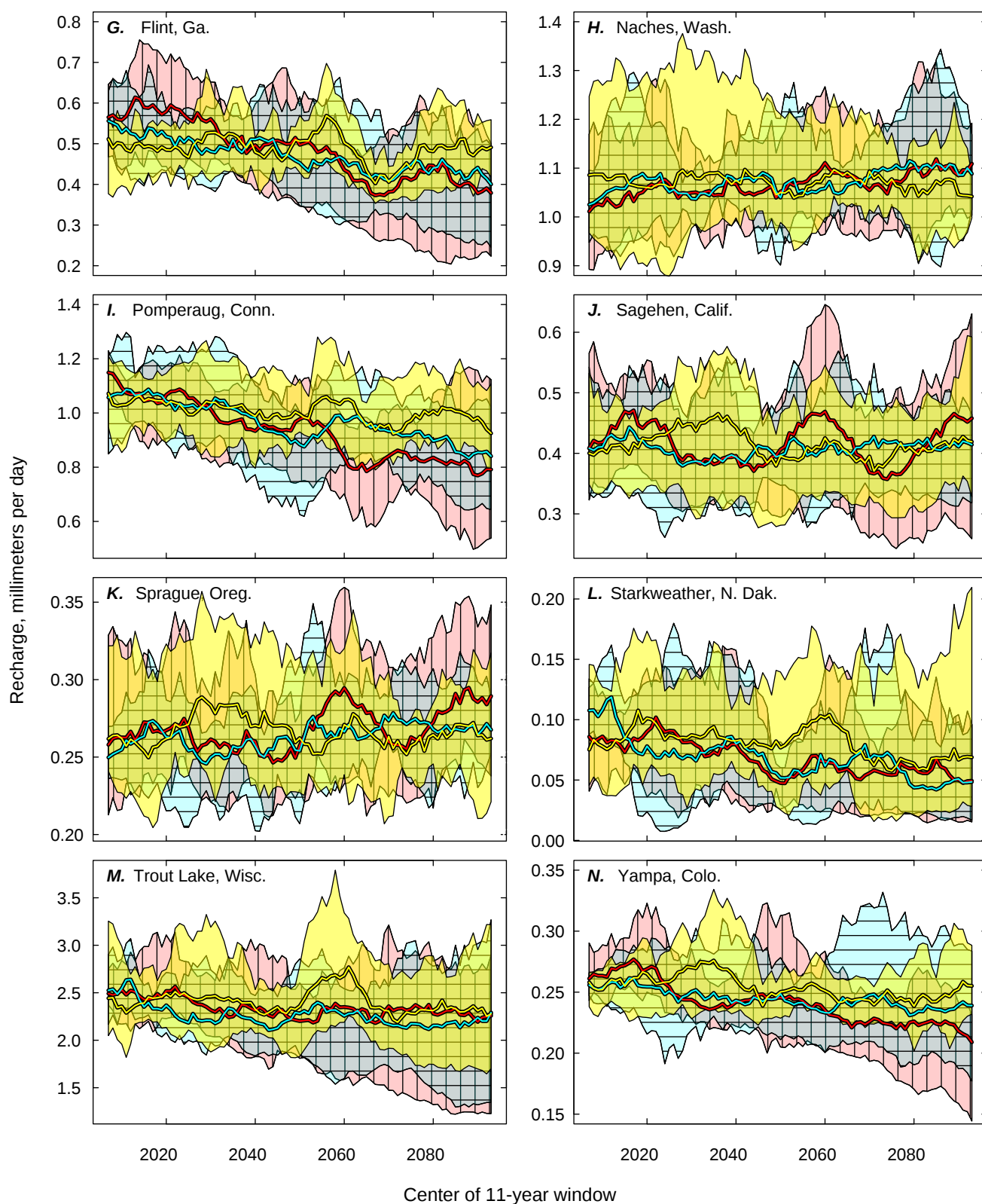


Figure 25. Projected range in 11-year moving mean daily groundwater reservoir recharge values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

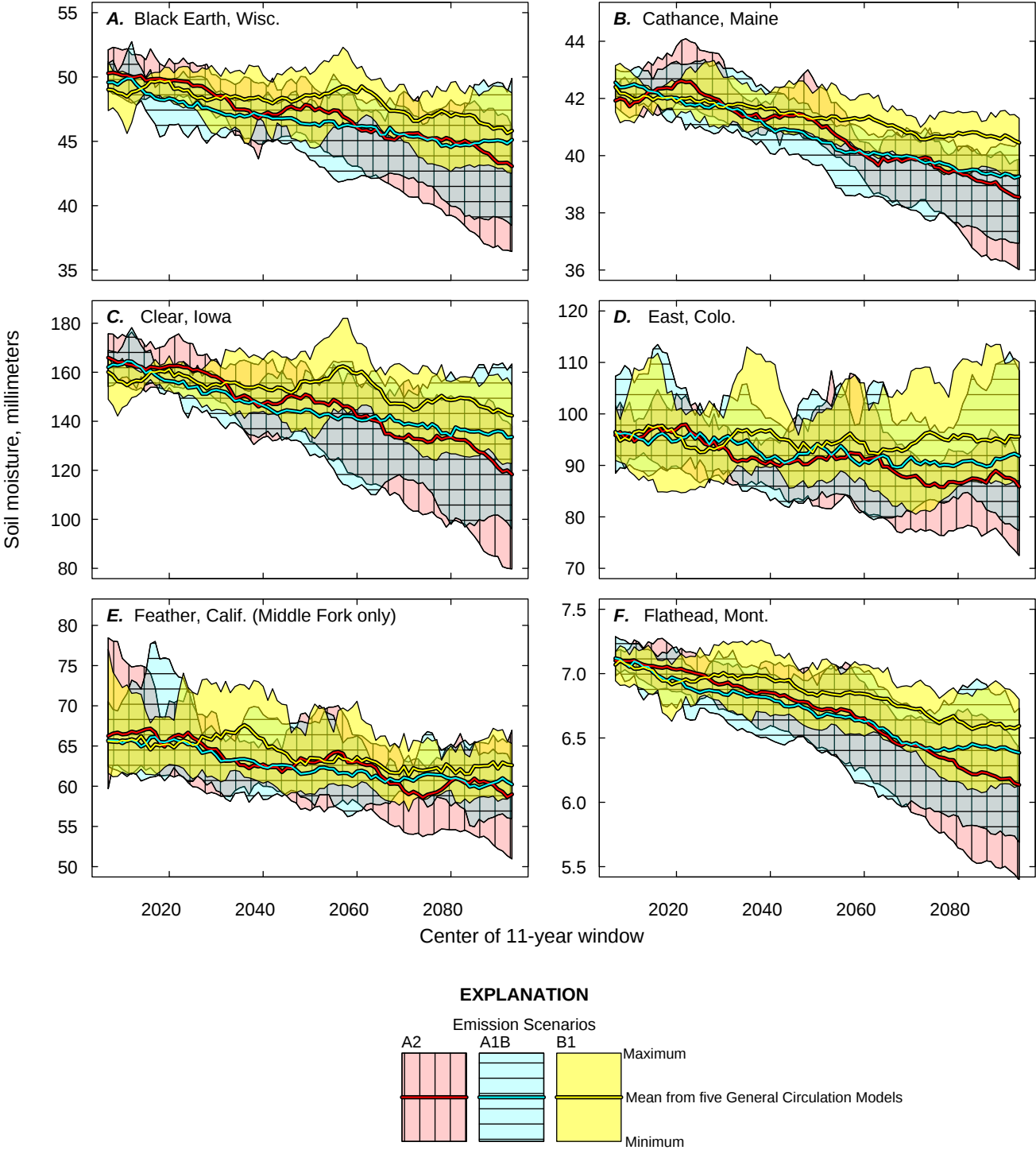


Figure 26. Projected range in 11-year moving mean daily soil moisture storage values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

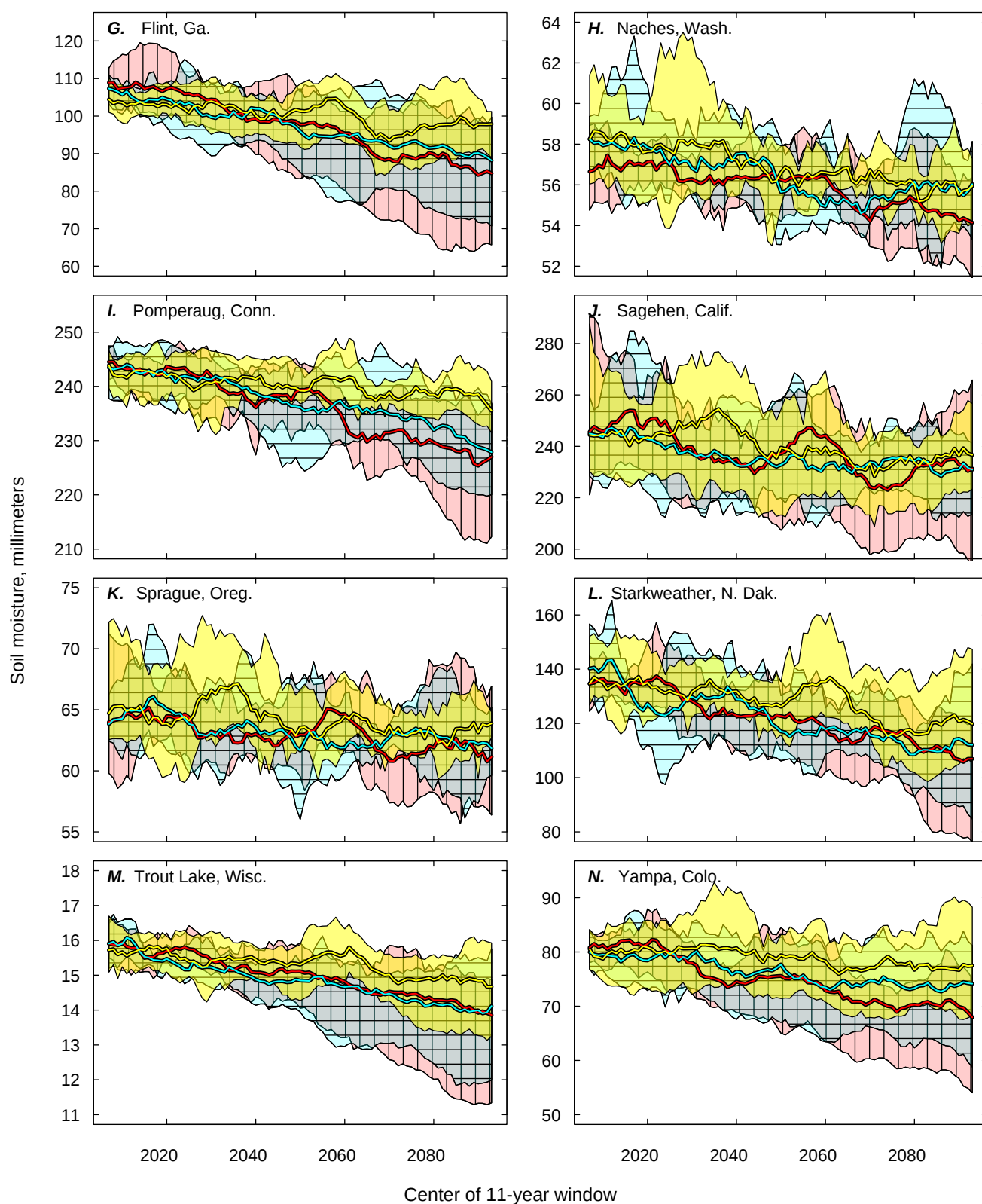


Figure 26. Projected range in 11-year moving mean daily soil moisture storage values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

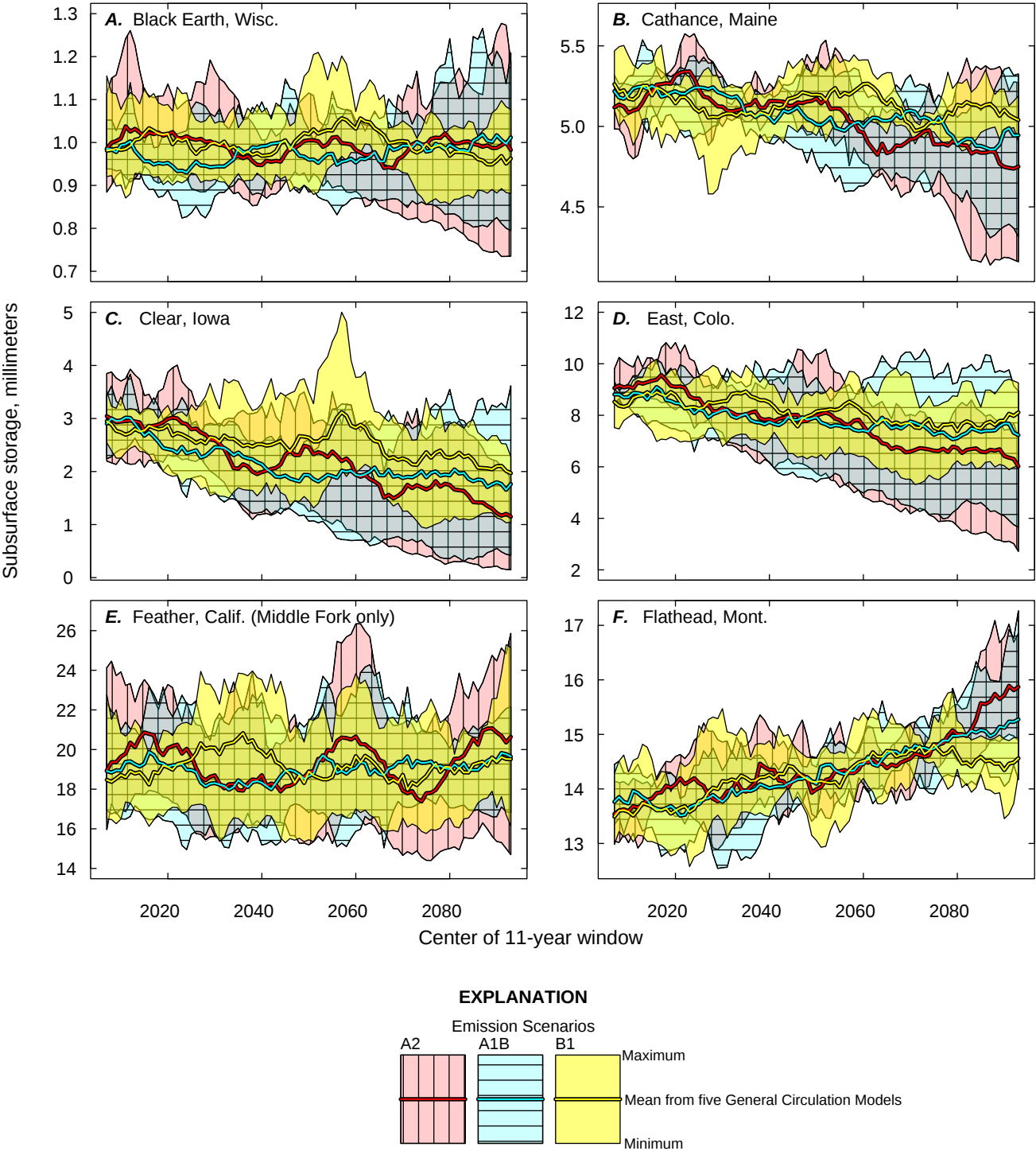


Figure 27. Projected range in 11-year moving mean daily subsurface reservoir storage values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

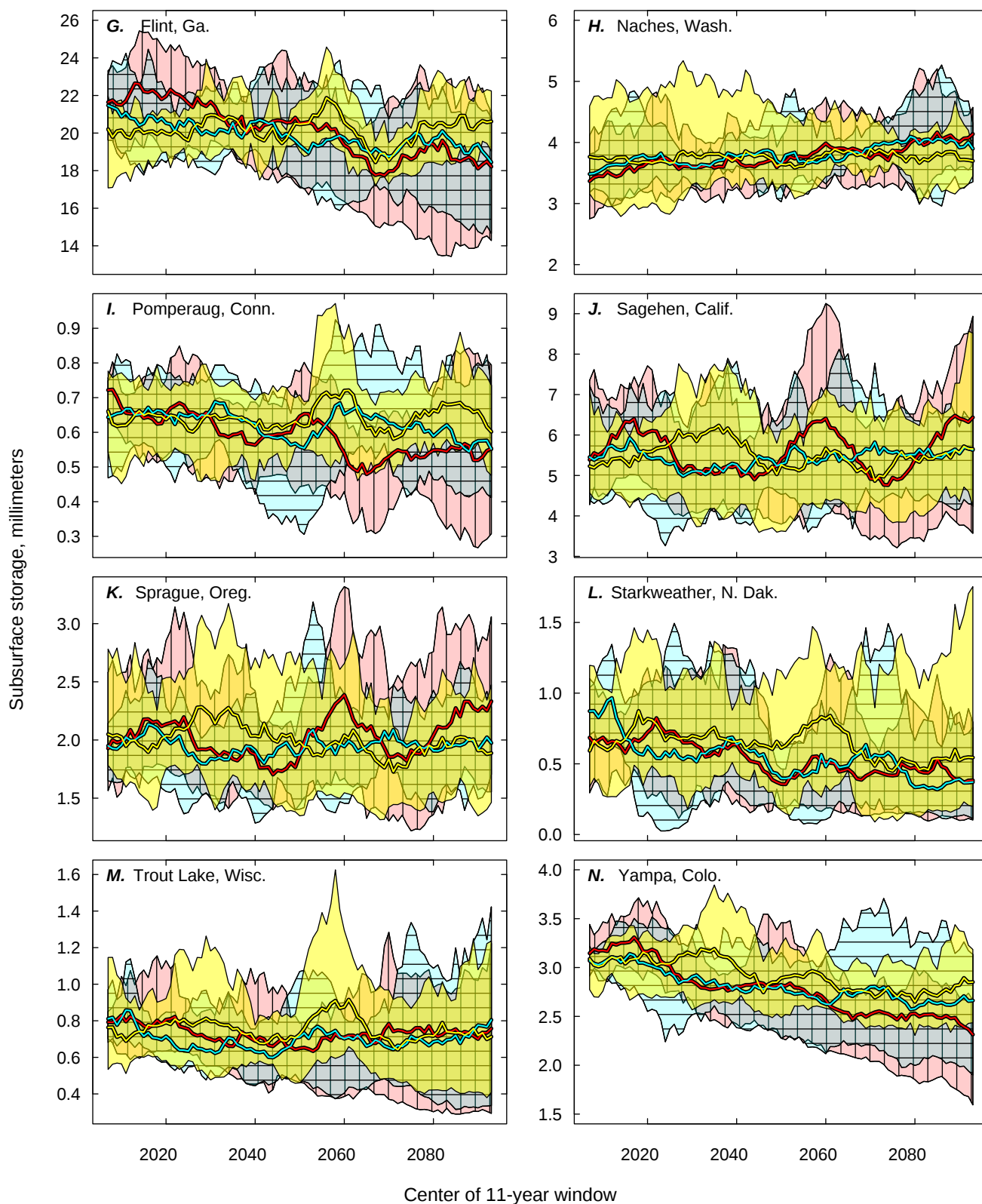


Figure 27. Projected range in 11-year moving mean daily subsurface reservoir storage values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

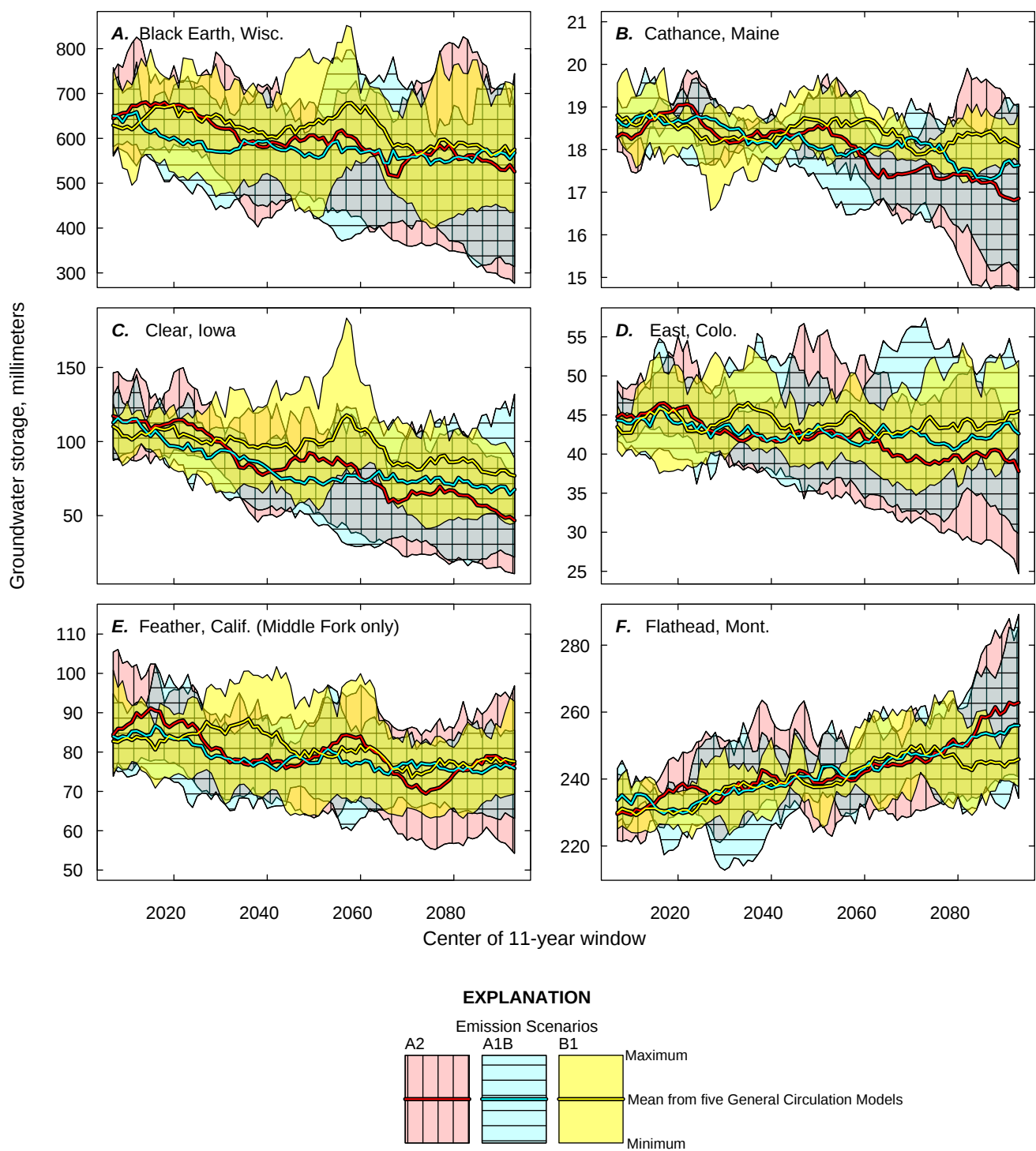


Figure 28. Projected range in 11-year moving mean daily groundwater reservoir storage values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).

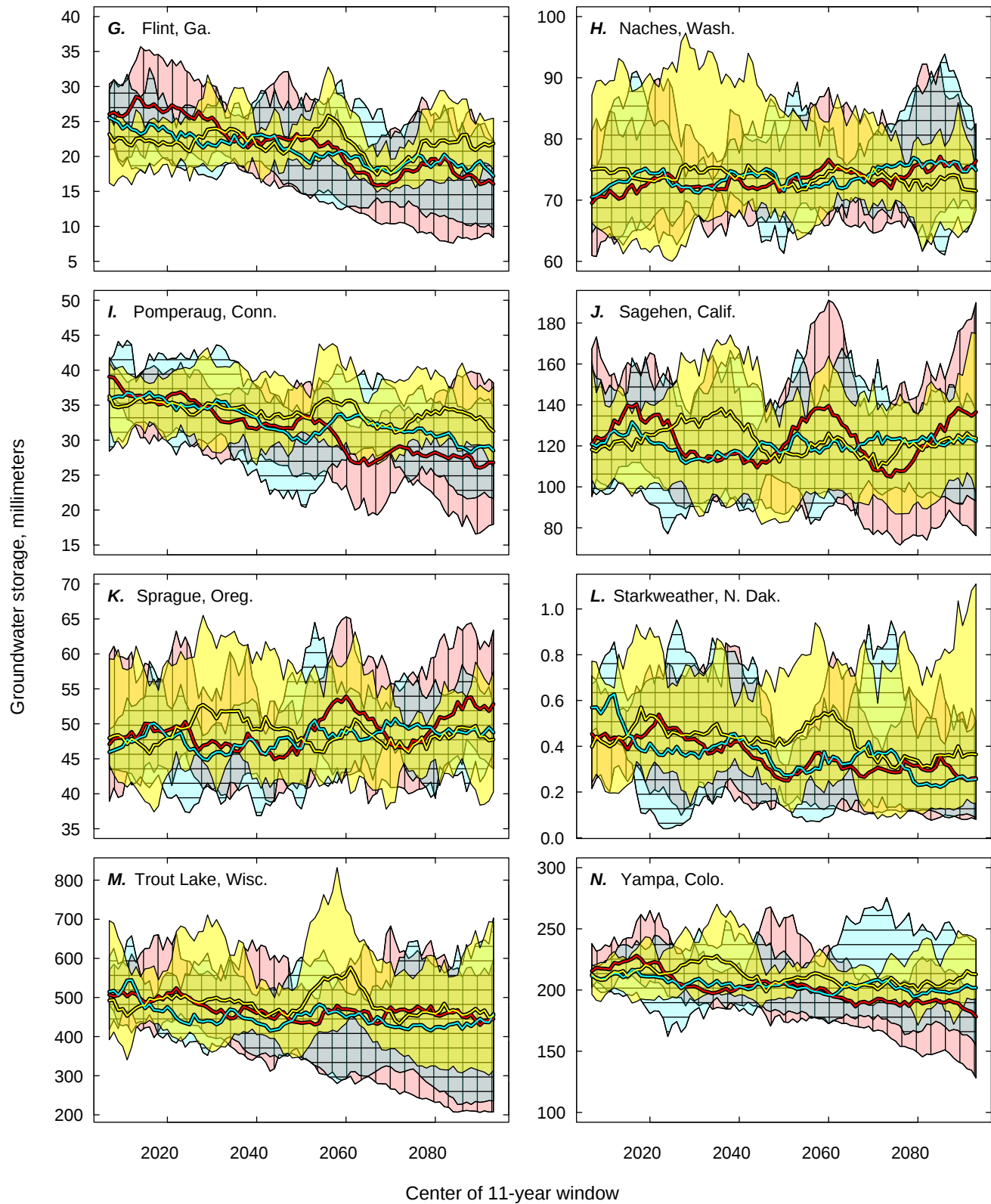


Figure 28. Projected range in 11-year moving mean daily groundwater reservoir storage values by emission scenarios for the 14 basins and means based on the 5 General Circulation Models (GCM).—Continued

Mean Monthly Plots

All mean monthly plots (figs. 29–46) show by month the mean daily values for baseline (red line, 1989–1999) and the range in PRMS projections for: 2030 (green, 2025–2035), 2060 (tan, 2055–2065), and 2090 (blue, 2085–2095). The range of values indicated by the boxplots illustrates the degree of uncertainty associated with the magnitude of projected changes on a monthly basis for the five GCMs and three scenarios. Each of these plots is described, basin by basin, in the following sections of this report.

Changes in the timing of peak flow are evident in projected mean monthly hydrographs (fig. 47) for baseline conditions (red dashed line, 1989–1999) and projections of 2030 (green line, 2025–2035), 2060 (tan line, 2055–2065), and 2090 (2085–2095) for the 14 PRMS applications. The timing of peak flow remains unchanged in cold weather basins (for example, April in the Starkweather projections) to shifts of several months for others (for example, May to March in the Sagehen projections). In many of the basins, projections of streamflow indicate a decrease in the effect of snowmelt and an increase in the effect of evapotranspiration, though the uncertainty is large.

Table 5 shows the change (slope) and R^2_a for selected output variables from the PRMS projections. The slope indicates the amount of change in the selected variable by year; blue colored values indicates a significant negative trend and red colored values indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation on the degrees of freedom (Lettenmaier, 1976; McCabe and Wolock, 1997). The R^2_a value gives an indication of the variability in the central tendency of the trend.

Precipitation Runoff Modeling System Projections for Black Earth Creek, Wisconsin

The Black Earth Creek basin is located in northwest Dane County, Wisconsin. At USGS streamflow-gaging station 05406500, the basin has a drainage area of 118 km² (table 1). The basin was simulated with a fully-coupled groundwater/surface-water model (GSFLOW; Markstrom and others, 2008). Only the results from the PRMS portion of the GSFLOW application are presented here. The Black Earth Creek basin application contains 780 HRUs with elevations that range from 718 to 1,551 m (table 1).

Mean monthly baseline conditions for the Black Earth Creek basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from just below –12 °C during January to just below 16 °C during July. Mean monthly maximum temperatures range from approximately –2 °C during January to above 27 °C during July. Mean monthly precipitation ranges from 1.8 mm/day (during December and February) to almost 6.0 mm/day (during July). Because of freezing temperatures

during the winter months, precipitation falls as snow and accumulates throughout the winter, with peak mean monthly runoff occurring during March because of snowmelt.

The Black Earth Creek basin encompasses both glaciated and unglaciated settings. The basin contains a trout stream located near an urbanizing fringe of the state capital, Madison. As a result of the conversions from agricultural to residential and commercial land uses, long-term evaluation of the effect of urbanization on the base flow and stormflow is desired. Urbanization may increase both stormflow (Steuer and Hunt, 2001) and nonpoint source loads of nutrients, pesticides, and sediments. Increased surface flows reduce the amount of water that would normally recharge to the groundwater system (Hunt and Steuer, 2001); thus, urbanization can result in a reduction of groundwater flows into streams.

Black Earth Creek Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Black Earth Creek basin (fig. 10A). The large range in the precipitation projections indicate a high level of uncertainty with no significant overall trend (table 5A). Seasonal variability is apparent also in these precipitation projections (fig. 29A).

All PRMS projections indicate steady increases in temperature (figs. 11A, 12A, 30A, 31A, and tables 5B and 5C), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario and largest for the A2 scenario (tables 5B and 5C).

All projections indicate steady increases in annual mean evapotranspiration (fig. 13A and table 5D). High variability in seasonal evapotranspiration is projected, with increases in all months, except August (fig. 32A).

All PRMS projections indicate steady increases in the length of the growing season (mid-April to mid-October) (fig. 14A and table 5E). This increase is because of the projected increase of annual temperature (figs. 11A and 12A and tables 5B and 5C). The projected changes in the length of the growing season are smallest for the B1 scenario and largest for the A2 scenario (table 5E).

Black Earth Creek Streamflow Variables

The PRMS projections indicate a decrease in annual mean streamflow for the A1B and A2 scenarios (fig. 15A and table 5F), though the uncertainties associated with these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, particularly in March (fig. 33A). Specifically, the PRMS simulation of baseline conditions (red line, fig. 33A) indicates that peak streamflow occurs in March, while the PRMS projections indicate that, by the end of the 21st century, the projected month of peak streamflow will no longer be discernible (fig. 33A).

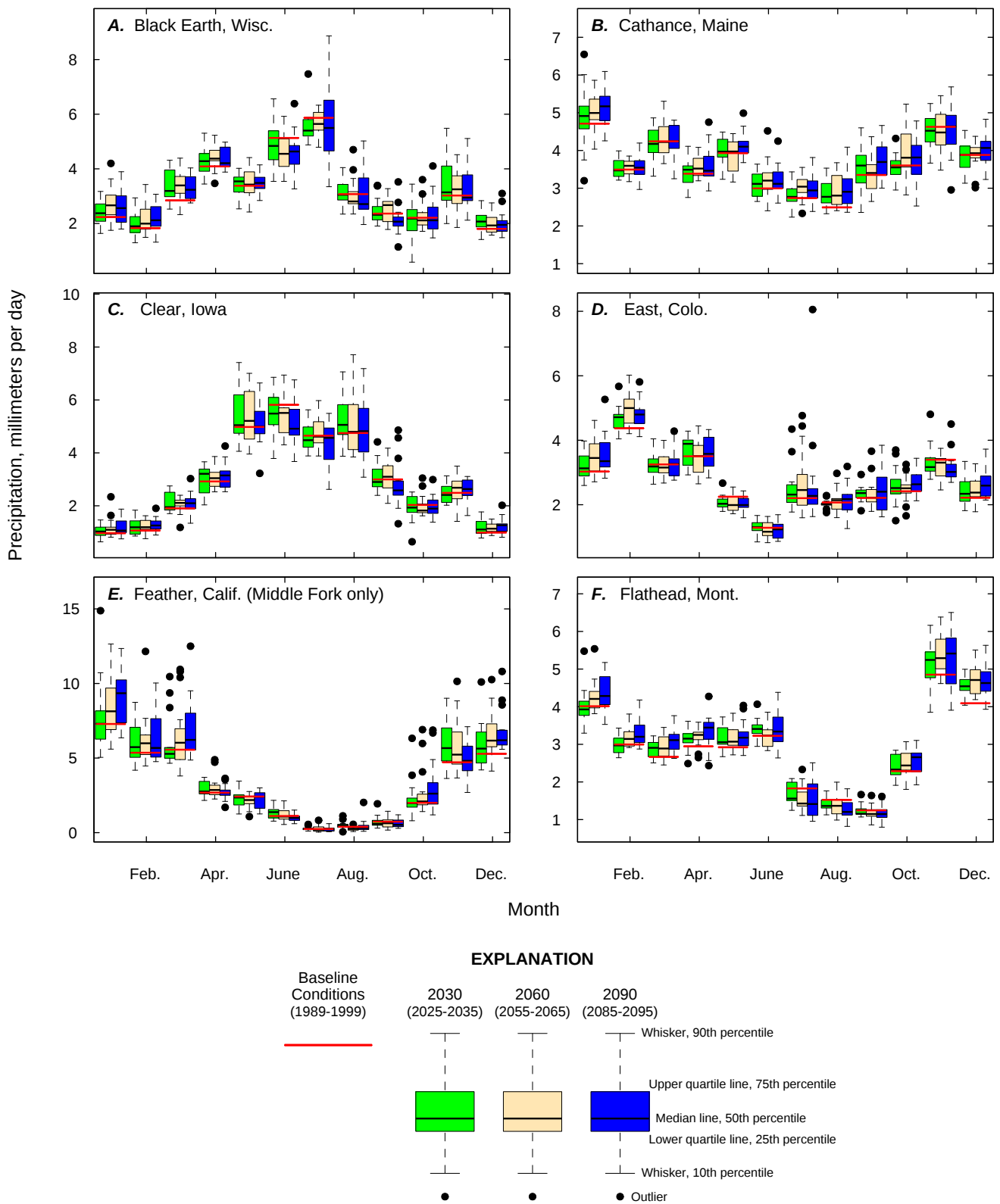


Figure 29. Mean daily precipitation values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

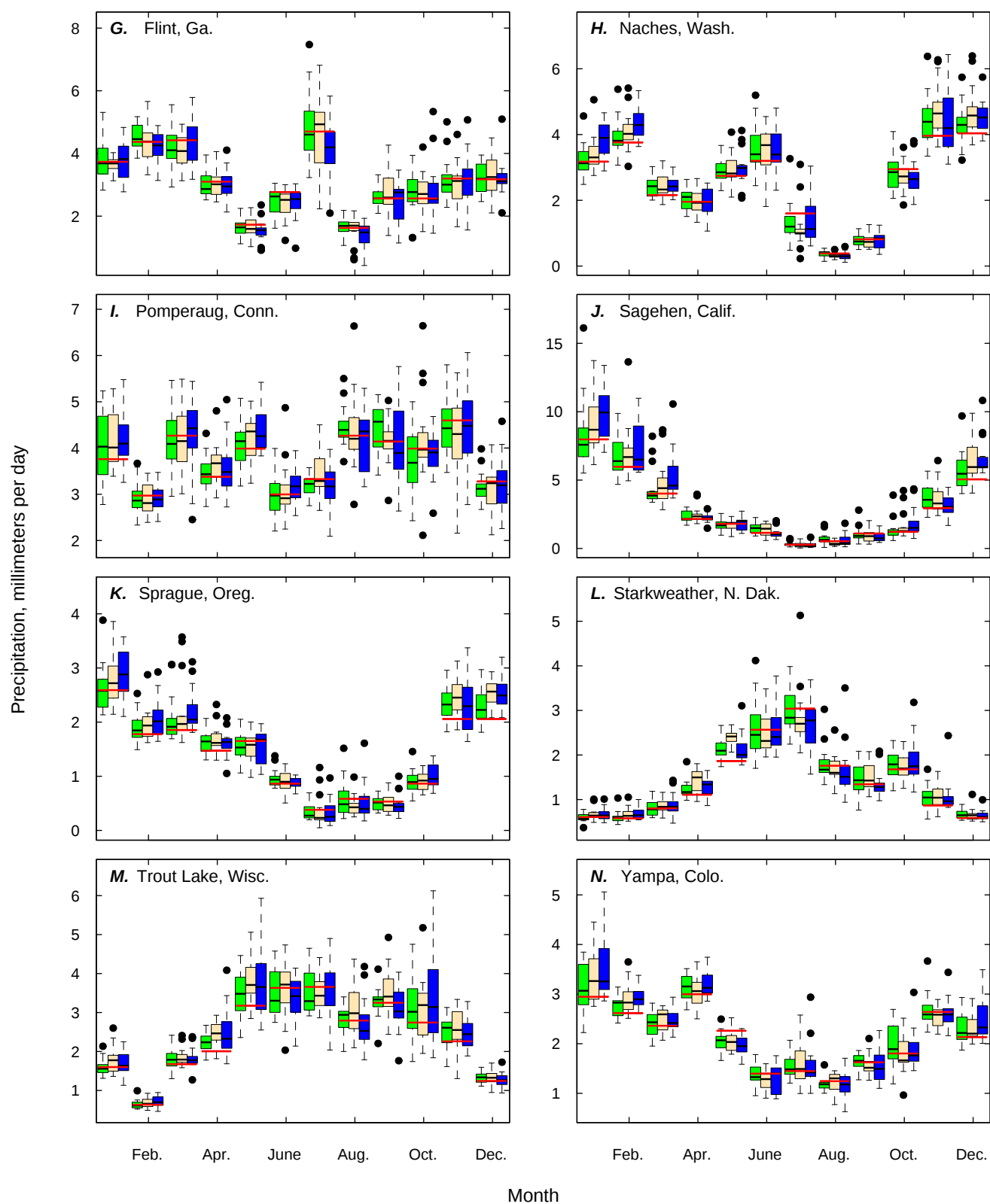


Figure 29. Mean daily precipitation values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

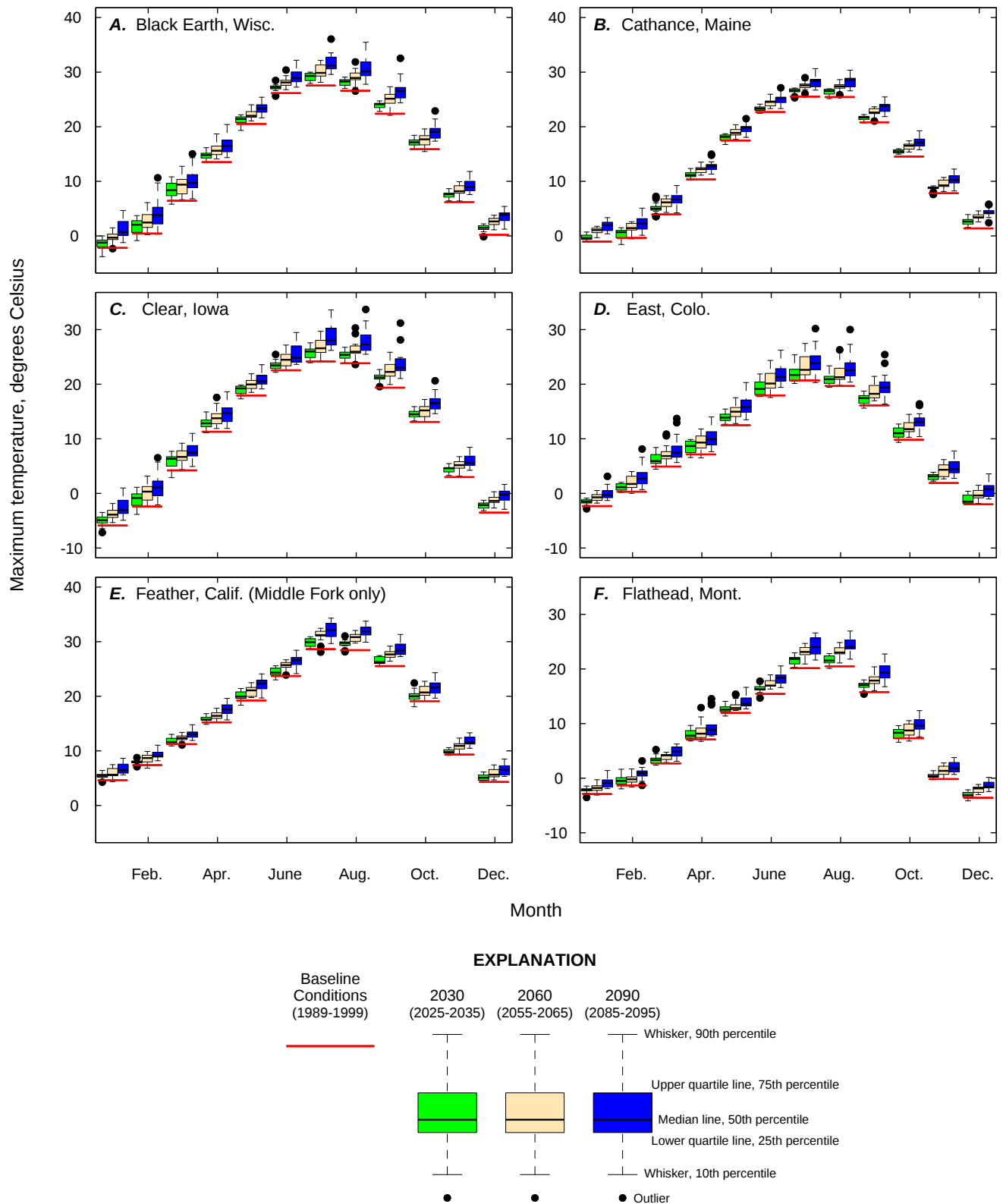


Figure 30. Mean daily maximum temperature values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

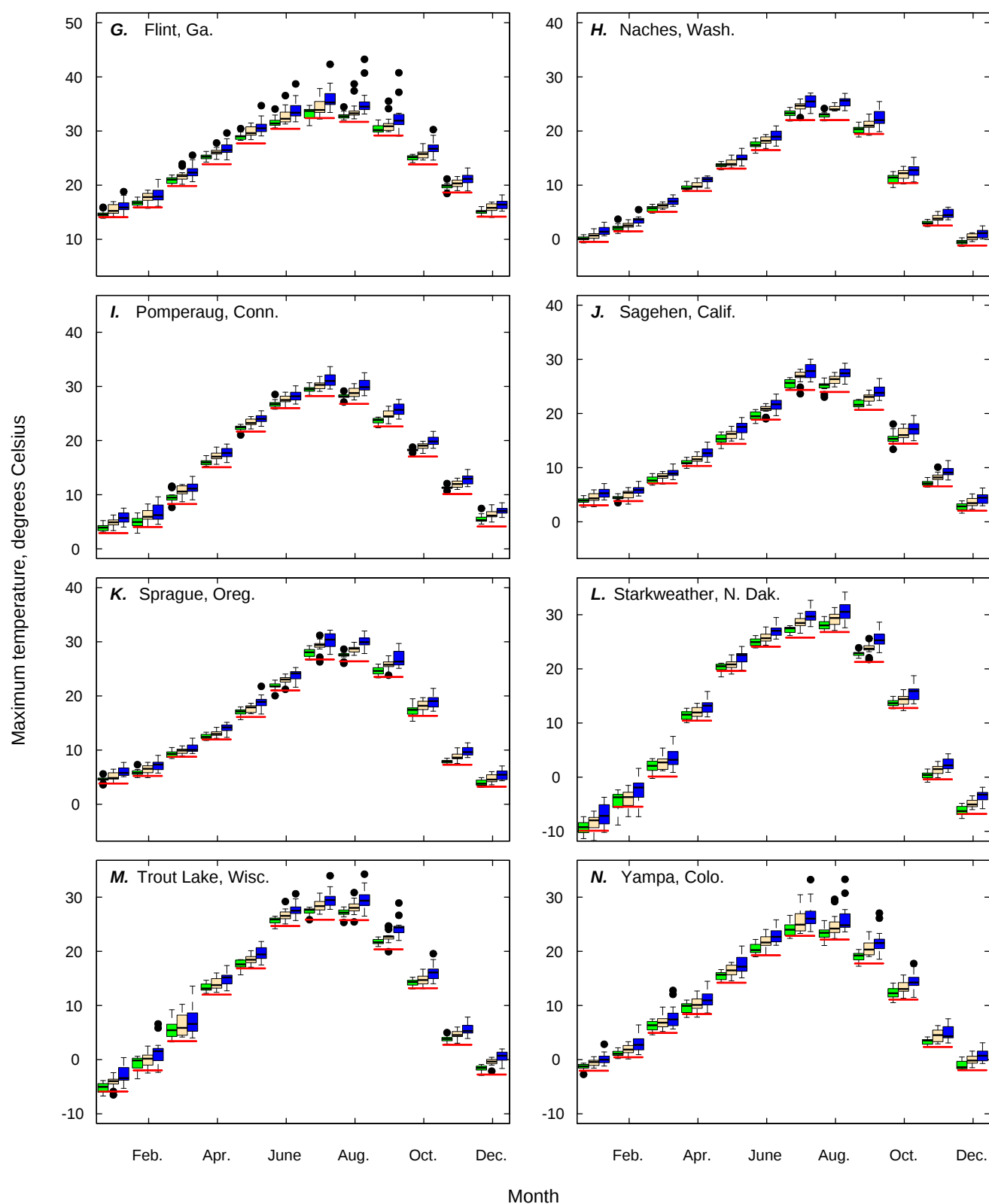


Figure 30. Mean daily maximum temperature values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

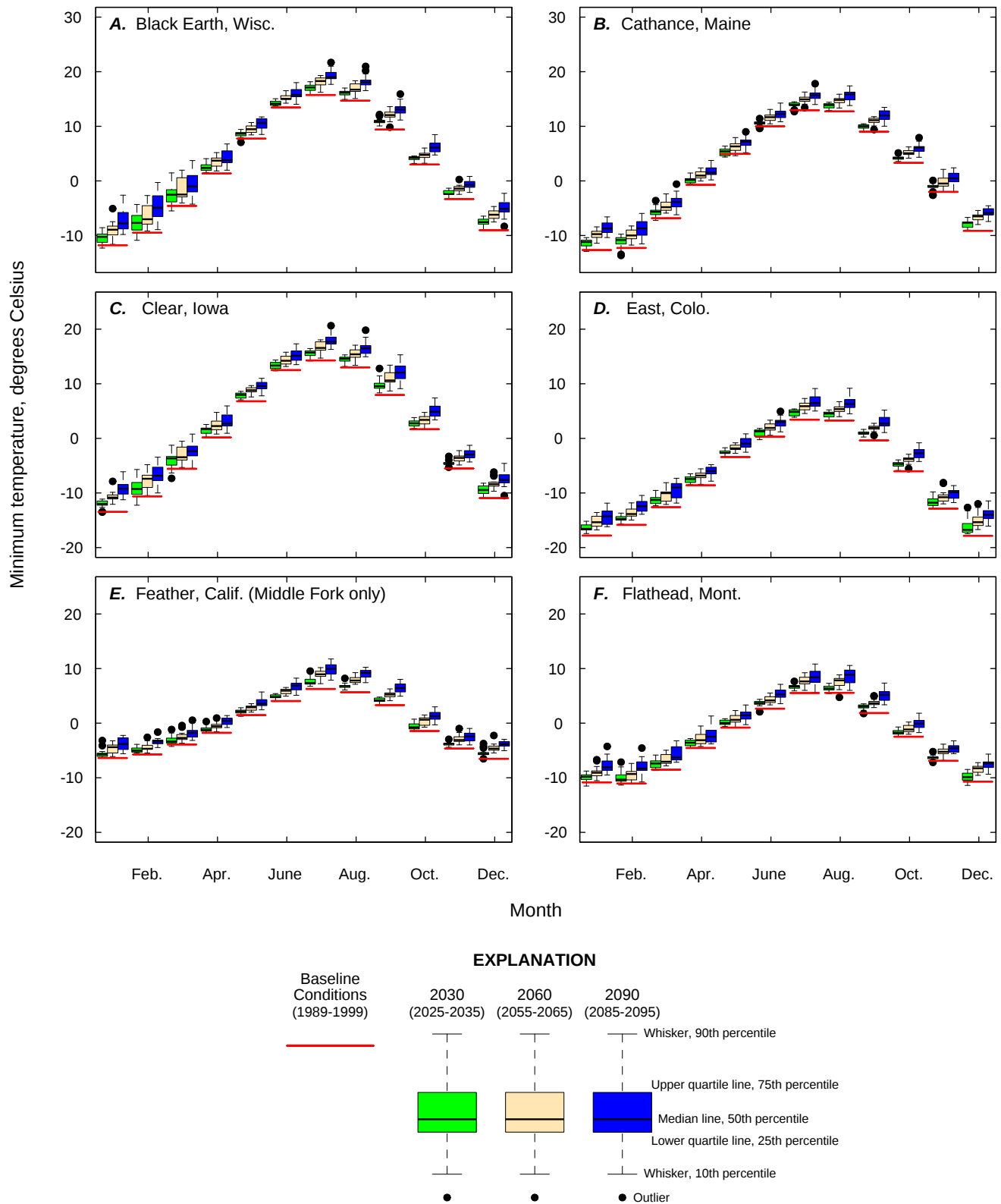


Figure 31. Mean daily minimum temperature values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

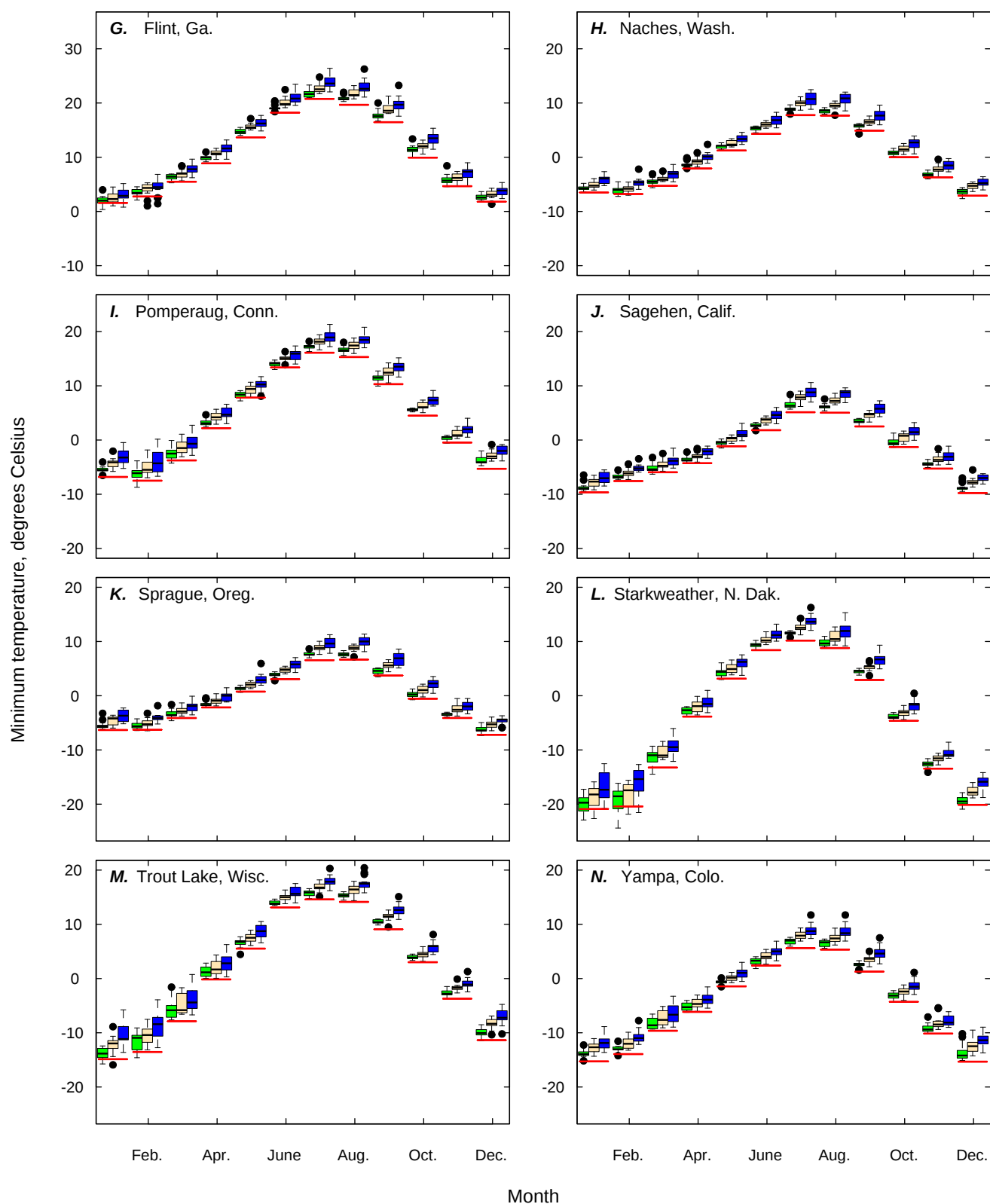


Figure 31. Mean daily minimum temperature values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

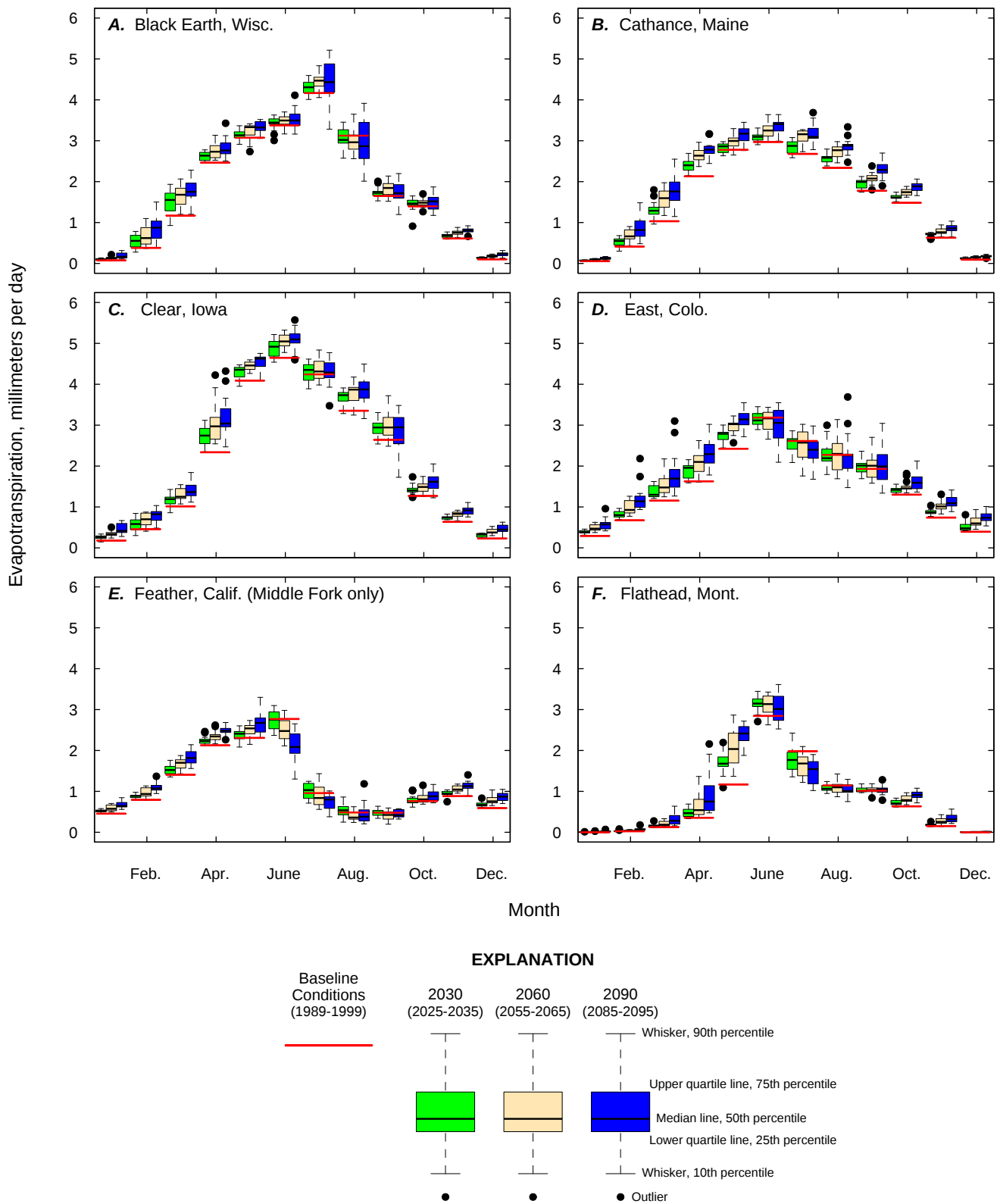


Figure 32. Mean daily evapotranspiration values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

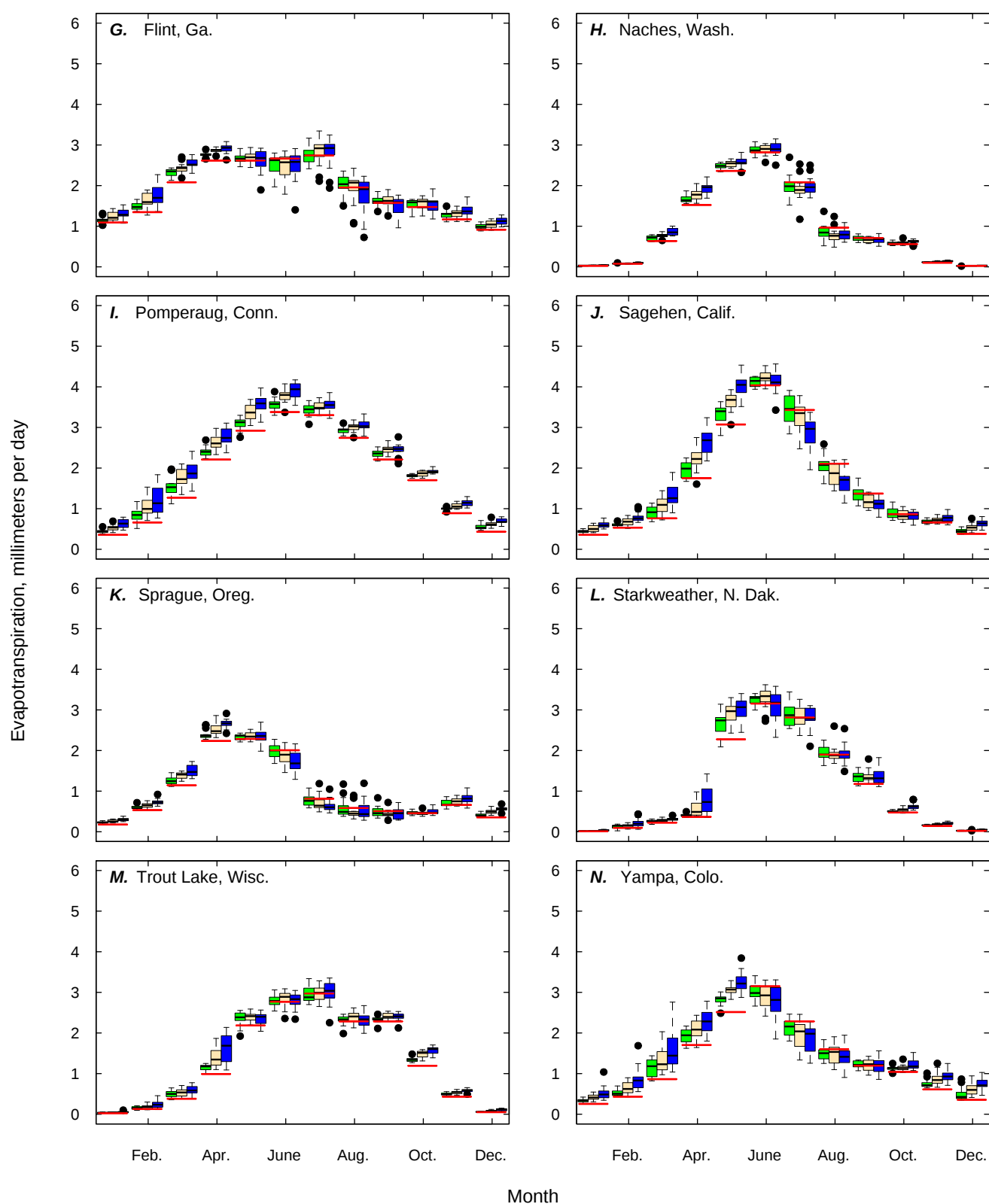


Figure 32. Mean daily evapotranspiration values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

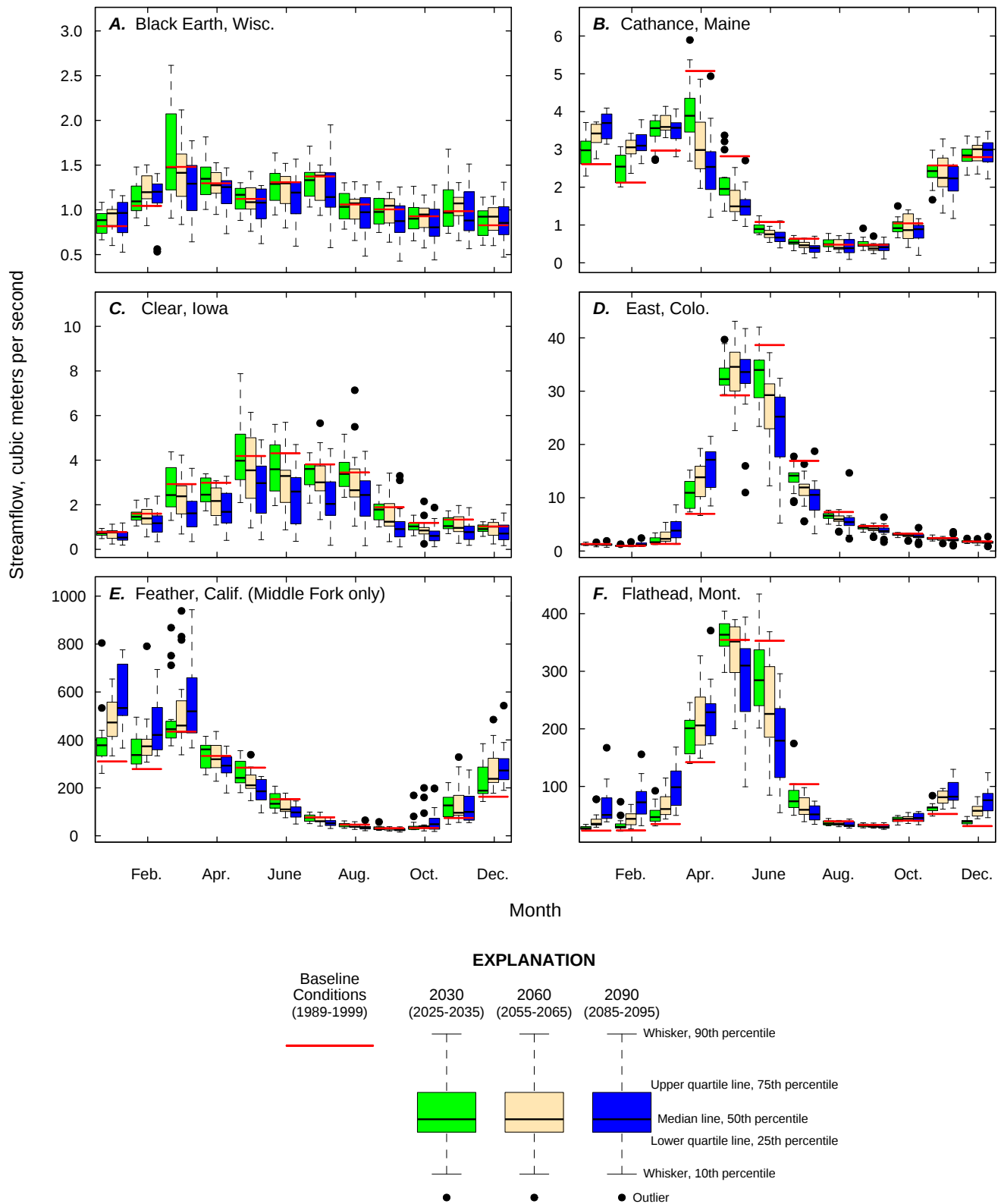


Figure 33. Mean daily streamflow values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

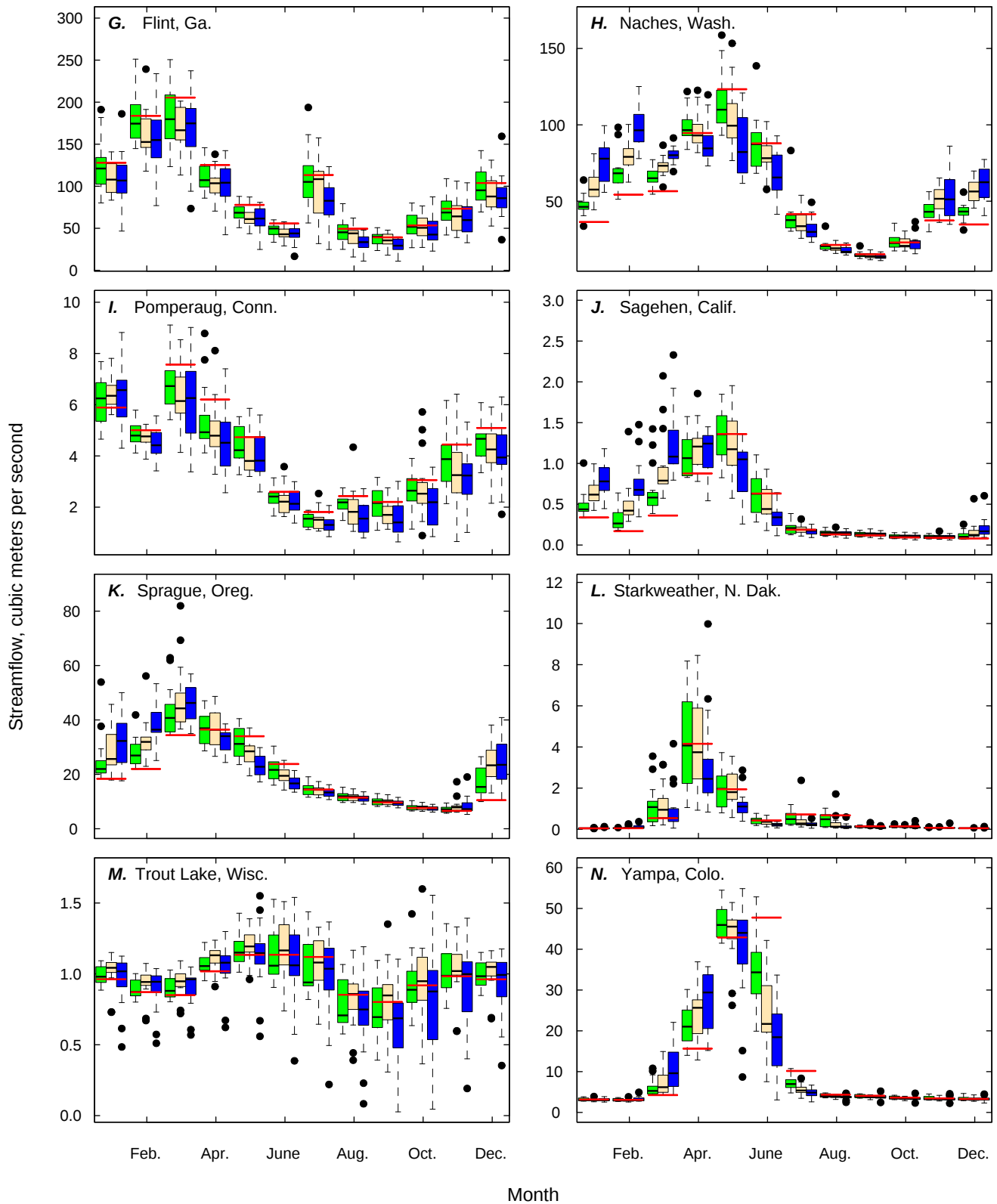


Figure 33. Mean daily streamflow values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

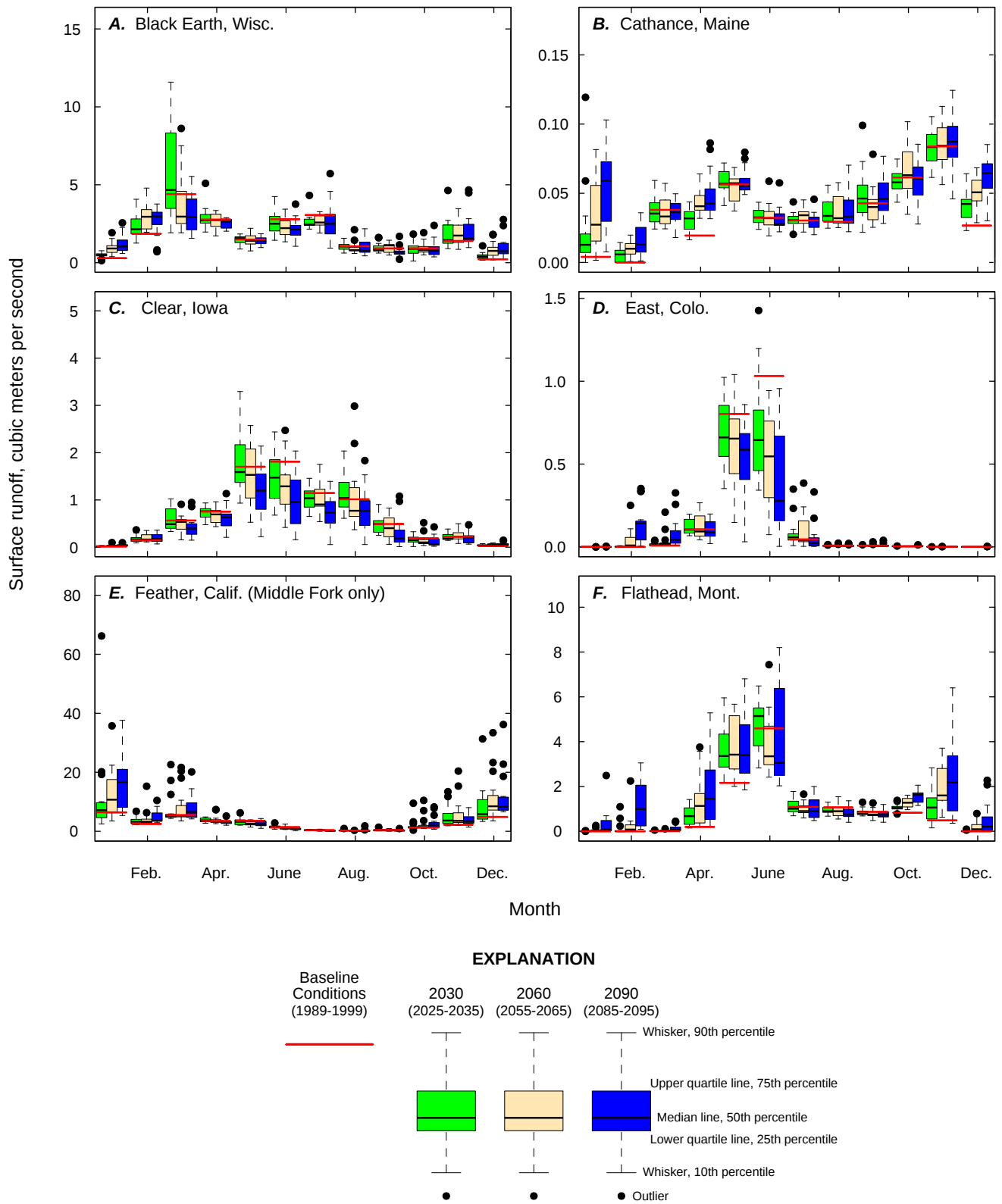


Figure 34. Mean daily surface runoff values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

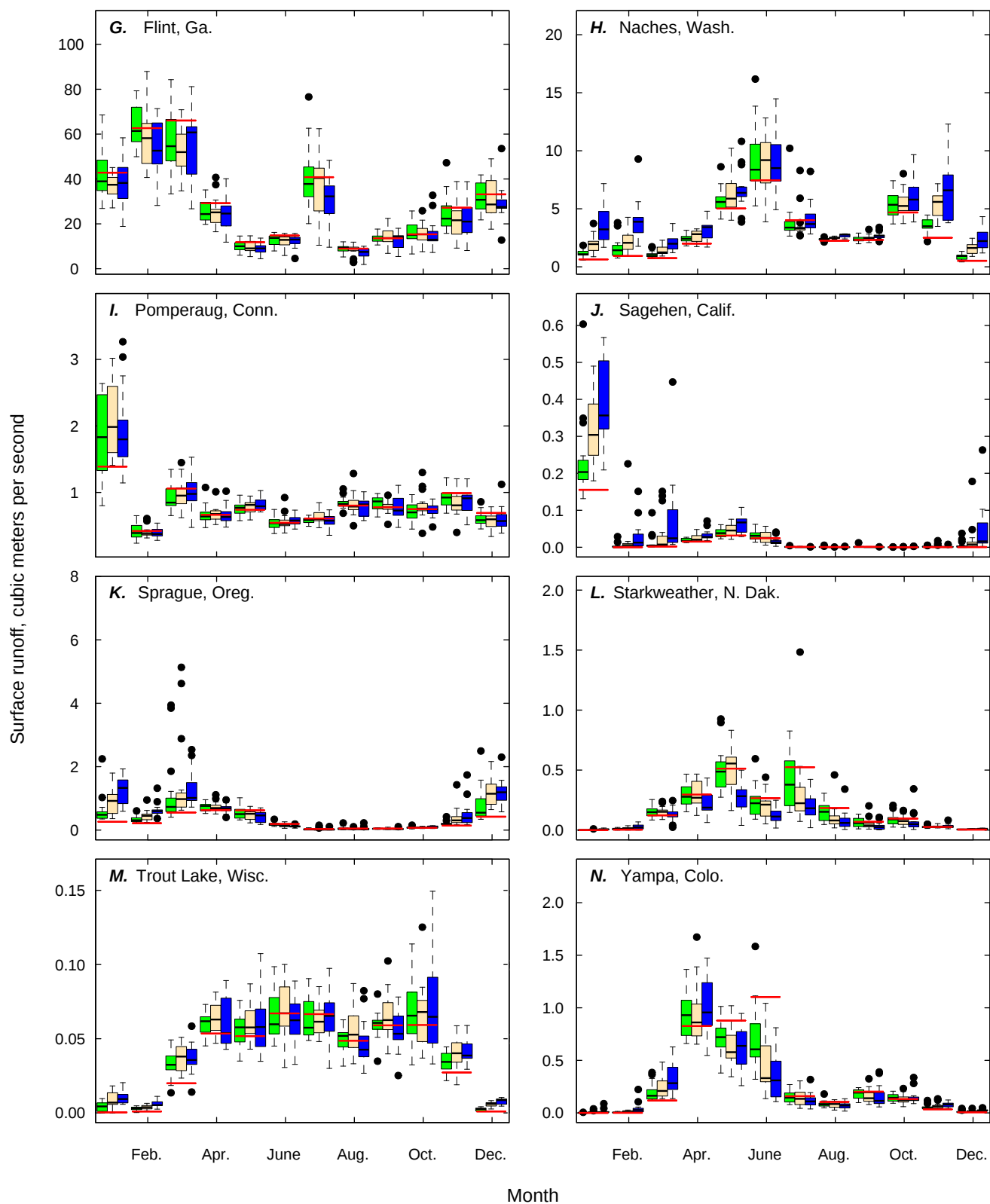


Figure 34. Mean daily surface runoff values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

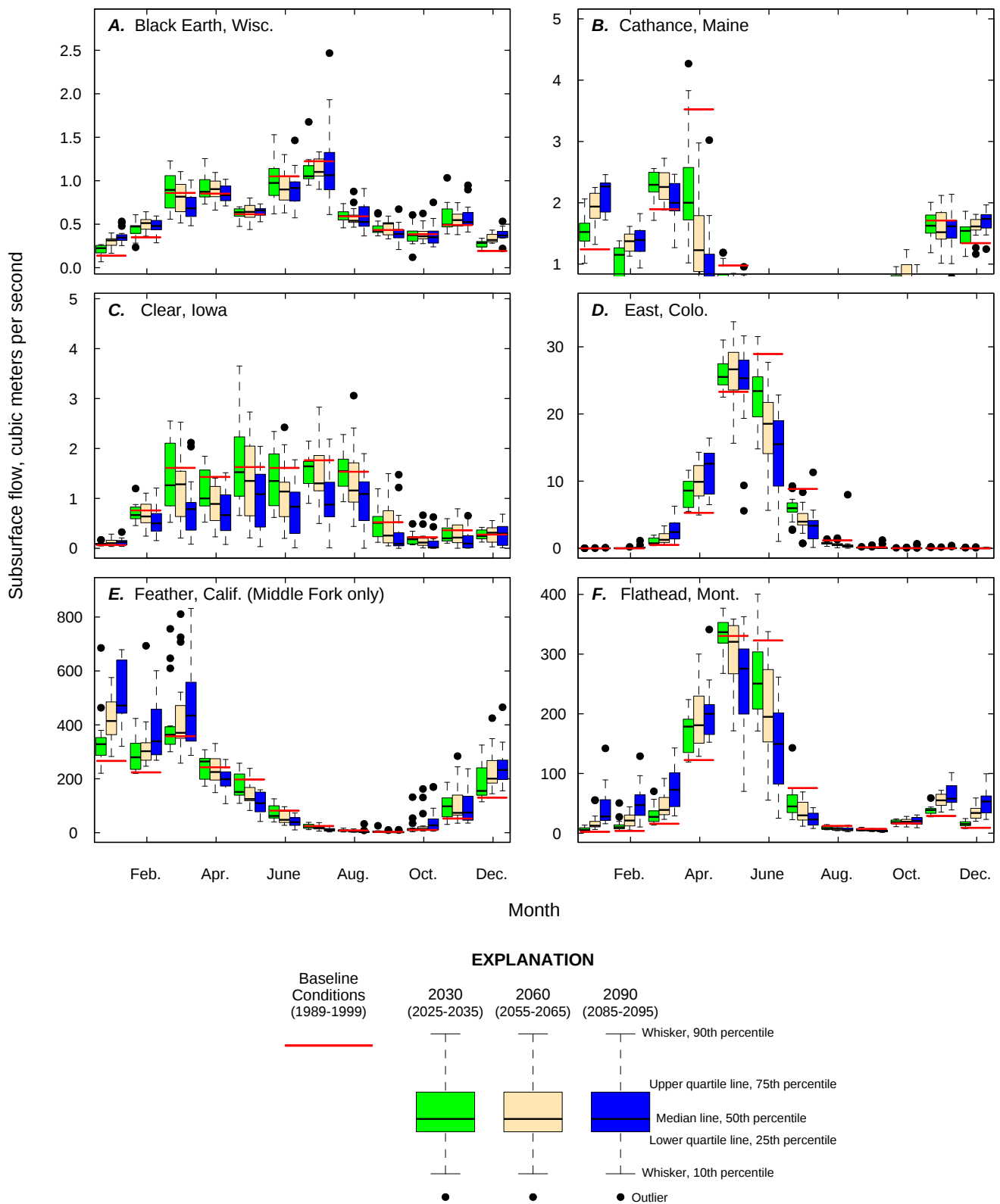


Figure 35. Mean daily subsurface flow values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

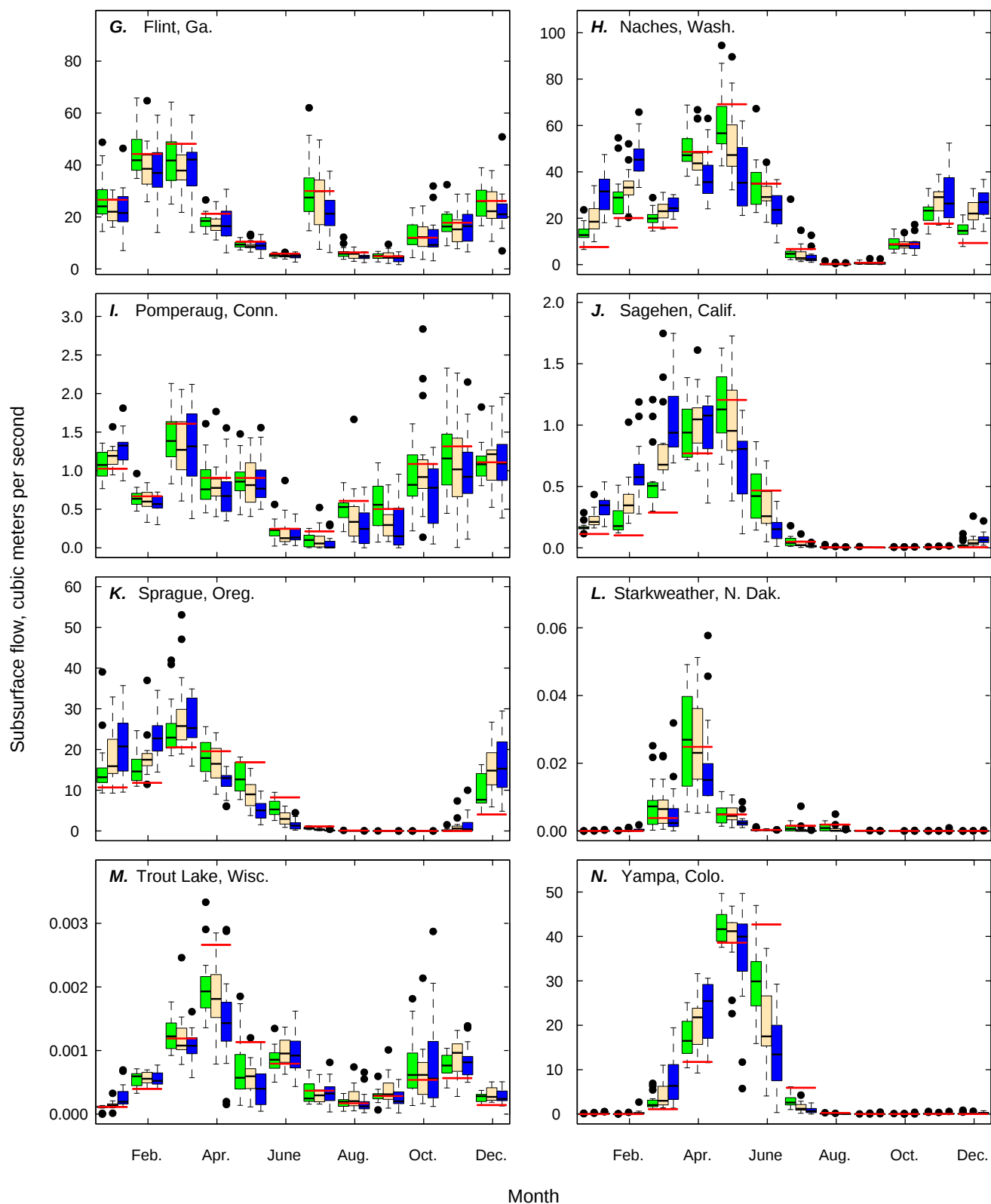


Figure 35. Mean daily subsurface flow values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

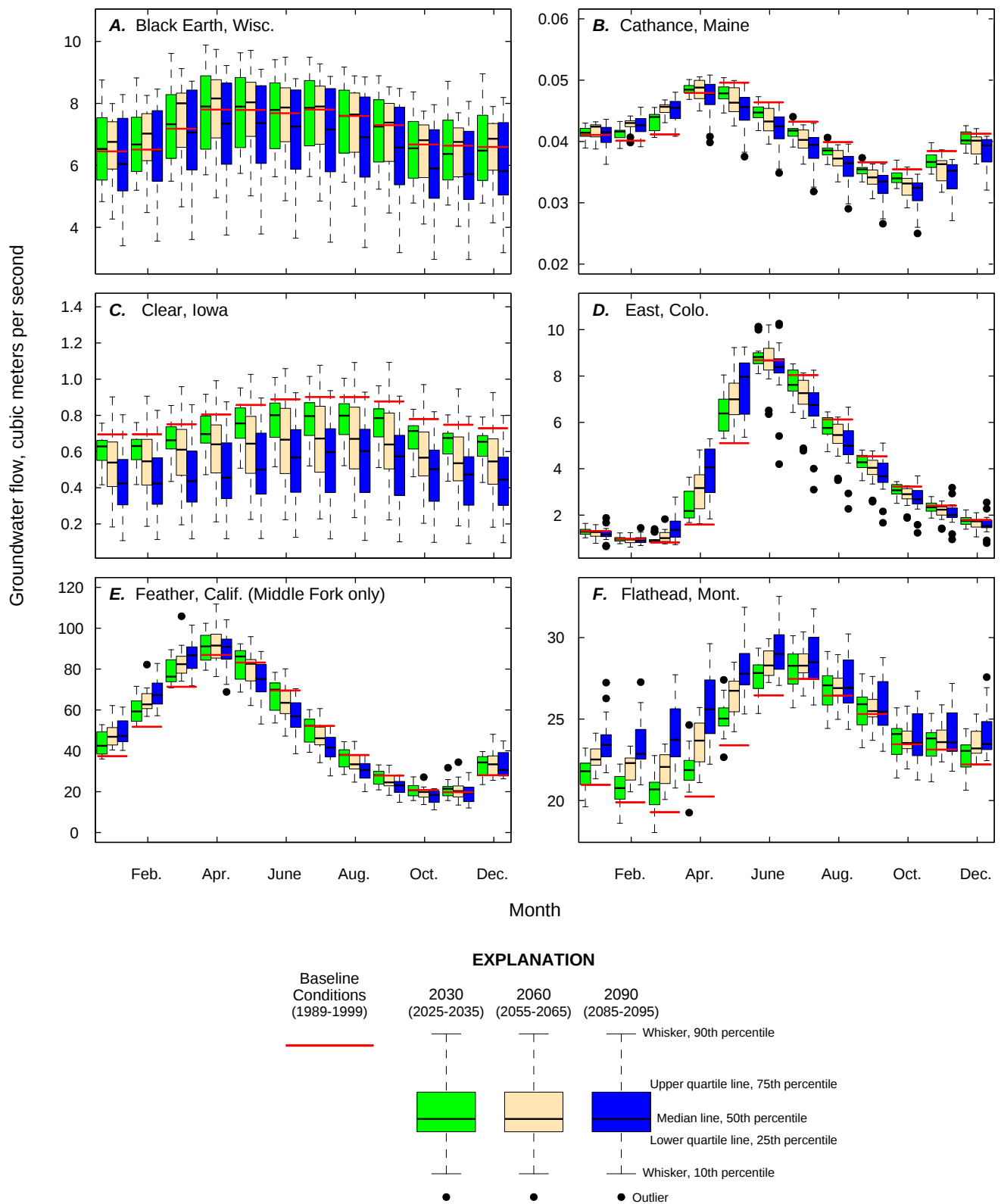


Figure 36. Mean daily groundwater flow values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

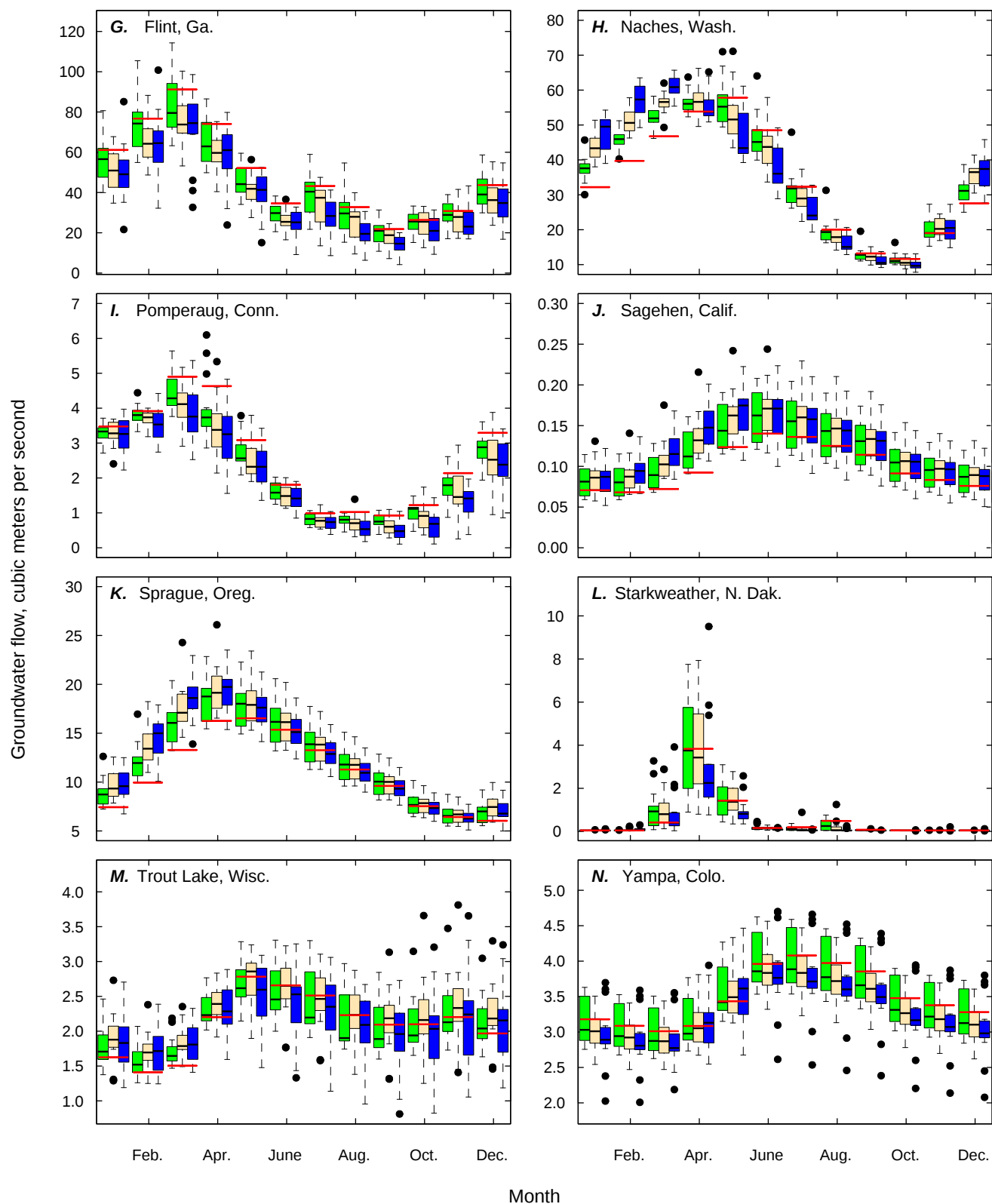


Figure 36. Mean daily groundwater flow values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

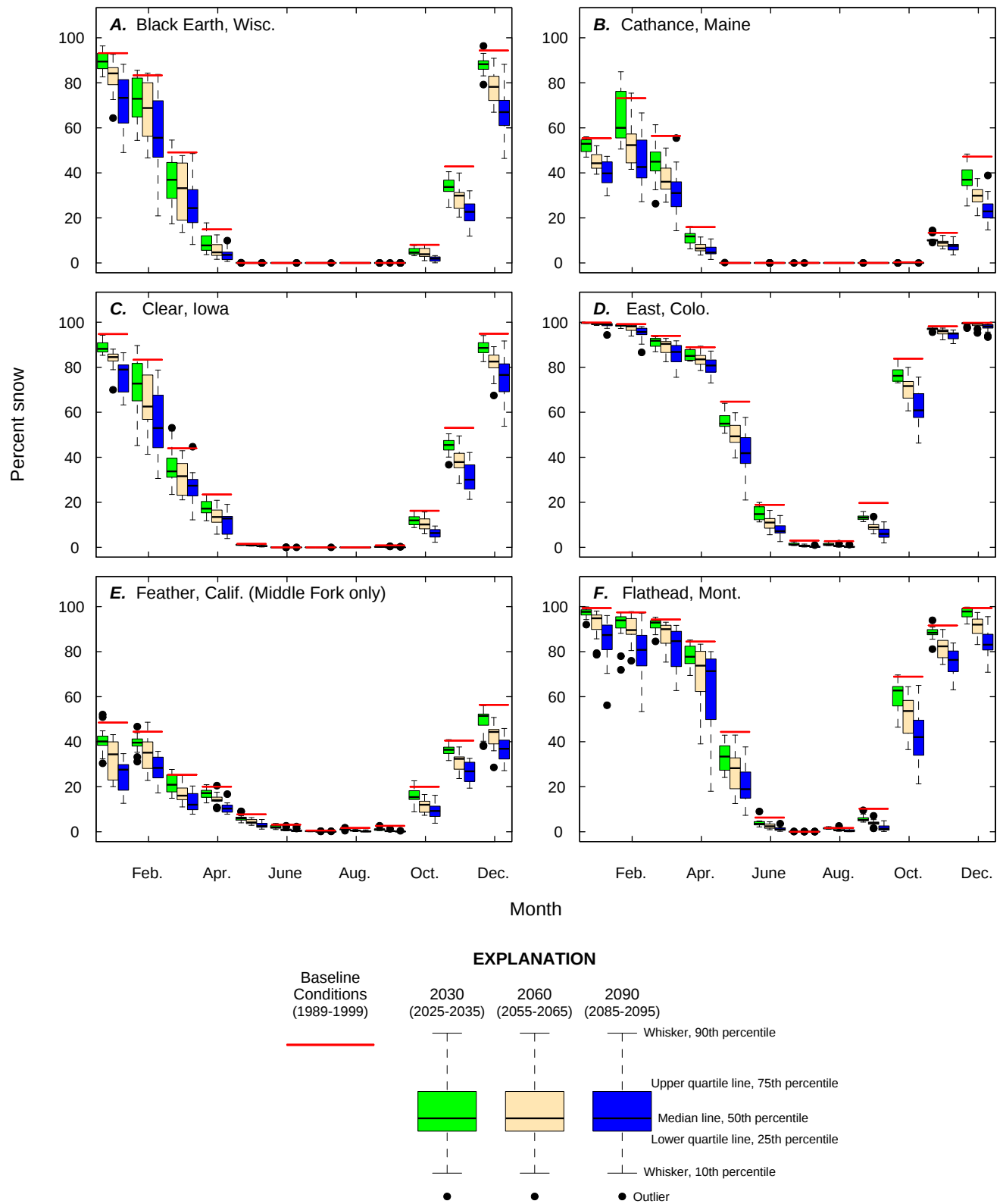


Figure 37. Mean daily percent of precipitation that falls as snow by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

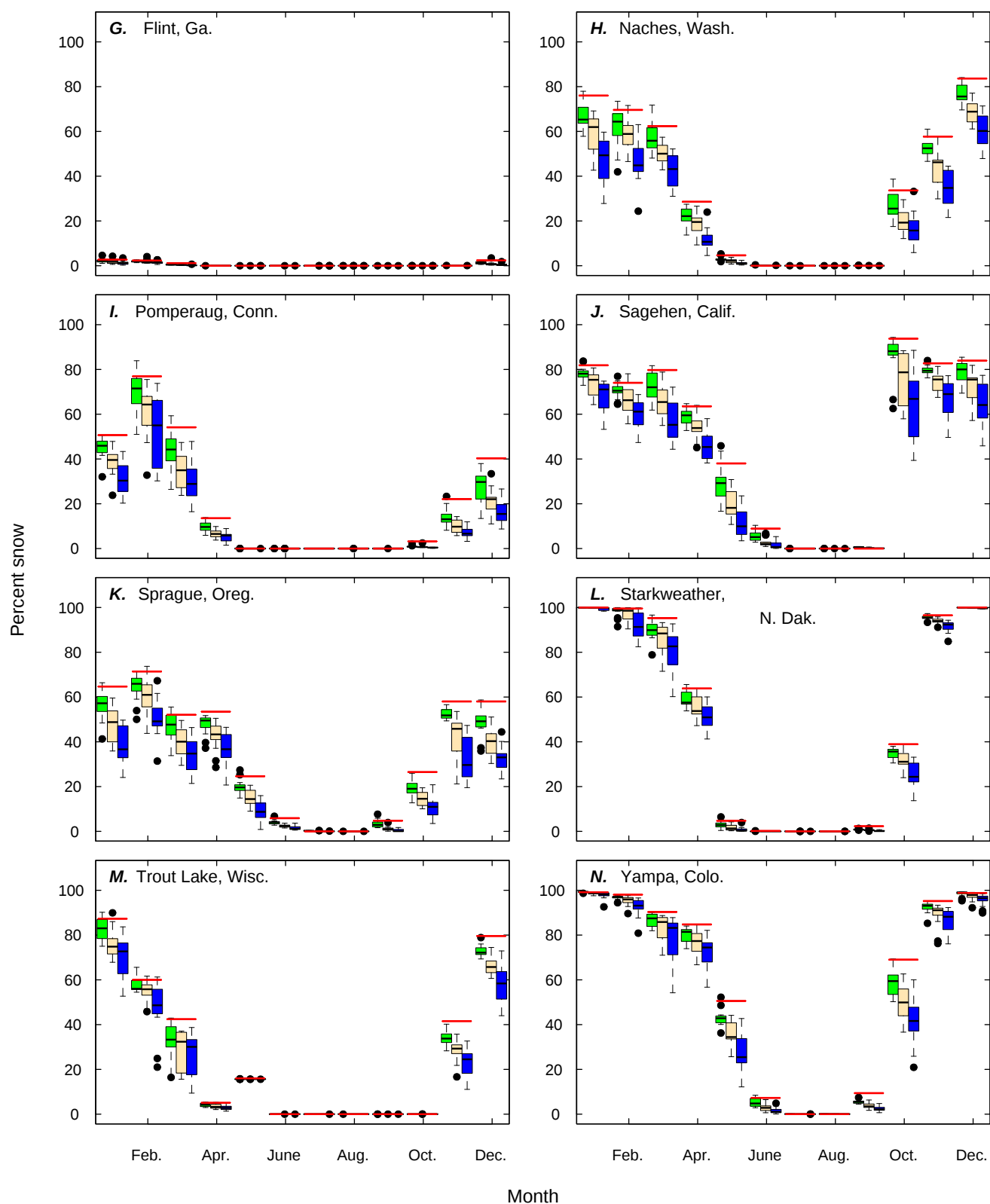


Figure 37. Mean daily percent of precipitation that falls as snow by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

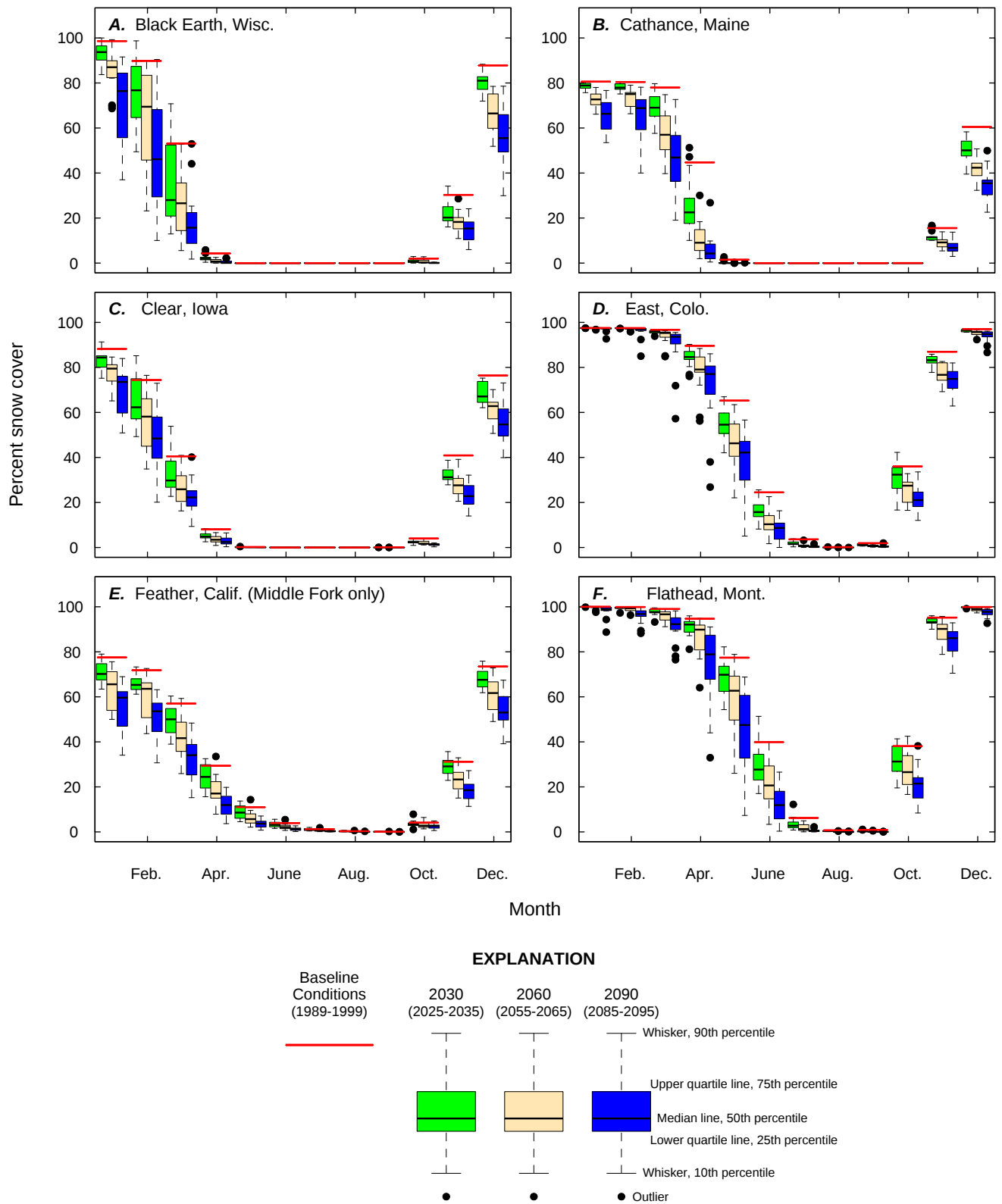


Figure 38. Mean daily percent of area covered by snow by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

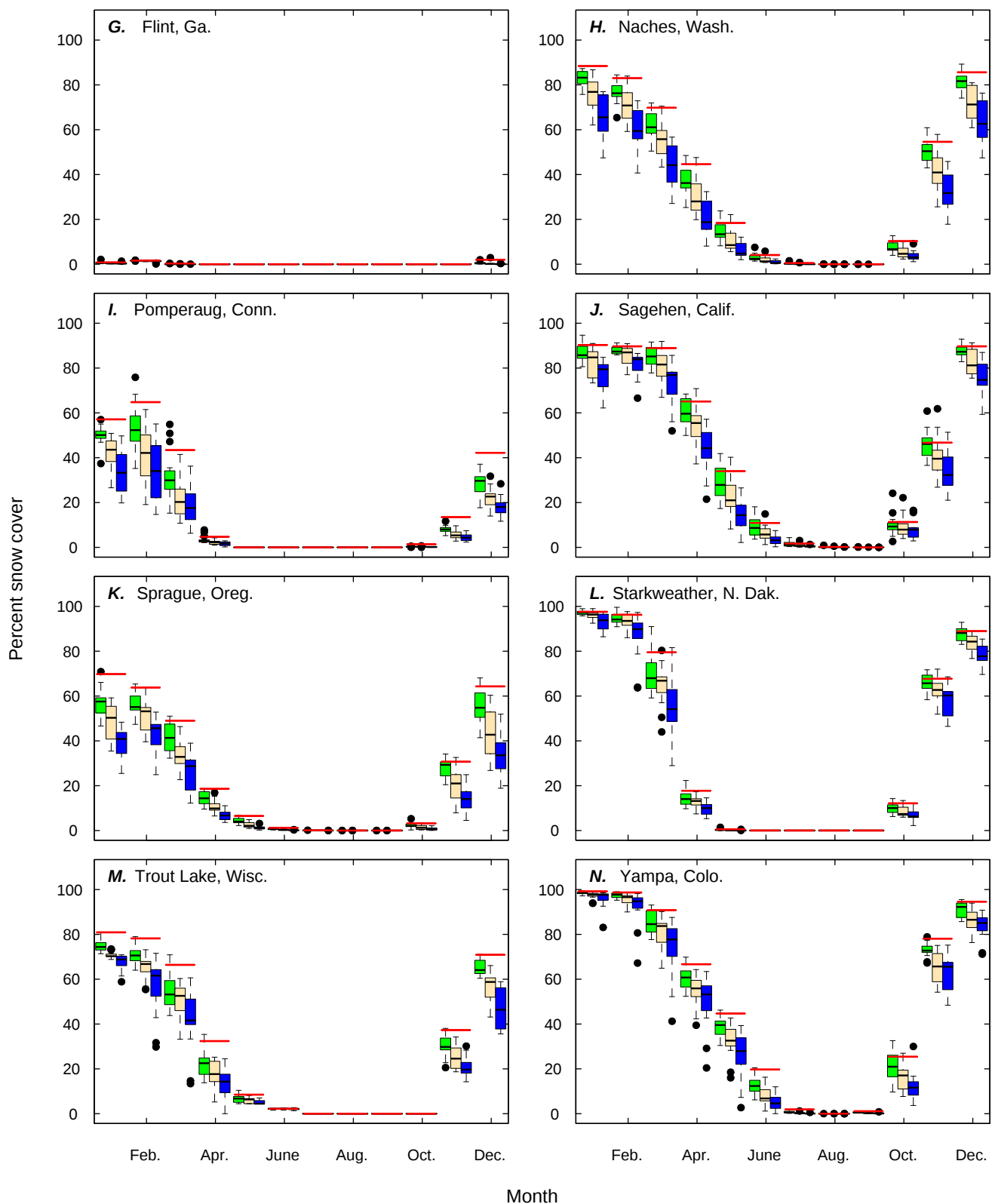


Figure 38. Mean daily percent of area covered by snow by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

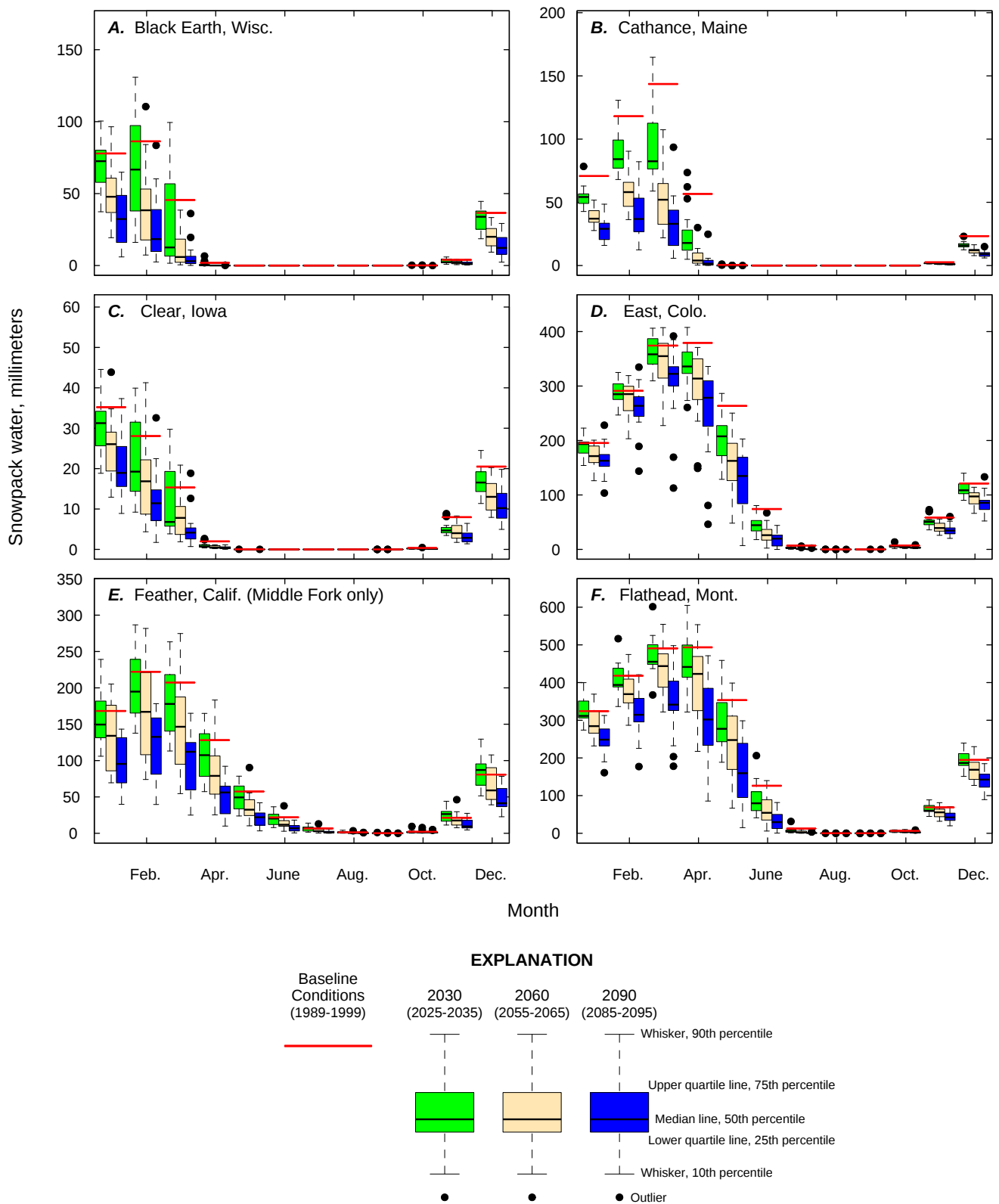


Figure 39. Mean daily snowpack water equivalent by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

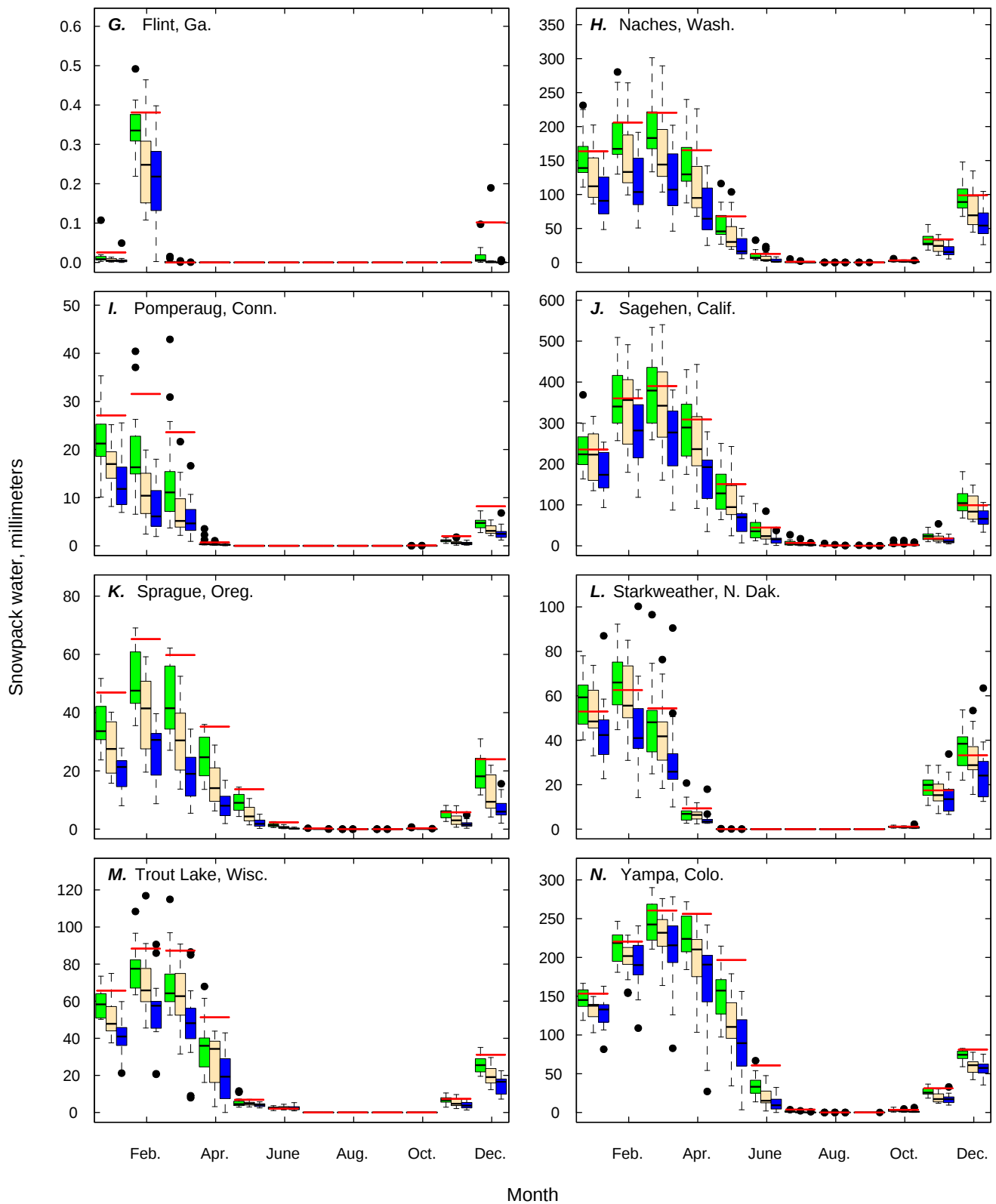


Figure 39. Mean daily snowpack water equivalent by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

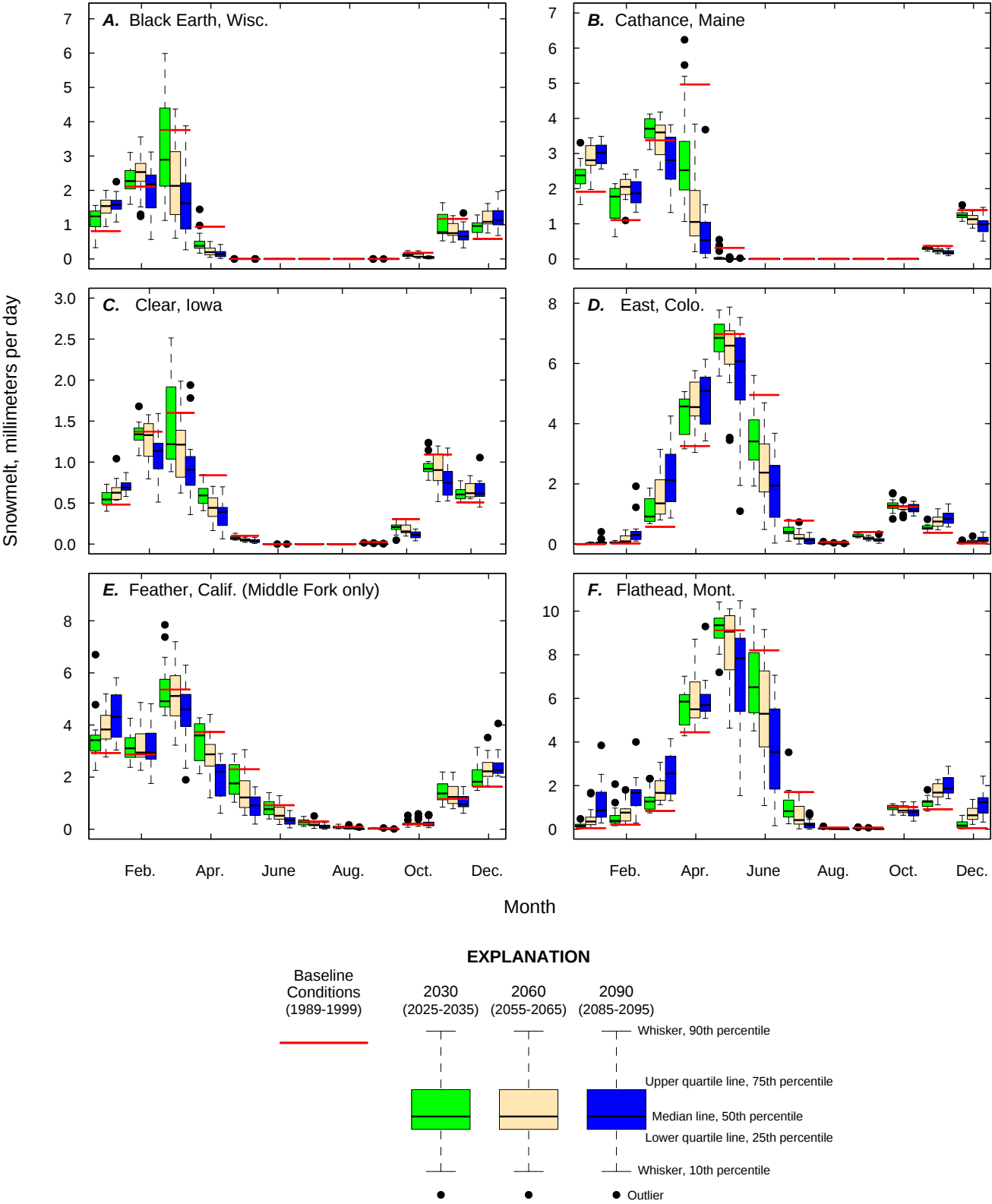


Figure 40. Mean daily snowmelt values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

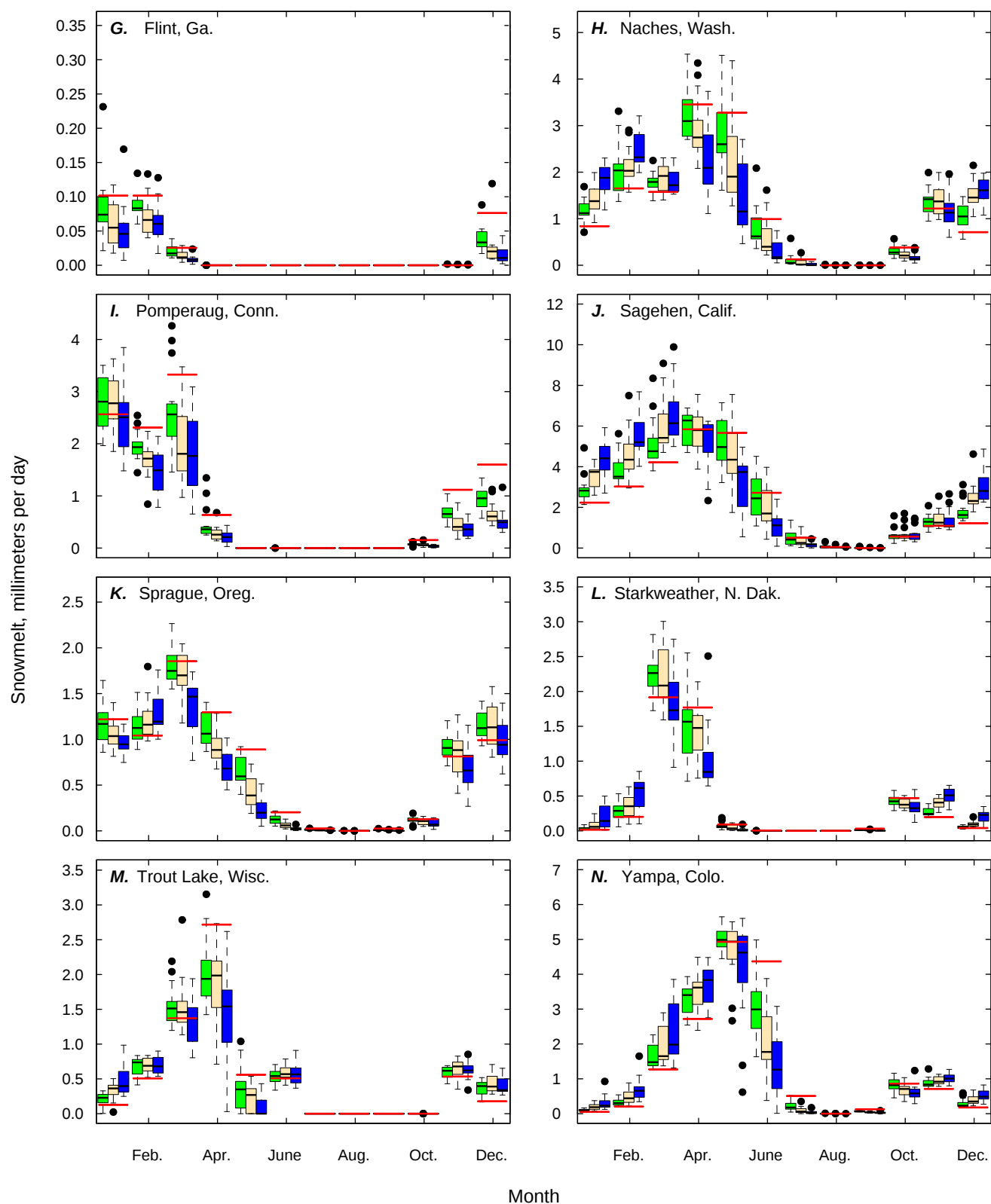


Figure 40. Mean daily snowmelt values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

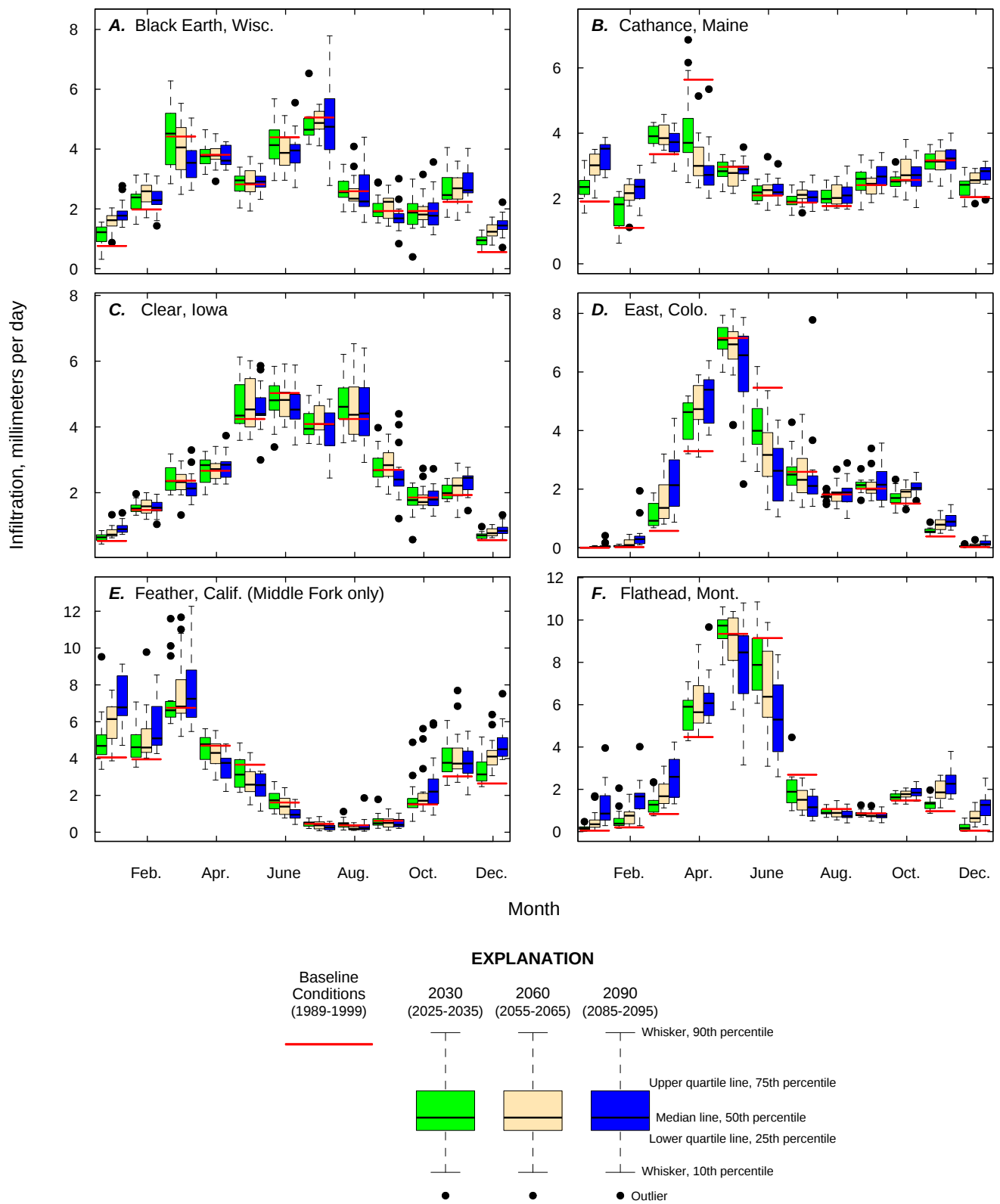


Figure 41. Mean daily infiltration values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

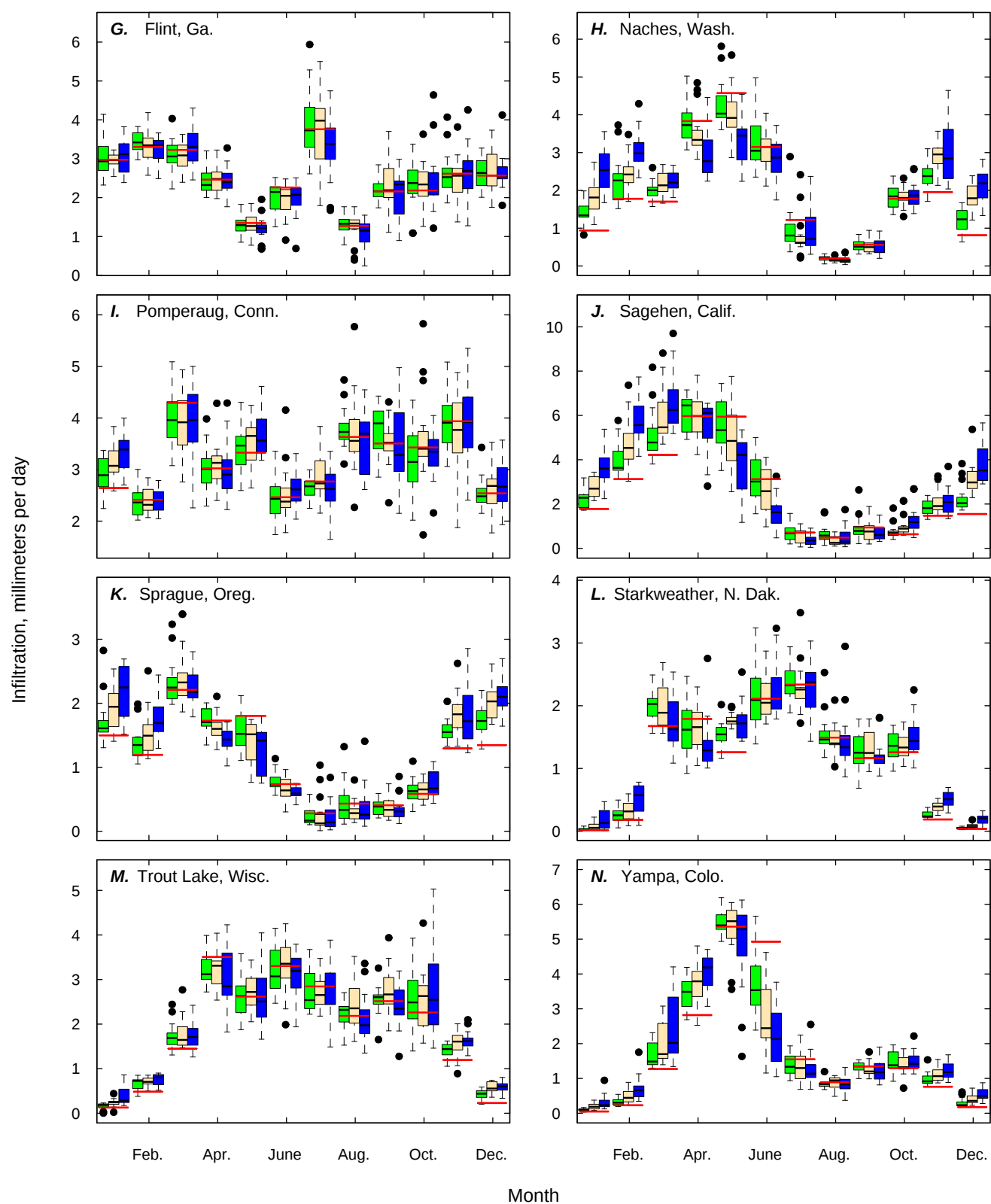


Figure 41. Mean daily infiltration values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

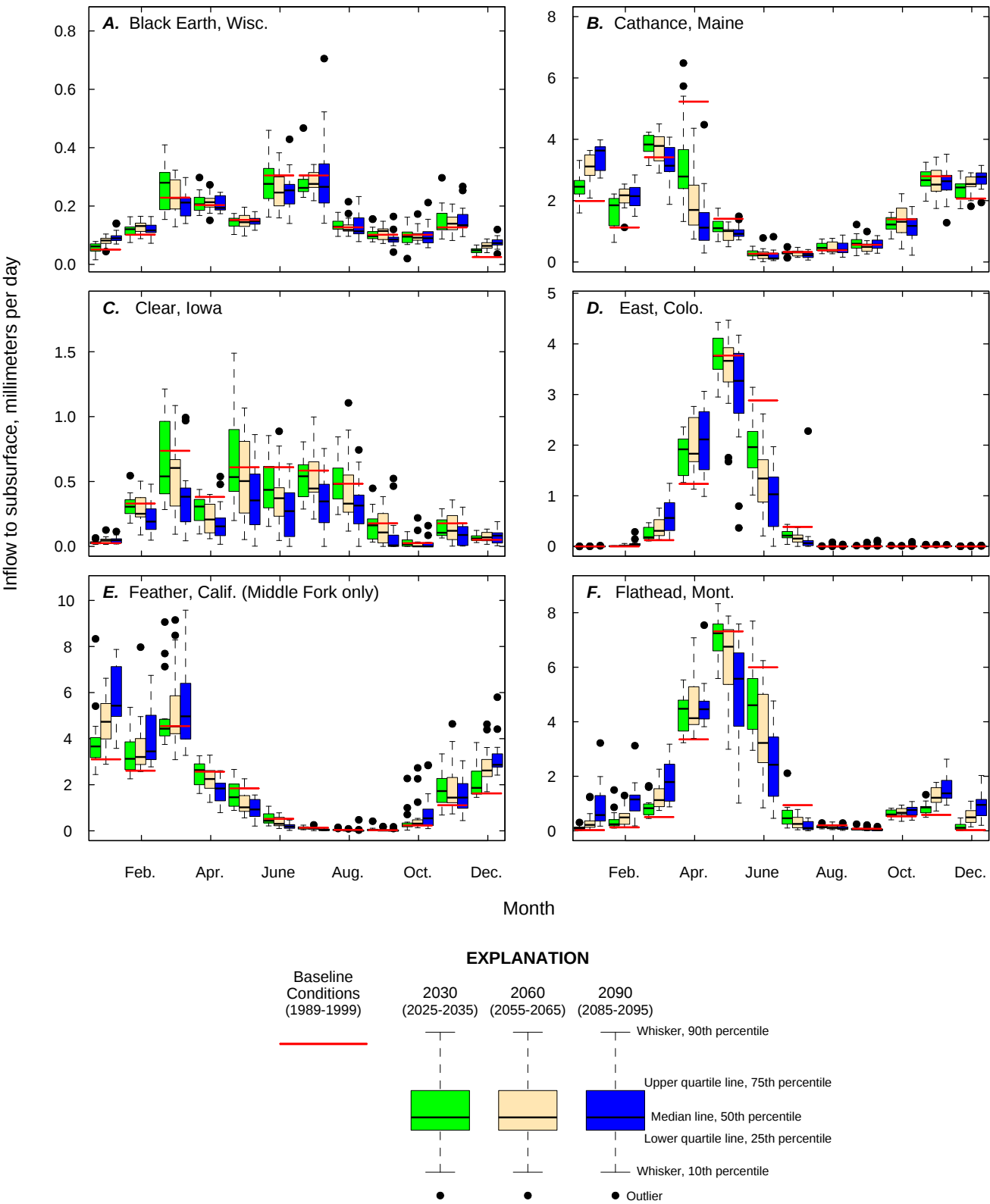


Figure 42. Mean daily subsurface reservoir inflow values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

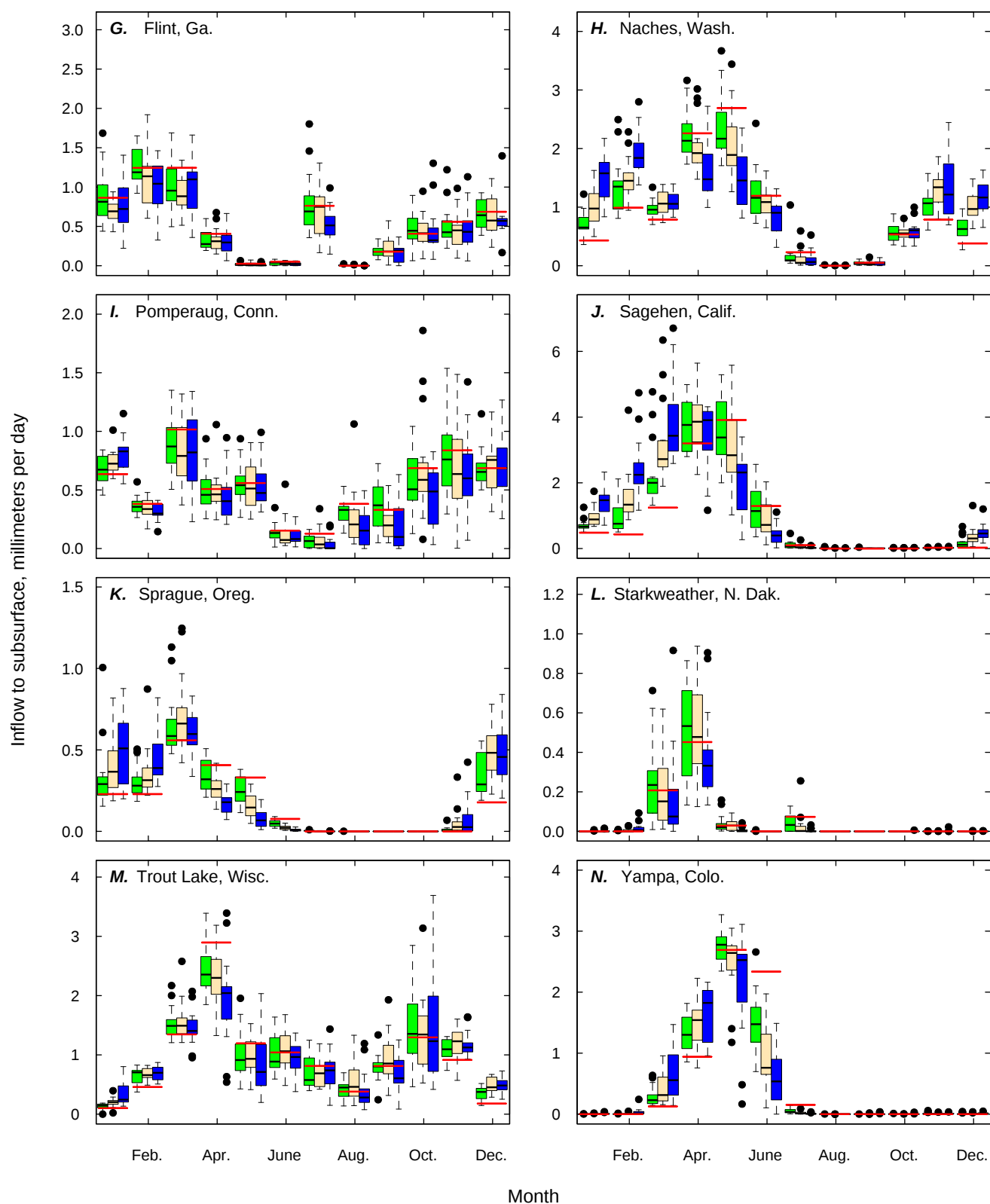


Figure 42. Mean daily subsurface reservoir inflow values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

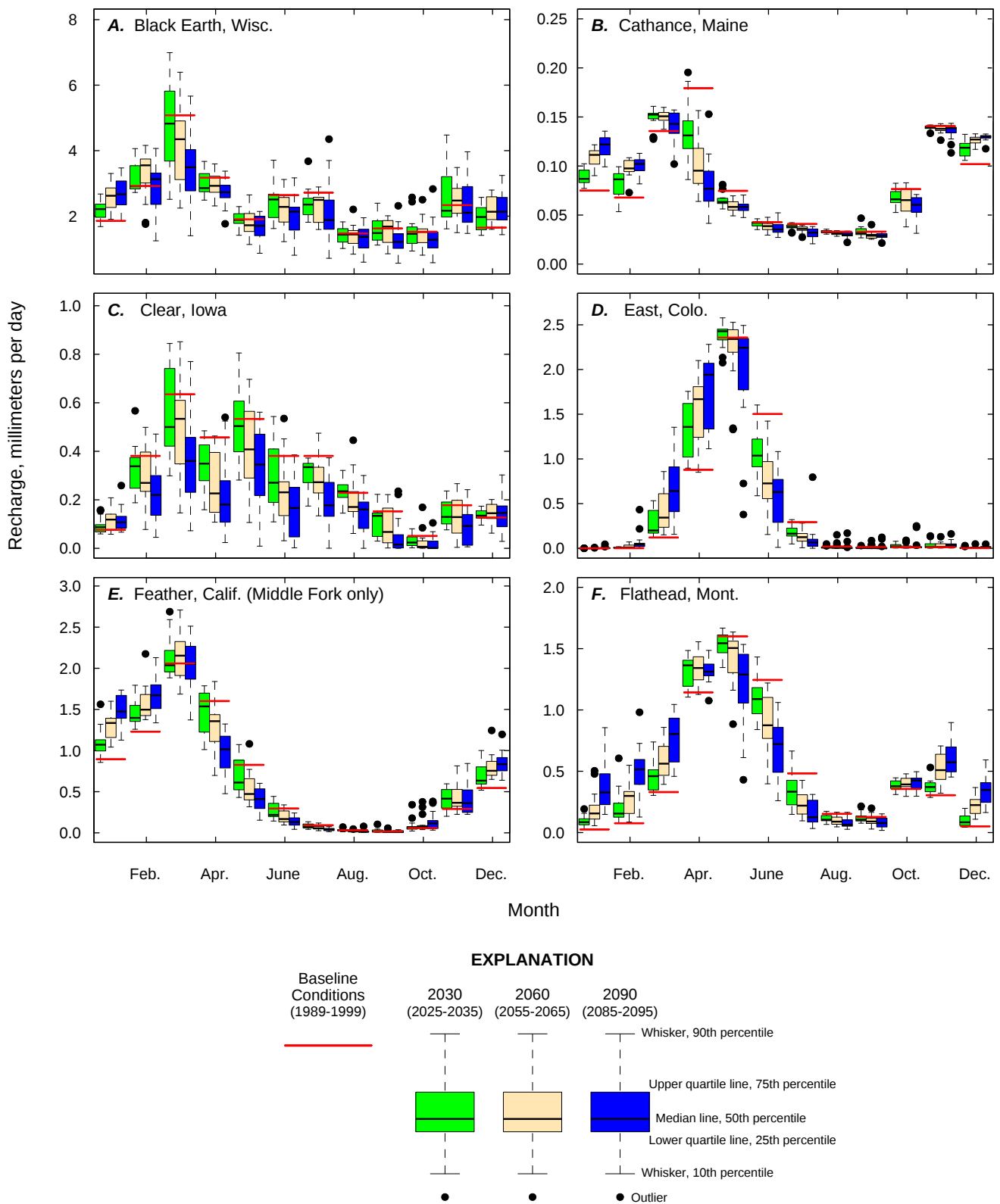


Figure 43. Mean daily groundwater reservoir recharge values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

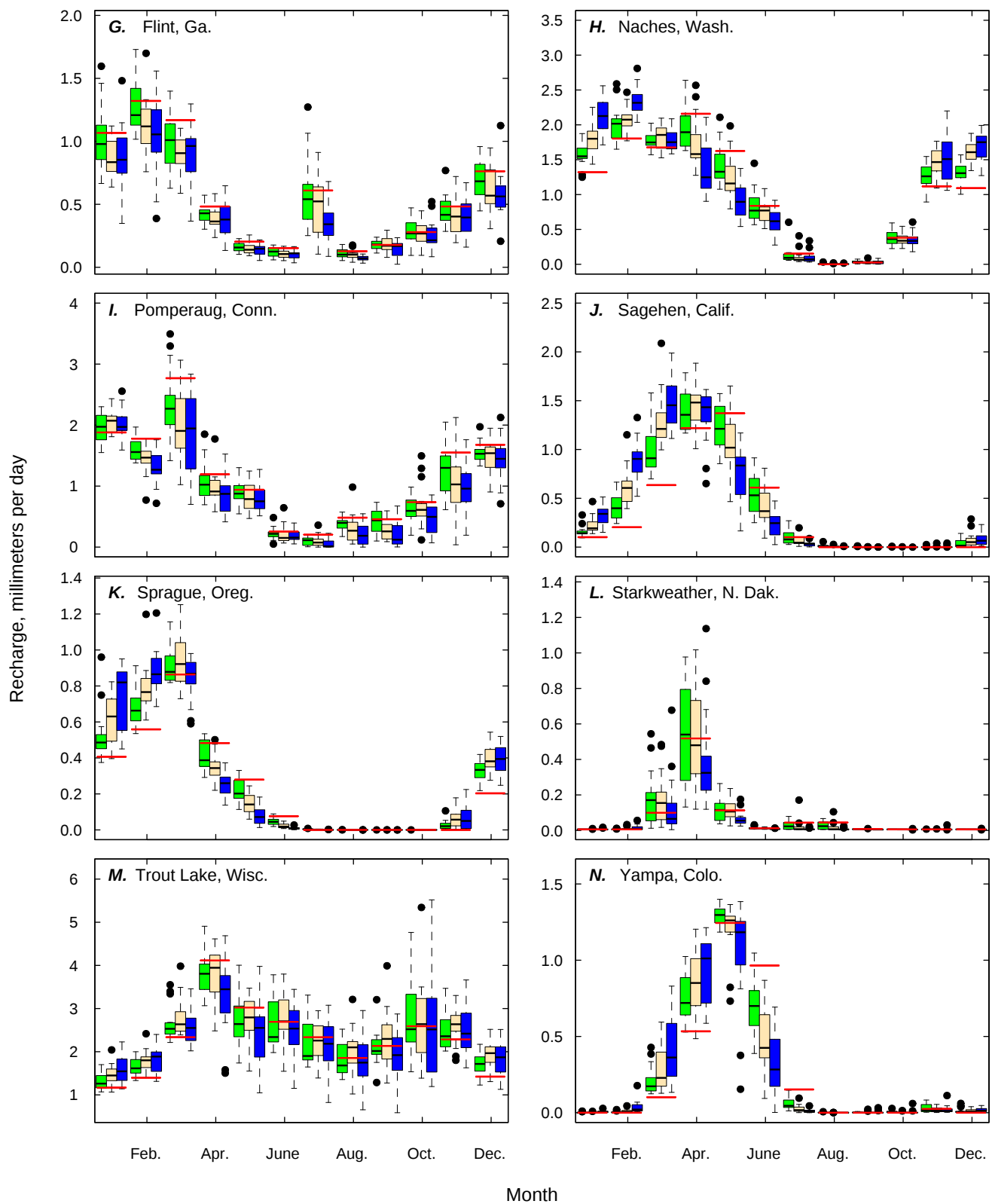


Figure 43. Mean daily groundwater reservoir recharge values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

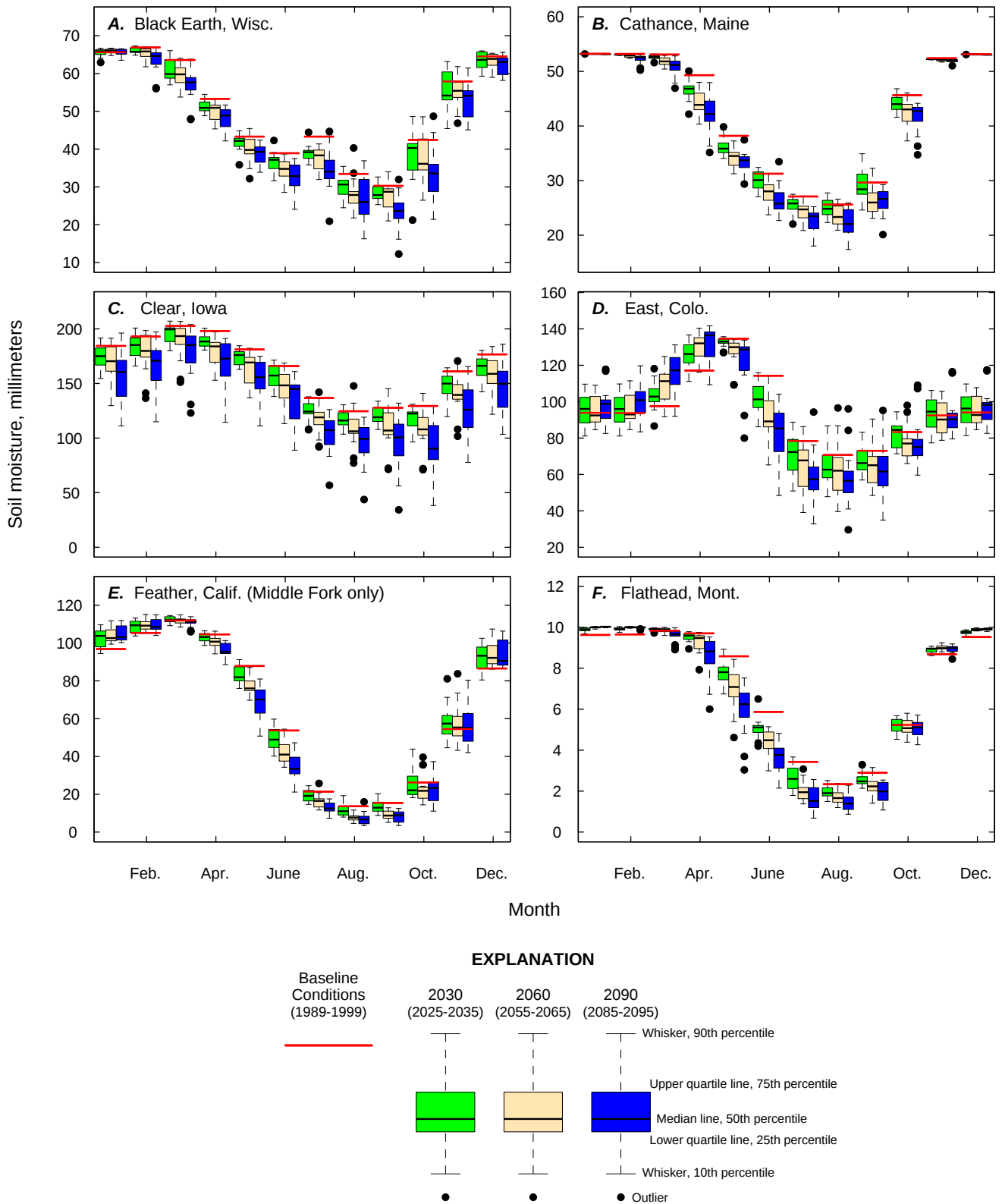


Figure 44. Mean daily soil moisture storage values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

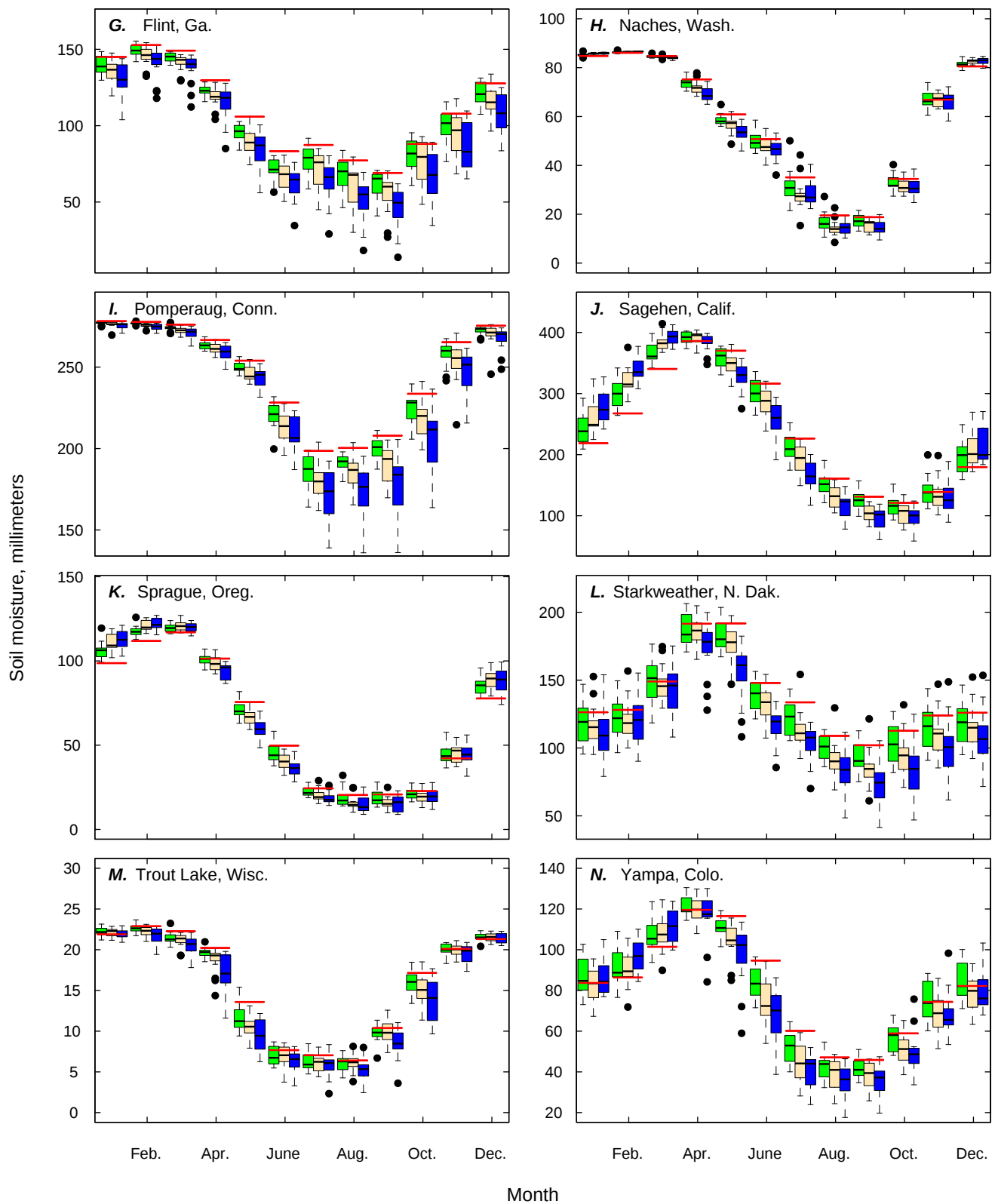


Figure 44. Mean daily soil moisture storage values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

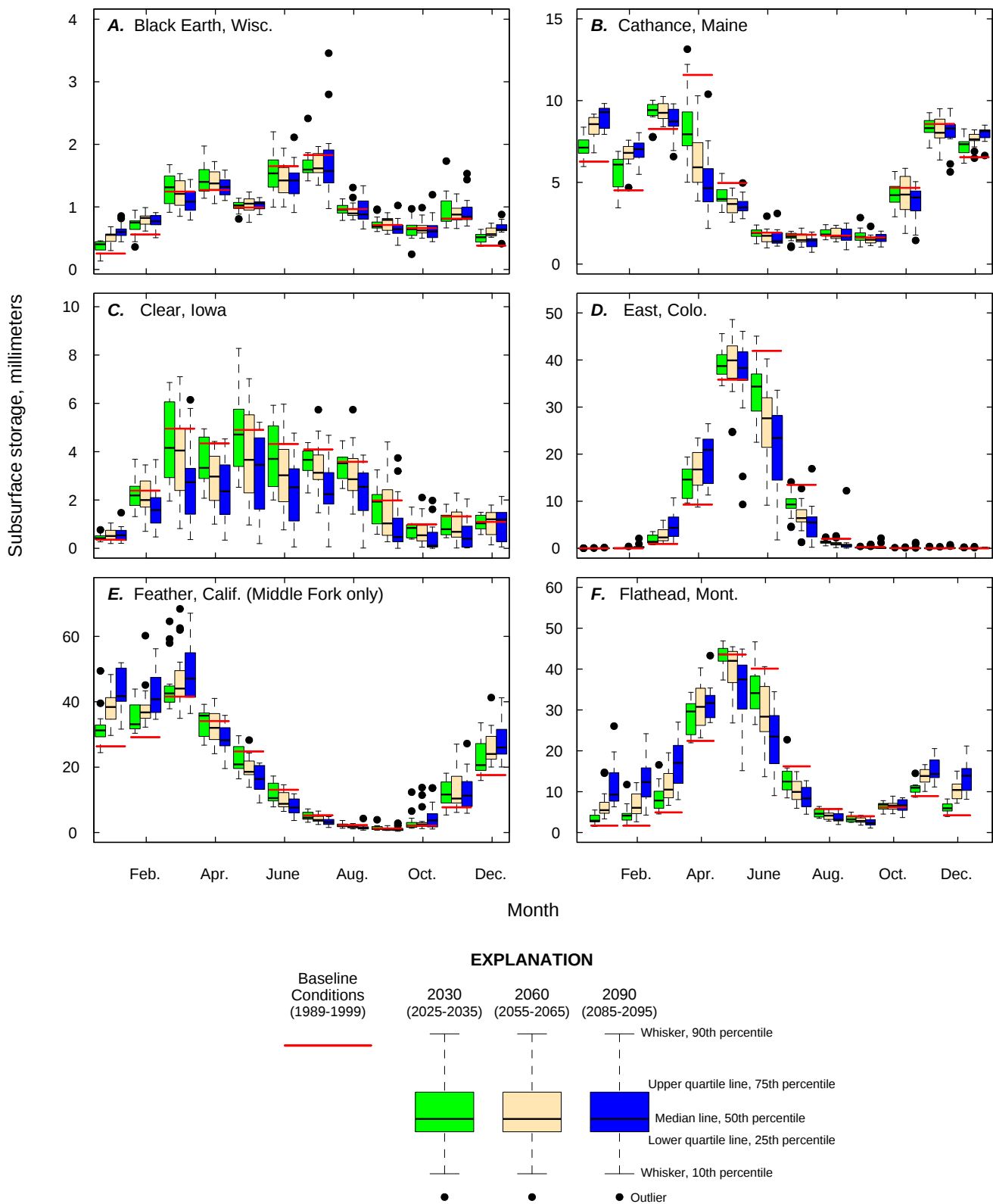


Figure 45. Mean daily subsurface reservoir storage values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

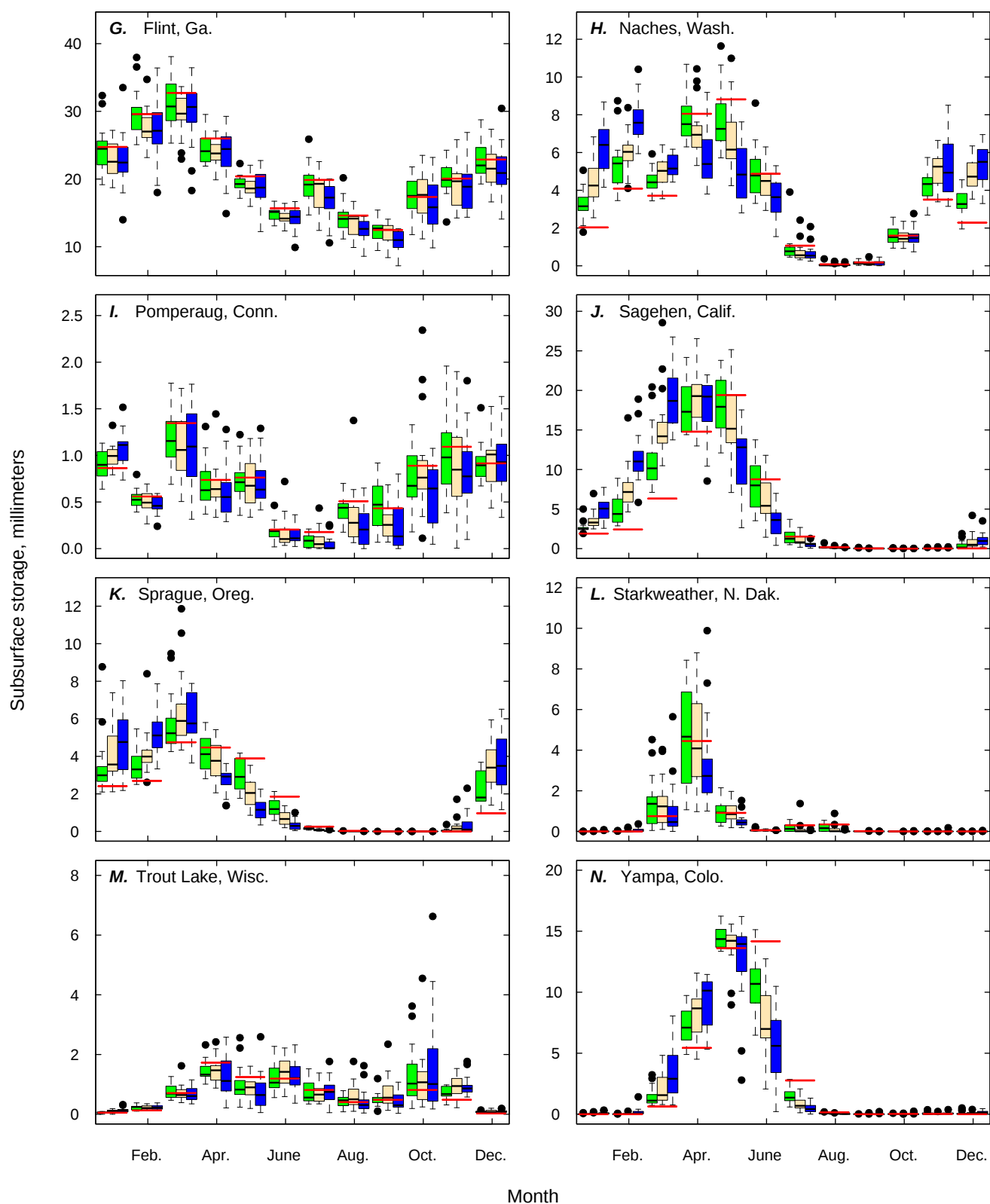


Figure 45. Mean daily subsurface reservoir storage values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

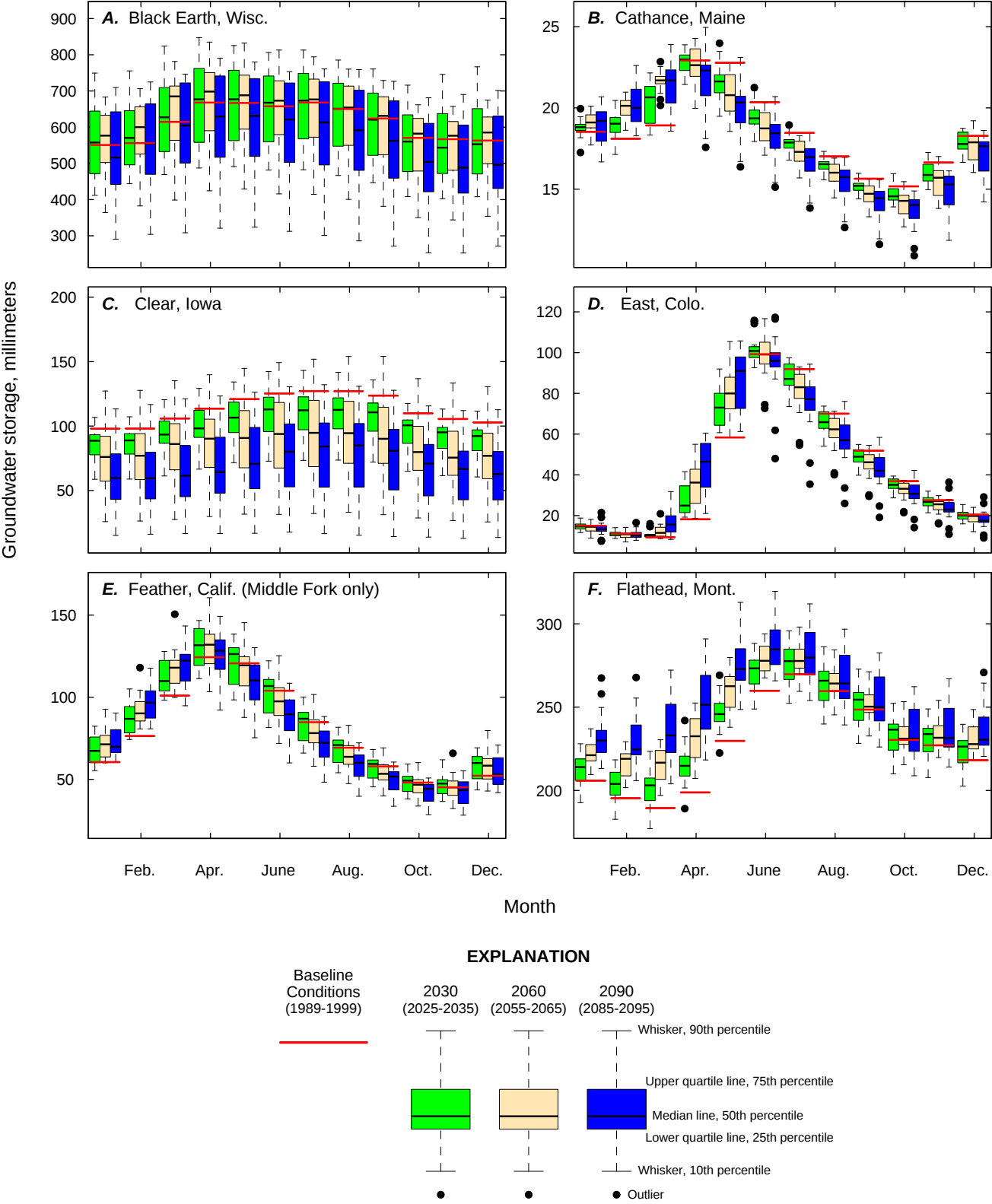


Figure 46. Mean daily groundwater reservoir storage values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.

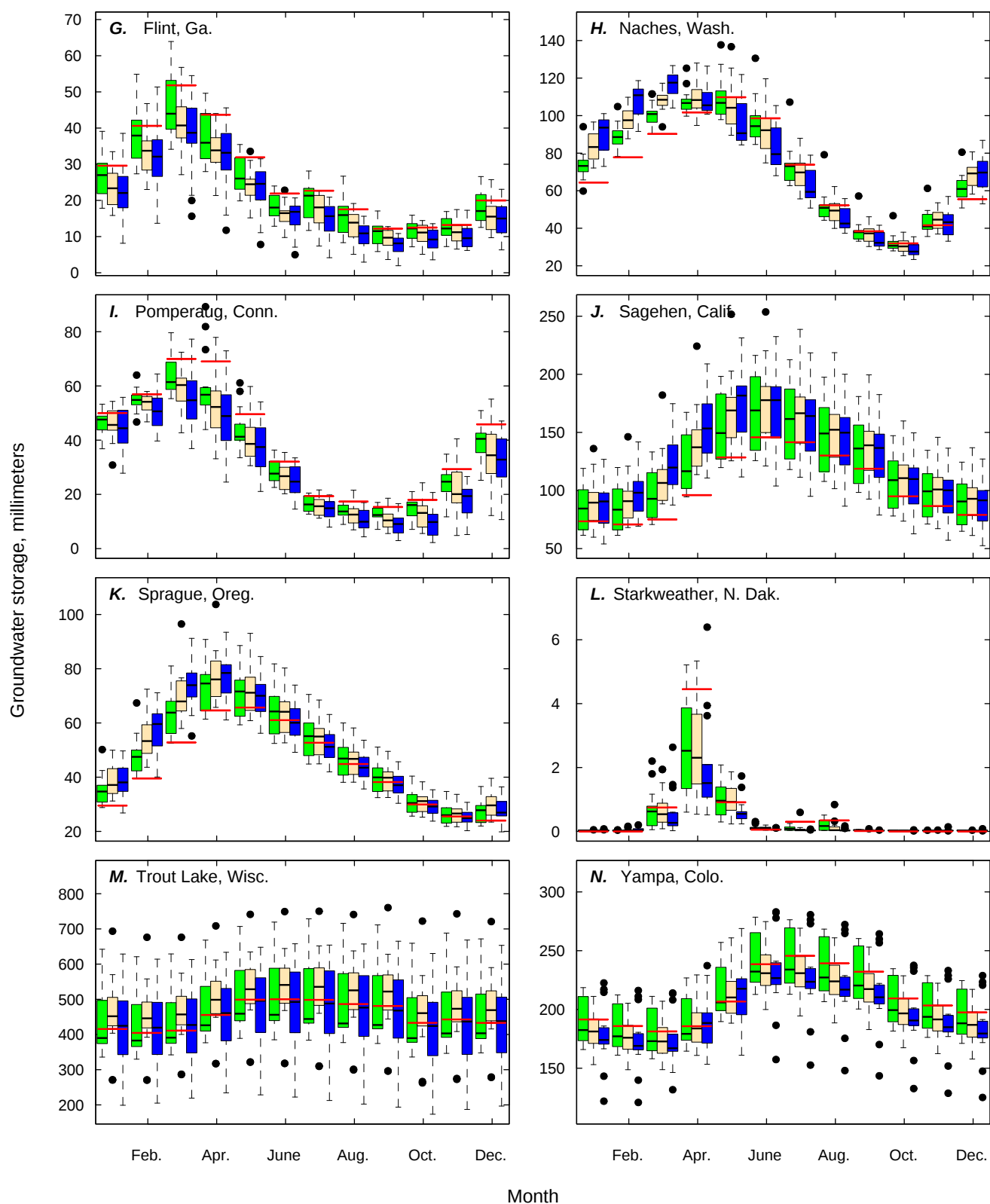


Figure 46. Mean daily groundwater reservoir storage values by month for baseline conditions and projected range (2030, 2060, and 2090) using the 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

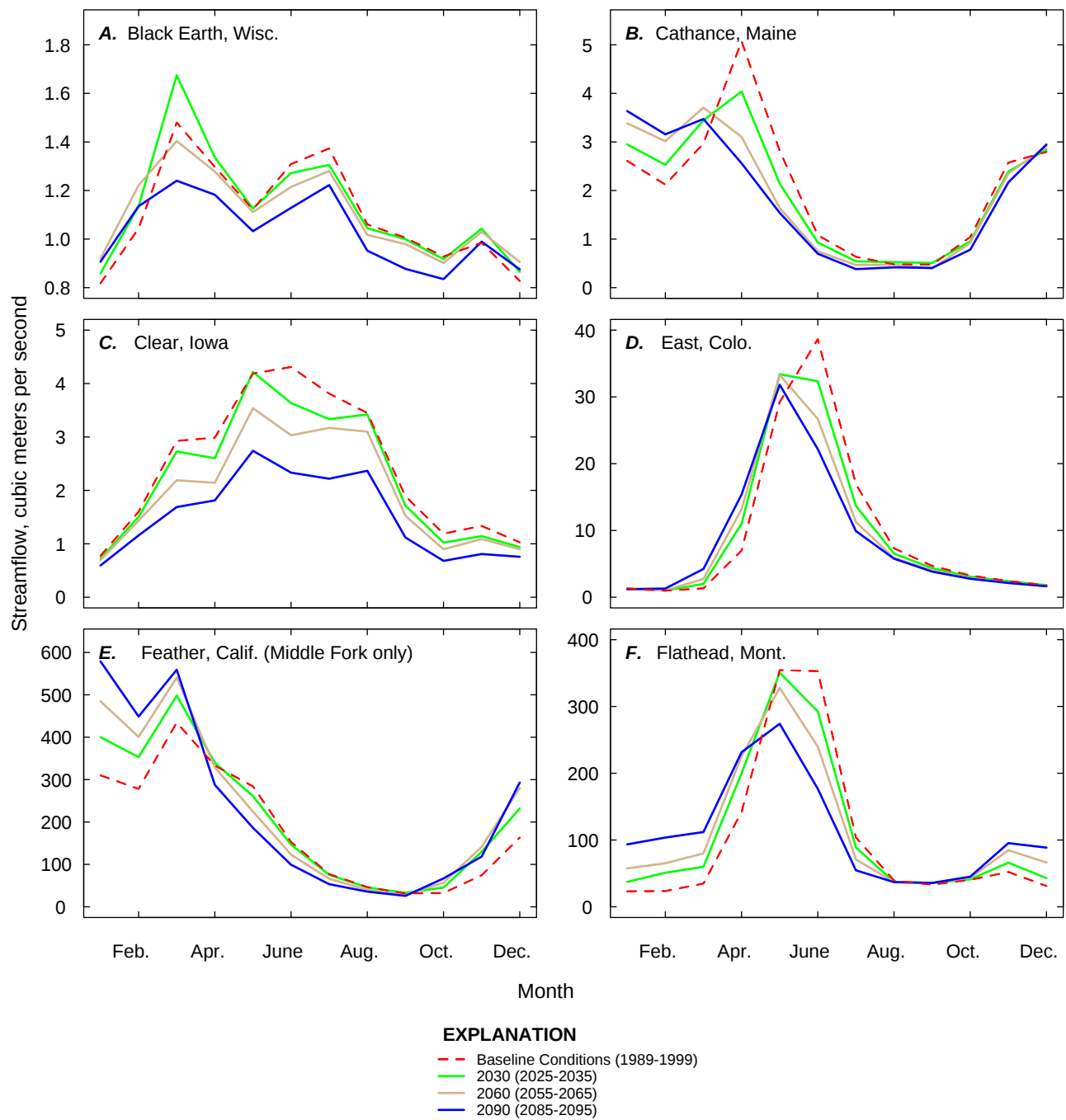


Figure 47. Mean monthly hydrographs for baseline conditions and projected conditions (2030, 2060, and 2090) using 5 general circulation models and 3 emission scenarios for the 14 basins.

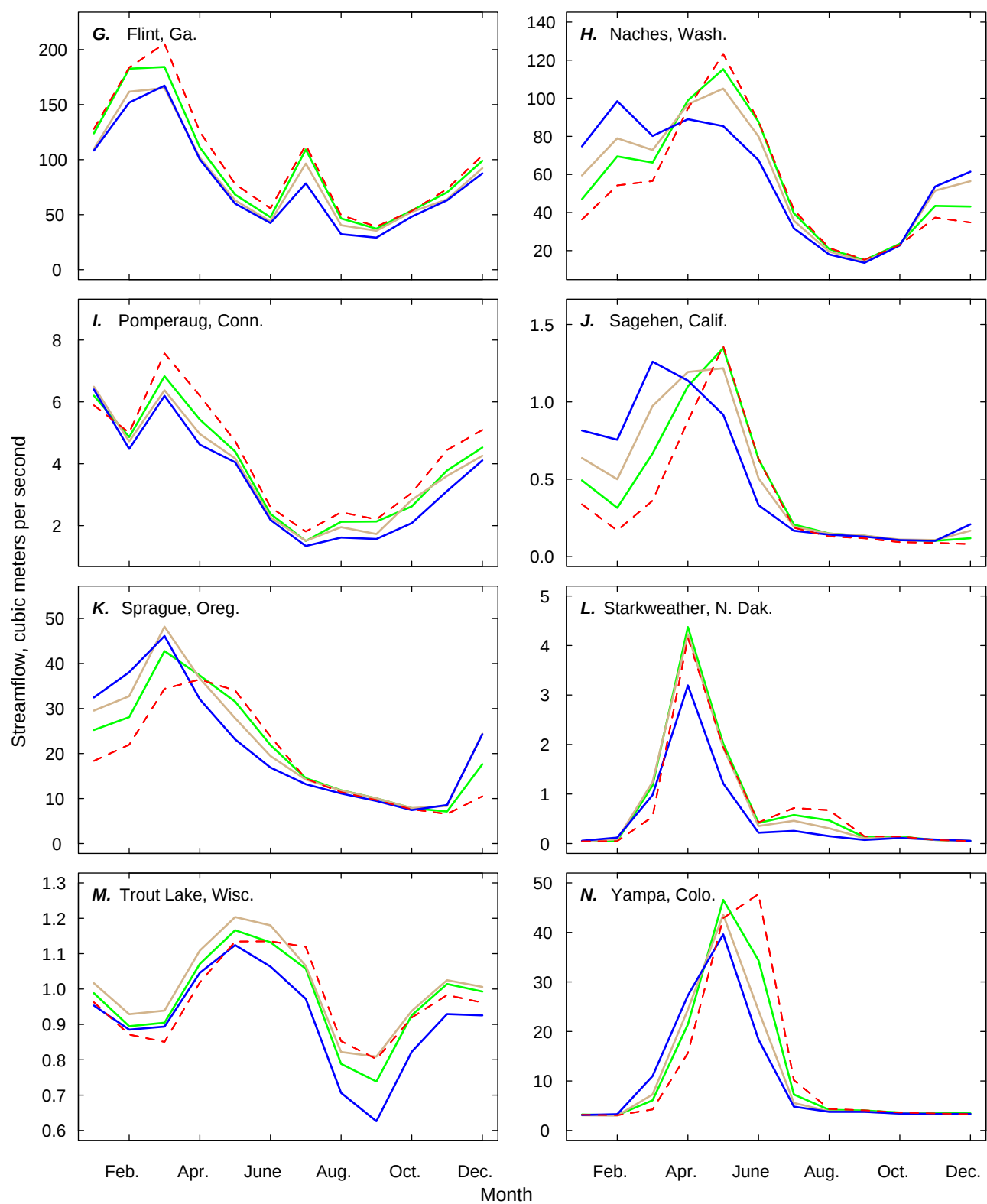


Figure 47. Mean monthly hydrographs for baseline conditions and projected conditions (2030, 2060, and 2090) using 5 general circulation models and 3 emission scenarios for the 14 basins.—Continued

Table 5. Projected change by year (slope) and adjusted R^2 (R_a^2) based on the central tendencies of the five General Circulation Models for the three carbon emission scenarios by basin for selected Precipitation Runoff Modeling System output variables.

[Blue shading indicates a significant negative trend and red shading indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation]

(A) Precipitation (millimeters): Area-weighted adjusted precipitation for basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Black Earth	0.0009	0.21	-0.0007	0.06	-0.0003	0.00
Cathance	0.0019	0.56	0.0019	0.58	0.0016	0.40
Clear	0.0008	0.14	-0.0018	0.25	0.0006	0.02
East	0.0013	0.50	-0.0007	0.11	0.0011	0.31
Feather	0.0021	0.38	0.0018	0.04	0.0003	-0.01
Flathead	0.0040	0.87	0.0049	0.69	0.0022	0.51
Flint	-0.0014	0.35	-0.0047	0.72	0.0008	0.04
Naches	0.0018	0.49	0.0029	0.77	0.0002	0.00
Pomperaug	0.0019	0.38	0.0002	-0.00	0.0024	0.57
Sagehen	0.0031	0.60	0.0033	0.16	0.0009	0.03
Sprague	0.0009	0.51	0.0015	0.43	0.0005	0.10
Starkweather	-0.0006	0.11	0.0004	0.08	0.0000	-0.01
Trout Lake	0.0004	0.02	0.0008	0.20	0.0007	0.05
Yampa	0.0011	0.56	-0.0001	-0.01	0.0004	0.10

(B) Maximum temperature (degrees celsius): Basin area-weighted daily maximum temperature						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Black Earth	0.040	1.00	0.049	0.99	0.024	0.97
Cathance	0.033	0.99	0.038	0.98	0.018	0.98
Clear	0.041	0.99	0.051	0.99	0.023	0.96
East	0.037	0.99	0.049	0.99	0.022	0.95
Feather	0.030	0.98	0.038	0.99	0.023	0.98
Flathead	0.031	0.98	0.040	0.99	0.024	0.98
Flint	0.034	0.99	0.043	0.98	0.018	0.95
Naches	0.028	0.97	0.033	0.97	0.023	0.99
Pomperaug	0.032	1.00	0.039	0.99	0.017	0.99
Sagehen	0.031	0.98	0.040	0.98	0.024	0.98
Sprague	0.030	0.97	0.037	0.98	0.024	0.99
Starkweather	0.040	0.99	0.046	0.98	0.026	0.96
Trout Lake	0.039	0.99	0.047	0.99	0.023	0.97
Yampa	0.035	0.99	0.047	0.99	0.023	0.97

(C) Minimum temperature (degrees celsius): Basin area-weighted daily minimum temperature						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Black Earth	0.04	1.00	0.048	0.99	0.024	0.97
Cathance	0.035	0.99	0.041	0.98	0.020	0.98
Clear	0.039	0.99	0.046	0.99	0.022	0.96
East	0.03	0.98	0.040	0.98	0.022	0.98

Table 5. Projected change by year (slope) and adjusted R^2 (R_a^2) based on the central tendencies of the five General Circulation Models for the three carbon emission scenarios by basin for selected Precipitation Runoff Modeling System output variables.—Continued

[Blue shading indicates a significant negative trend and red shading indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation]

(C) Minimum temperature (degrees celsius): Basin area-weighted daily minimum temperature—Continued						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Feather	0.030	0.99	0.036	0.98	0.022	0.99
Flathead	0.033	0.99	0.040	0.98	0.024	0.98
Flint	0.030	1.00	0.034	0.98	0.017	0.99
Naches	0.029	0.98	0.034	0.98	0.021	0.99
Pomperaug	0.034	1.00	0.040	0.99	0.019	0.99
Sagehen	0.031	0.99	0.037	0.98	0.023	0.99
Sprague	0.030	0.99	0.036	0.98	0.022	0.99
Starkweather	0.042	0.99	0.049	0.98	0.028	0.97
Trout Lake	0.041	1.00	0.048	0.99	0.026	0.97
Yampa	0.033	0.99	0.042	0.99	0.022	0.98

(D) Actual evapotranspiration (millimeters): Evapotranspiration on basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Black Earth	0.0027	0.92	0.0029	0.93	0.0018	0.69
Cathance	0.0049	0.99	0.0058	0.98	0.0029	0.98
Clear	0.0044	0.97	0.0042	0.94	0.0029	0.89
East	0.0031	0.98	0.0034	0.91	0.0025	0.92
Feather	0.0012	0.91	0.0013	0.76	0.0010	0.69
Flathead	0.0028	0.97	0.0035	0.94	0.0023	0.94
Flint	0.0013	0.75	0.0008	0.54	0.0012	0.70
Naches	0.0005	0.38	0.0010	0.77	0.0006	0.61
Pomperaug	0.0045	0.99	0.0053	0.99	0.0029	0.99
Sagehen	0.0020	0.94	0.0021	0.84	0.0013	0.82
Sprague	0.0007	0.58	0.0011	0.65	0.0007	0.49
Starkweather	0.0012	0.75	0.0018	0.88	0.001	0.66
Trout Lake	0.0018	0.93	0.0021	0.95	0.0013	0.82
Yampa	0.0024	0.97	0.0026	0.91	0.0017	0.96

(E) Growing season (days): Length of growing season						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Black Earth	0.3057	0.97	0.3536	0.98	0.1747	0.96
Cathance	0.3569	0.99	0.4008	0.97	0.2148	0.97
Clear	0.3095	0.99	0.3546	0.98	0.1597	0.95
East	0.5515	0.97	0.5849	0.99	0.3173	0.97
Feather	0.6285	0.98	0.7145	0.99	0.4385	0.99
Flathead	0.4506	0.98	0.5746	0.98	0.3720	0.97
Flint	0.3221	0.97	0.3207	0.95	0.1359	0.85
Naches	0.4343	0.99	0.5173	0.98	0.2974	0.99

Table 5. Projected change by year (slope) and adjusted R^2 (R^2_a) based on the central tendencies of the five General Circulation Models for the three carbon emission scenarios by basin for selected Precipitation Runoff Modeling System output variables.—Continued

[Blue shading indicates a significant negative trend and red shading indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation]

(E) Growing season (days): Length of growing season—Continued						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Pomperaug	0.4049	0.97	0.4248	0.99	0.2300	0.98
Sagehen	0.7947	0.95	0.8562	0.99	0.5801	0.98
Sprague	0.6852	0.96	0.7669	0.99	0.5400	0.98
Starkweather	0.2533	0.99	0.3093	0.98	0.1780	0.97
Trout Lake	0.3995	0.98	0.4699	0.98	0.2378	0.94
Yampa	0.4625	0.97	0.5116	1.00	0.2786	0.97
(F) Streamflow (cubic meters per second): Streamflow from basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	-0.0014	0.59	-0.0028	0.75	-0.0016	0.49
Cathance	-0.0029	0.78	-0.0038	0.84	-0.0012	0.35
Clear	-0.0101	0.73	-0.0169	0.86	-0.0065	0.48
East	-0.0150	0.70	-0.0347	0.91	-0.0118	0.51
Feather	0.1132	0.16	0.0820	0.00	-0.0378	-0.00
Flathead	0.0547	0.41	0.0676	0.22	-0.0010	-0.01
Flint	-0.2162	0.76	-0.4359	0.83	-0.0373	0.02
Naches	0.0561	0.59	0.0789	0.79	-0.0023	-0.01
Pomperaug	-0.0061	0.59	-0.0116	0.82	-0.0012	0.05
Sagehen	0.0004	0.23	0.0005	0.04	-0.0001	-0.01
Sprague	0.0107	0.15	0.0225	0.15	-0.0128	0.10
Starkweather	-0.0056	0.66	-0.0047	0.75	-0.0031	0.37
Trout Lake	-0.0015	0.61	-0.0014	0.74	-0.0004	0.04
Yampa	-0.0205	0.77	-0.0415	0.89	-0.0211	0.67
(G) Surface runoff (cubic meters per second): Surface runoff from basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	-0.00186	0.29	-0.00424	0.55	-0.00260	0.38
Cathance	0.00013	0.79	0.00014	0.78	0.00007	0.67
Clear	-0.00213	0.64	-0.00396	0.79	-0.00125	0.26
East	-0.00038	0.38	-0.00120	0.84	-0.00050	0.57
Feather	0.00450	0.15	0.00810	0.07	0.00074	-0.01
Flathead	0.02048	0.87	0.02764	0.76	0.01043	0.73
Flint	-0.05164	0.73	-0.11173	0.82	-0.00588	-0.00
Naches	0.01745	0.91	0.02563	0.93	0.01146	0.91
Pomperaug	0.00051	0.24	-0.00030	0.08	0.00048	0.30
Sagehen	0.00027	0.90	0.00044	0.66	0.00022	0.75
Sprague	0.00182	0.83	0.00277	0.67	0.00099	0.38
Starkweather	-0.00141	0.73	-0.00112	0.85	-0.00084	0.47

Table 5. Projected change by year (slope) and adjusted R^2 (R^2_a) based on the central tendencies of the five General Circulation Models for the three carbon emission scenarios by basin for selected Precipitation Runoff Modeling System output variables.—Continued[Blue shading indicates a significant negative trend and red shading indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation]

(G) Surface runoff (cubic meters per second): Surface runoff from basin—Continued						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Trout Lake	0.00004	0.37	0.00004	0.62	0.00003	0.15
Yampa	-0.00044	0.37	-0.00118	0.68	-0.00066	0.61
(H) Subsurface flow (cubic meters per second): Subsurface flow from basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	0.00022	0.15	-0.00027	0.13	-0.00024	0.09
Cathance	-0.00137	0.76	-0.00177	0.83	-0.00059	0.32
Clear	-0.00448	0.71	-0.00745	0.84	-0.00311	0.53
East	-0.01231	0.78	-0.02649	0.92	-0.01065	0.66
Feather	0.15873	0.37	0.14958	0.06	0.01897	-0.01
Flathead	-0.00625	0.00	-0.00338	-0.01	-0.03355	0.23
Flint	-0.04200	0.66	-0.09640	0.79	-0.00282	-0.01
Naches	0.02159	0.41	0.03062	0.63	-0.01020	0.22
Pomperaug	-0.00103	0.32	-0.00230	0.66	0.00026	0.02
Sagehen	0.00012	0.04	0.00007	-0.01	-0.00020	0.05
Sprague	0.00125	-0.00	0.00872	0.07	-0.01045	0.22
Starkweather	-0.00002	0.59	-0.00002	0.62	-0.00001	0.30
Trout Lake	-0.00000	0.54	-0.00000	0.67	-0.00000	0.46
Yampa	-0.01766	0.81	-0.03324	0.90	-0.01784	0.73
(I) Groundwater flow (cubic meters per second): Groundwater flow from basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	-0.00960	0.61	-0.01894	0.79	-0.01076	0.48
Cathance	-0.00004	0.90	-0.00005	0.82	-0.00001	0.37
Clear	-0.00349	0.80	-0.00550	0.91	-0.00217	0.57
East	-0.00232	0.37	-0.00698	0.81	-0.00063	0.02
Feather	-0.04999	0.59	-0.07566	0.43	-0.05747	0.52
Flathead	0.04074	0.88	0.04375	0.69	0.02210	0.56
Flint	-0.12264	0.81	-0.22784	0.86	-0.02873	0.06
Naches	0.02027	0.46	0.02671	0.67	-0.00580	0.08
Pomperaug	-0.00554	0.79	-0.00895	0.89	-0.00193	0.37
Sagehen	0.00002	0.01	-0.00003	-0.01	-0.00008	0.10
Sprague	0.00765	0.29	0.01102	0.20	-0.00335	0.03
Starkweather	-0.00417	0.61	-0.00357	0.68	-0.00222	0.33
Trout Lake	-0.00244	0.37	-0.00208	0.42	-0.00081	0.03
Yampa	-0.00241	0.54	-0.00709	0.81	-0.00260	0.31

Table 5. Projected change by year (slope) and adjusted R^2 (R^2_a) based on the central tendencies of the five General Circulation Models for the three carbon emission scenarios by basin for selected Precipitation Runoff Modeling System output variables.—Continued

[Blue shading indicates a significant negative trend and red shading indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation]

(J) Percent snow: Percent precipitation that falls as snow for total basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	-0.09	0.92	-0.10	0.95	-0.06	0.70
Cathance	-0.12	0.98	-0.13	0.99	-0.07	0.95
Clear	-0.06	0.89	-0.06	0.93	-0.04	0.69
East	-0.09	0.95	-0.10	0.95	-0.07	0.87
Feather	-0.15	0.96	-0.19	0.99	-0.14	0.96
Flathead	-0.18	0.97	-0.21	0.95	-0.13	0.94
Flint	-0.01	0.90	-0.00	0.92	-0.00	0.94
Naches	-0.17	0.96	-0.22	0.97	-0.15	0.97
Pomperaug	-0.10	0.98	-0.12	0.98	-0.07	0.97
Sagehen	-0.17	0.96	-0.23	0.97	-0.14	0.96
Sprague	-0.19	0.98	-0.25	0.98	-0.16	0.97
Starkweather	-0.02	0.11	-0.03	0.45	-0.01	0.09
Trout Lake	-0.06	0.92	-0.06	0.92	-0.04	0.72
Yampa	-0.08	0.96	-0.10	0.96	-0.06	0.93

(K) Snow-covered area (percent): Average snow-covered area for total basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	-0.16	0.97	-0.18	0.98	-0.11	0.89
Cathance	-0.13	0.99	-0.16	0.98	-0.08	0.97
Clear	-0.11	0.97	-0.12	0.98	-0.07	0.88
East	-0.10	0.98	-0.15	0.98	-0.07	0.94
Feather	-0.13	0.97	-0.17	0.97	-0.12	0.95
Flathead	-0.14	0.98	-0.18	0.96	-0.11	0.96
Flint	-0.00	0.89	-0.00	0.93	-0.00	0.93
Naches	-0.16	0.97	-0.20	0.97	-0.14	0.98
Pomperaug	-0.10	0.98	-0.12	0.98	-0.06	0.97
Sagehen	-0.13	0.97	-0.17	0.94	-0.11	0.92
Sprague	-0.13	0.98	-0.17	0.98	-0.12	0.98
Starkweather	-0.08	0.96	-0.09	0.96	-0.05	0.88
Trout Lake	-0.12	0.98	-0.13	0.98	-0.08	0.90
Yampa	-0.10	0.98	-0.16	0.98	-0.08	0.94

(L) Snowpack water equivalent (millimeters): Average snowpack water equivalent for total basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	-0.19	0.86	-0.23	0.90	-0.16	0.73
Cathance	-0.28	0.95	-0.29	0.98	-0.17	0.94
Clear	-0.06	0.82	-0.07	0.90	-0.04	0.70
East	-0.47	0.93	-0.77	0.96	-0.38	0.82

Table 5. Projected change by year (slope) and adjusted R^2 (R^2_a) based on the central tendencies of the five General Circulation Models for the three carbon emission scenarios by basin for selected Precipitation Runoff Modeling System output variables.—Continued[Blue shading indicates a significant negative trend and red shading indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation]

(L) Snowpack water equivalent (millimeters): Average snowpack water equivalent for total basin—Continued						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Feather	-0.54	0.92	-0.6	0.88	-0.53	0.87
Flathead	-1.21	0.95	-1.45	0.95	-1.04	0.92
Flint	-0.00	0.90	-0.00	0.92	-0.00	0.89
Naches	-0.46	0.91	-0.58	0.95	-0.51	0.97
Pomperaug	-0.06	0.91	-0.06	0.97	-0.03	0.88
Sagehen	-0.79	0.88	-1.04	0.77	-0.79	0.80
Sprague	-0.16	0.95	-0.19	0.96	-0.15	0.96
Starkweather	-0.10	0.84	-0.11	0.82	-0.07	0.67
Trout Lake	-0.17	0.91	-0.18	0.92	-0.12	0.80
Yampa	-0.34	0.93	-0.54	0.96	-0.32	0.87

(M) Snowmelt (millimeters): Average snowmelt for total basin						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	-0.0030	0.89	-0.0035	0.91	-0.0021	0.70
Cathance	-0.0039	0.95	-0.0047	0.97	-0.0021	0.93
Clear	-0.0017	0.86	-0.0021	0.93	-0.0012	0.74
East	-0.0014	0.69	-0.0031	0.86	-0.0012	0.52
Feather	-0.0052	0.84	-0.0074	0.71	-0.0052	0.72
Flathead	-0.0008	0.30	-0.0010	0.17	-0.0010	0.18
Flint	-0.0002	0.90	-0.0002	0.96	-0.0001	0.94
Naches	-0.0014	0.57	-0.0020	0.70	-0.0022	0.87
Pomperaug	-0.0040	0.96	-0.0053	0.98	-0.0025	0.94
Sagehen	0.0004	0.01	-0.0005	-0.01	-0.0011	0.07
Sprague	-0.0025	0.97	-0.0032	0.94	-0.0022	0.89
Starkweather	-0.0004	0.29	-0.0003	0.30	-0.0002	0.11
Trout Lake	-0.0012	0.80	-0.0014	0.77	-0.0009	0.61
Yampa	-0.0010	0.64	-0.0023	0.82	-0.0011	0.59

(M) Infiltration (millimeters)						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	0.0011	0.32	-0.0002	-0.00	0.0001	-0.01
Cathance	0.0016	0.63	0.0017	0.65	0.0013	0.43
Clear	0.0015	0.48	-0.0006	0.04	0.0010	0.13
East	0.0009	0.34	-0.0010	0.25	0.0007	0.18
Feather	0.0022	0.43	0.0019	0.06	0.0005	0.00
Flathead	0.0030	0.82	0.0037	0.62	0.0016	0.40
Flint	-0.0007	0.19	-0.0031	0.67	0.0008	0.09
Naches	0.0026	0.74	0.0037	0.87	0.0010	0.39

Table 5. Projected change by year (slope) and adjusted R^2 (R_a^2) based on the central tendencies of the five General Circulation Models for the three carbon emission scenarios by basin for selected Precipitation Runoff Modeling System output variables.—Continued

[Blue shading indicates a significant negative trend and red shading indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation]

(M) Infiltration (millimeters)—Continued						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Pomperaug	0.0016	0.40	0.0004	0.02	0.0022	0.59
Sagehen	0.0023	0.51	0.0021	0.09	0.0003	-0.00
Sprague	0.0008	0.53	0.0014	0.45	0.0005	0.14
Starkweather	0.0006	0.30	0.0014	0.73	0.0007	0.31
Trout Lake	0.0002	0.00	0.0006	0.14	0.0006	0.06
Yampa	0.0005	0.23	-0.0007	0.20	-0.0000	-0.01

(O) Subsurface reservoir inflow (millimeters)						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Black Earth	0.0000	0.09	-0.0001	0.20	-0.0001	0.17
Cathance	-0.0022	0.75	-0.0028	0.83	-0.0009	0.31
Clear	-0.0016	0.71	-0.0027	0.85	-0.0011	0.53
East	-0.0015	0.78	-0.0032	0.92	-0.0013	0.66
Feather	0.0016	0.34	0.0015	0.05	0.0001	-0.01
Flathead	0.0008	0.26	0.0009	0.12	-0.0002	-0.00
Flint	-0.0011	0.68	-0.0025	0.80	-0.0001	-0.01
Naches	0.0010	0.49	0.0014	0.69	-0.0003	0.09
Pomperaug	-0.0006	0.30	-0.0014	0.65	0.0002	0.02
Sagehen	0.0006	0.09	0.0006	0.00	-0.0005	0.02
Sprague	0.0000	-0.00	0.0002	0.07	-0.0002	0.22
Starkweather	-0.0005	0.60	-0.0004	0.66	-0.0003	0.32
Trout Lake	-0.0011	0.40	-0.0010	0.47	-0.0004	0.05
Yampa	-0.0011	0.81	-0.0021	0.90	-0.0011	0.73

(P) Groundwater reservoir recharge						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Black Earth	-0.0033	0.62	-0.0066	0.80	-0.0037	0.49
Cathance	-0.0001	0.88	-0.0001	0.79	-0.0000	0.30
Clear	-0.0013	0.79	-0.0021	0.91	-0.0008	0.58
East	-0.0003	0.36	-0.0008	0.81	-0.0001	0.02
Feather	-0.0007	0.53	-0.0010	0.34	-0.0008	0.46
Flathead	0.0009	0.87	0.0010	0.67	0.0005	0.52
Flint	-0.0014	0.81	-0.0026	0.86	-0.0003	0.06
Naches	0.0006	0.42	0.0008	0.64	-0.0002	0.12
Pomperaug	-0.0025	0.79	-0.0040	0.89	-0.0008	0.36
Sagehen	0.0001	0.01	-0.0001	-0.01	-0.0003	0.10
Sprague	0.0002	0.34	0.0003	0.24	-0.0001	0.02
Starkweather	-0.0005	0.61	-0.0005	0.68	-0.0003	0.33

Table 5. Projected change by year (slope) and adjusted R^2 (R^2_a) based on the central tendencies of the five General Circulation Models for the three carbon emission scenarios by basin for selected Precipitation Runoff Modeling System output variables.—Continued[Blue shading indicates a significant negative trend and red shading indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation]

(P) Groundwater reservoir recharge—Continued						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Trout Lake	-0.0028	0.38	-0.0025	0.44	-0.0011	0.04
Yampa	-0.0002	0.66	-0.0006	0.83	-0.0002	0.38
(Q) Soil moisture						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	-0.0549	0.93	-0.0785	0.95	-0.0326	0.66
Cathance	-0.0405	0.98	-0.0456	0.93	-0.0204	0.96
Clear	-0.3386	0.93	-0.5071	0.95	-0.1537	0.55
East	-0.0725	0.72	-0.1258	0.82	-0.0113	0.03
Feather	-0.0710	0.89	-0.0908	0.83	-0.0564	0.69
Flathead	-0.0456	0.69	-0.0760	0.86	-0.0322	0.38
Flint	-0.2076	0.95	-0.2995	0.97	-0.0902	0.56
Naches	-0.0344	0.69	-0.0325	0.81	-0.0343	0.86
Pomperaug	-0.1650	0.95	-0.2205	0.93	-0.0534	0.64
Sagehen	-0.1596	0.68	-0.2340	0.51	-0.1792	0.53
Sprague	-0.0315	0.56	-0.0333	0.48	-0.0289	0.37
Starkweather	-0.3260	0.82	-0.3400	0.93	-0.2087	0.63
Trout Lake	-0.0217	0.95	-0.0227	0.97	-0.0106	0.67
Yampa	-0.0852	0.84	-0.1596	0.90	-0.0476	0.63
(R) Subsurface reservoir storage						
Basin	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R^2_a	slope	R^2_a	slope	R^2_a
Black Earth	0.0005	0.26	-0.0002	0.06	-0.0002	0.04
Cathance	-0.0040	0.79	-0.0056	0.76	-0.0011	0.13
Clear	-0.0120	0.74	-0.0204	0.88	-0.0077	0.53
East	-0.0172	0.77	-0.0367	0.93	-0.0143	0.65
Feather	0.0059	0.09	0.0018	-0.01	-0.0027	-0.00
Flathead	0.0080	0.87	0.0087	0.66	0.0044	0.52
Flint	-0.0223	0.67	-0.0508	0.83	0.0003	-0.01
Naches	0.0050	0.68	0.0068	0.82	0.0008	0.06
Pomperaug	-0.0008	0.32	-0.0019	0.66	0.0002	0.02
Sagehen	0.0023	0.07	0.0010	-0.01	-0.0029	0.04
Sprague	0.0002	-0.01	0.0018	0.06	-0.0024	0.24
Starkweather	-0.0044	0.60	-0.0038	0.66	-0.0024	0.32
Trout Lake	-0.0001	-0.01	-0.0005	0.07	-0.0004	0.02
Yampa	-0.0054	0.81	-0.0099	0.90	-0.0052	0.71

Table 5. Projected change by year (slope) and adjusted R^2 (R_a^2) based on the central tendencies of the five General Circulation Models for the three carbon emission scenarios by basin for selected Precipitation Runoff Modeling System output variables.—Continued

[Blue shading indicates a significant negative trend and red shading indicates a significant positive trend ($p < 0.05$) accounting for lag-1 autocorrelation]

Basin	(S) Groundwater reservoir storage					
	Emission scenario A1B		Emission scenario A2		Emission scenario B1	
	slope	R_a^2	slope	R_a^2	slope	R_a^2
Black Earth	-0.8165	0.61	-1.6218	0.79	-0.9220	0.47
Cathance	-0.0149	0.81	-0.0214	0.79	-0.0048	0.23
Clear	-0.4921	0.80	-0.7756	0.91	-0.3054	0.57
East	-0.0265	0.36	-0.0798	0.81	-0.0072	0.02
Feather	-0.1055	0.68	-0.1568	0.53	-0.1169	0.61
Flathead	0.7227	0.88	0.7760	0.69	0.3923	0.56
Flint	-0.0781	0.83	-0.1400	0.86	-0.0197	0.08
Naches	0.0385	0.37	0.0519	0.60	-0.0203	0.19
Pomperaug	-0.0835	0.78	-0.1365	0.89	-0.0288	0.34
Sagehen	0.0250	0.01	-0.0309	-0.01	-0.0875	0.10
Sprague	0.0304	0.29	0.0438	0.20	-0.0133	0.03
Starkweather	-0.0028	0.61	-0.0024	0.68	-0.0015	0.33
Trout Lake	-0.6985	0.38	-0.6133	0.44	-0.2613	0.04
Yampa	-0.1453	0.54	-0.4269	0.81	-0.1565	0.31

The PRMS projections indicate minimal changes in annual mean surface runoff and subsurface flow (figs. 16*A* and 17*A*) with no significant trends (table 5*G* and 5*H*). High variability (indicating uncertainty) in the seasonal timing of these variables is projected (figs. 34*A* and 35*A*), with decreasing flows in March, June, and July, and increasing flows in December through February. Groundwater flow is the dominant component of streamflow and has a similar seasonal pattern (figs. 15*A* and 18*A*, table 5*F* and 5*J*). High variability in the seasonal timing of groundwater flow also is projected (fig. 36*A*).

Black Earth Creek Snow Variables

Because of increases in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19*A*), percentage of area covered by snow (fig. 20*A*), snowpack water equivalent (fig. 21*A*), and snowmelt (fig. 22*A*) will decrease (tables 5*J*–5*M*). Projected warmer temperatures will: (1) result in less snowfall, which will result in less snow available to melt in the spring; (2) decrease the amount of precipitation that falls as snow, percentage of area covered by snow, and snowpack water equivalent in every month (figs. 37*A*–39*A*); and (3) increase the snowmelt in December and January, leaving less snow to melt in March and April (fig. 40*A*).

Black Earth Creek Soil Moisture Variables

The PRMS projections of infiltration (figs. 23*A* and 41*A*, respectively) are similar to those of precipitation (figs. 10*A* and 29*A*). The large range in the infiltration projections indicates a high level of uncertainty with no significant overall trend (table 5*N*). High variability in seasonal infiltration is also projected (fig. 41*A*), with large increases projected for the months of November through February and large decreases projected for the months of March and June.

Projected changes in the inflow to the subsurface reservoir (fig. 24*A*) and recharge to the groundwater reservoir (fig. 25*A*) are similar to the changes seen in the subsurface (fig. 17*A*) and groundwater (fig. 18*A*) flow components, respectively. The projections show no significant trends in inflow to the subsurface reservoir (table 5*O*). The large range in the projections of annual mean recharge indicates a high level of uncertainty, with decreases projected for the A1B and A2 scenarios (table 5*P*). High variability in seasonal recharge is projected (figs. 42*A* and 43*A*).

There is a decrease in the projected annual mean soil moisture (fig. 26*A* and table 5*Q*). High variability in seasonal timing of soil moisture is projected (fig. 44*A*), with decreases in nearly every month. The largest decreases occur during the growing season months.

The PRMS projections indicate that the storage in the subsurface (fig. 27*A*) and the groundwater reservoirs (fig. 28*A*) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5*O* and 5*P*; 5*R* and 5*S*, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45*A* and 46*A*).

Discussion of Black Earth Creek Projections

In the Black Earth Creek Basin, the PRMS projections indicate an overall decrease in snowmelt and increase in evapotranspiration in the basin, although the large range of values indicate a large degree of uncertainty in these trends. Annual mean streamflow may not change substantially; however, the seasonal timing of flows, such as reduced spring runoff, and local conditions, such as soil moisture, may be affected. This could affect in-stream water quality and temperature. This study did not consider future changes in land use or the demand for water from irrigation or urbanization in the basin. The interaction of land use and climate change may affect the watershed hydrology in ways that were not accounted for by this study.

Precipitation Runoff Modeling System Projections for Cathance Stream, Maine

The Cathance Stream basin is a tributary to the Dennys River Basin in Washington County, eastern Maine, on the coast of the Atlantic Ocean. At USGS streamflow-gaging station 01021230 the basin has a drainage area of 85 km² (table 1). The initial PRMS application was developed by Dudley (2008). The PRMS application contains 85 HRUs with elevations that range from 46 to 174 m (table 1).

Mean monthly baseline conditions in the Cathance Stream Basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from just above -13 °C during January to just below 13 °C during July and August. Mean monthly maximum temperatures range from -1 °C during January to above 25 °C during July and August. On a mean monthly basis, precipitation ranges from 2.5 mm/day (during August) to 4.7 mm/day (during January). Because of freezing temperatures during the winter months (December, January, and February), precipitation falls as snow and accumulates throughout the winter, with peak mean monthly streamflow occurring during April because of snowmelt.

The Cathance Stream Basin is rural and its rolling topography is predominantly forested with wetlands, lakes, and ponds, some blueberry agriculture fields, clear cuts, partial cuts, regenerating forest, and light residential development. The Dennys River and its tributary, Cathance Stream, are important habitat for wild Atlantic salmon. Wild Atlantic salmon populations are protected under the U.S. Endangered Species Act of 1973 and are the subject of a comprehensive recovery program (Dudley, 2008). The USGS, in cooperation with the Maine Department of Marine Resources Bureau

of Sea Run Fisheries and Habitat, began a study in 2004 to characterize the quantity, variability, and timing of streamflow in the Dennys River (Dudley, 2008). The study included the development and evaluation of a distributed-parameter watershed model (PRMS). The watershed modeling work supports several tasks directed at developing water-use management plans for Atlantic salmon rivers and development of comprehensive flow monitoring in those basins. Improved water-resources information supports Atlantic salmon protection efforts.

Cathance Stream Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Cathance Stream Basin (fig. 10*B*), with similar increases in the central tendencies for the A1B and A2 scenarios (table 5*A*). Seasonal variability is apparent also in these precipitation projections (fig. 29*B*), with the largest projected increases in January, July, and August.

All PRMS projections indicate steady increases in temperature (figs. 11*B*, 12*B*, 30*B*, and 31*B* and tables 5*B* and 5*C*), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5*B* and 5*C*).

All projections indicate steady increases in annual mean evapotranspiration (fig. 13*B* and table 5*D*). High variability in seasonal evapotranspiration is projected, with increases in all months (fig. 32*B*).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14*B* and table 5*E*). This increase is because of the projected increase of annual temperature (figs. 11*B* and 12*B* and tables 5*B* and 5*C*). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5*E*).

Cathance Stream Streamflow Variables

The PRMS projections indicate a decrease in annual mean streamflow for the A1B and A2 scenarios (fig. 15*B* and table 5*F*), though the uncertainties associated with these projections are large. Large increases in seasonal streamflow are projected from January through March, whereas large decreases are projected from April through June. The PRMS simulation of baseline conditions (red line, fig. 33*B*) indicate that peak streamflow occurs in April, whereas the PRMS projections indicate that, by the end of the 21st century, the peak of streamflow could occur as early as January (figs. 33*B*). High variability (indicating uncertainty) is projected in the seasonal timing of streamflow in all but the summer months.

The PRMS projections indicate increases in annual mean surface runoff for all scenarios (fig. 16*B* and table 5*G*). Seasonal projections indicate that the increases in surface runoff are because of increased surface runoff in December, January, February, and April (fig. 34*B*). Projections of subsurface and groundwater flow indicate the same decreasing pattern as the

annual mean streamflow (figs. 15*B*, 17*B* and 18*B*, and tables 5*F*, 5*H* and 5*I*). Projected decreases in April subsurface flow (fig. 35*B*) leads to an overall decrease in the annual mean subsurface flow (fig. 25*B*). Monthly decreases in groundwater are projected in May through December, whereas increases are projected from January through March (fig. 36*B*).

Cathance Stream Snow Variables

Because of increases in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19*B*), percentage of area covered by snow (fig. 20*B*), snowpack water equivalent (fig. 21*B*), and snowmelt (fig. 22*B*) will decrease (tables 5*J*–5*M*). Projected warmer temperatures will: (1) result in less snowfall, which will result in less snow available to melt in the spring; (2) decrease the amount of precipitation that falls as snow, percentage of area covered by snow, and snowpack water equivalent in every month (figs. 37*B*–39*B*); and (3) increase the snowmelt in January and February, leaving less snow to melt in April (fig. 40*B*).

Cathance Stream Soil Moisture Variables

The PRMS projections of infiltration (figs. 23*B* and 41*B*) are similar to those of precipitation (figs. 10*B* and 29*B*). The large range in the infiltration projections indicates a high level of uncertainty with an increasing trend for the A1B and A2 scenarios (table 5*N*). Seasonal infiltration is projected to decrease during April (fig. 41*B*) because of the decrease in snowmelt in that month (fig. 40*B*). The overall increase in infiltration is because of the increase in precipitation.

Projected changes in the inflow to the subsurface reservoir (fig. 24*B*) and recharge to the groundwater reservoir (fig. 25*B*) are similar to the changes seen in the subsurface (fig. 17*B*) and groundwater (fig. 18*B*) flow components, respectively. The projections show a negative trend in the inflows to the subsurface reservoirs and recharge for the A1B and A2 scenarios (figs. 24*B* and 25*B* and tables 5*O* and 5*P*). High variability in seasonal recharge is projected (figs. 42*B* and 43*B*).

There is a decrease in the projected annual mean soil moisture (fig. 26*B* and table 5*Q*). High variability in seasonal timing of soil moisture is projected (fig. 44*B*), with decreases in nearly every month. The largest decreases occur during the growing season months.

The PRMS projections indicate that the storage in the subsurface (fig. 27*B*) and the groundwater reservoirs (fig. 28*B*) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5*O* and 5*P*; 5*R* and 5*S*, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45*B* and 46*B*).

Discussion of Cathance Stream Projections

The Cathance Stream and other coastal river basins in Maine are important habitat for a variety of aquatic animals including endangered Atlantic salmon. The broader-scale consequences of climate change on the flow regime of the rivers in Maine, like Cathance Stream, likely will affect ecosystems. The ecological implications of changes in the timing of winter-spring streamflows in Maine are not well understood and possibly could affect Atlantic salmon survival rates. If the peak spring migration of juvenile salmon from freshwater rivers (which is controlled by photoperiod, temperature, and flow) becomes out of phase with optimal environmental conditions in rivers, estuaries, or the ocean, salmon survival could drop substantially (McCormick and others, 1998). Changes in the timing of spring river flows may affect fresh-water mixing in estuarine ecosystems. Earlier snowmelt and increases in summer evapotranspiration rates could shorten the annual periods of standing water used by amphibians for breeding in forested depressional wetlands (vernal pools).

Precipitation Runoff Modeling System Projections for Clear Creek, Iowa

The Clear Creek Basin is located in east-central Iowa, includes parts of Iowa and Johnson Counties, and is a tributary to the Iowa River. At USGS streamflow-gaging station 05454300, the basin has a drainage area of 254 km². The PRMS application contains 75 HRUs with elevations that range from 205 to 261 meters (table 1).

Mean monthly baseline conditions for the Clear Creek basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from –13.5 °C during January to above 14 °C during July. Mean monthly maximum temperatures range from approximately –6 °C during January to above 24 °C during July. On a mean monthly basis, precipitation ranges from approximately 1.0 mm/day (from December through February) to almost 6.0 mm/day (during June). Because of freezing temperatures during the winter months, precipitation falls as snow and accumulates throughout the winter. The majority of annual mean streamflow occurs during spring and summer, with the peak monthly streamflow occurring during May and June.

The topography of the Clear Creek Basin is characterized by uplands dissected by tributary streams (Schwob, 1964). The Clear Creek Basin has a wide and broad flood plain, with a meandering channel except along reaches that have been straightened (Barnes and Eash, 1990). Land use is predominantly agricultural with a growing urban population in Johnson County, Iowa. The Iowa Department of Natural Resources and the Natural Resources Conservation Service (NRCS) have created the Clear Creek Watershed Enhancement Project (Iowa Department of Natural Resources, 2011). The goal of this project is to restore Clear Creek to natural conditions.

Clear Creek Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Clear Creek basin (fig. 10C). The large range in the precipitation projections indicates a large amount of uncertainty with no apparent overall trend (table 5A). Seasonal variability also is apparent in these precipitation projections (fig. 29C), with the lowest precipitation amounts (September through April) having minimal changes and low variability. Months with the highest projected precipitation amounts (May through August) have both increasing and decreasing precipitation, but with much larger uncertainty.

All PRMS projections indicate steady increases in temperature (figs. 11C, 12C, 30C, and 31C, and tables 5B and 5C), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5B and 5C).

All projections indicate steady increases in annual mean evapotranspiration (figs. 13C and table 5D). High variability in seasonal evapotranspiration is projected, with increases in all months, except July (fig. 32C).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14C and table 5E). This increase is because of the projected increase of annual temperature (figs. 11C and 12C and tables 5B and 5C). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5E).

Clear Creek Streamflow Variables

The PRMS projections indicate a decrease in annual mean streamflow for the A1B and A2 scenarios (fig. 15C and table 5F), though the uncertainties associated with these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, particularly from March through August. Specifically, the PRMS simulation of baseline conditions (red line, fig. 33C) indicate that peak streamflow occurs in June. Whereas the PRMS projections indicate that, by the end of the 21st century, the projected month of peak streamflow will be May (fig. 33C).

The PRMS projections indicate decreases in all three components of flow (figs. 16C–18C and tables 5G–5I). High variability (indicating uncertainty) in the seasonal timing of these variables is projected (34C–36C), with the largest decreases in surface runoff in May and June, the largest decreases in subsurface flows in May through August, and groundwater flow shows relatively consistent decreases in every month.

Clear Creek Snow Variables

Because of increases in temperature, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19C), percentage of area covered by snow (fig. 20C), snowpack water equivalent (fig. 21C), and snowmelt (fig. 22C) will decrease (tables 5J–5M). Projected

warmer temperatures will: (1) result in less snowfall, which will result in less snow available to melt in the spring; (2) decrease the amount of precipitation that falls as snow, percentage of area covered by snow, and snowpack water equivalent in every month (figs. 37C–39C); and (3) increase the snowmelt in December and January, with large decreases in November, March, and April by the end of the 21st century (fig. 40C).

Clear Creek Soil Moisture Variables

The PRMS projections of infiltration (fig. 23C and 41C) are similar to those of precipitation (fig. 10C and 29C). The large range in the infiltration projections indicates a high level of uncertainty with no significant overall trend (table 5N). High variability in seasonal infiltration also is projected (fig. 41C).

Projected changes in the inflow to the subsurface reservoir (fig. 24C) and recharge to the groundwater reservoir (fig. 25C) are similar to the changes seen in the subsurface (fig. 17C) and groundwater (fig. 18C) flow components, respectively. The projections show decreases in the inflows to the subsurface reservoirs and recharge (tables 5O and 5P). High variability in seasonal recharge is projected (figs. 42C and 43C).

There is a decrease in the projected annual mean soil moisture (fig. 26C) for scenarios A1B and A2 (table 5Q). High variability in seasonal timing of soil moisture is projected (fig. 44C), with decreases in nearly every month.

The PRMS projections indicate that the storage in the subsurface (fig. 27C) and the groundwater reservoirs (fig. 28C) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5O and 5P; 5R and 5S, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45C and 46C).

Discussion of Clear Creek Projections

In the Clear Creek Basin, agricultural water consumption and population growth will result in increasing water demand. The broader-scale effects of climate change on the flow regime of the Clear Creek Basin indicate a slight overall drying of the basin as a consequence of increased evapotranspiration, but the uncertainty associated with the magnitude of this drying is large.

These results did not consider potential future population growth and land-use changes. They also do not answer the question of whether the potentially adverse effects because of climate change can be mitigated with careful land-use planning in the Clear Creek Basin.

The combined effects of climate change and urbanization in the Clear Creek Basin may alter both the quantity and timing of streamflow and has the potential to change the conditions of water quality that supports biological diversity in aquatic communities.

Precipitation Runoff Modeling System Projections for East River, Colorado

The East River Basin at Almont, Colorado is part of the Gunnison River Basin, and an important tributary of the Colorado River. The Gunnison River contributes approximately 40 percent of the streamflow of the Colorado River at the Utah/Colorado State line (Spahr and others, 1999), and the East River accounts for approximately 25 percent of the streamflow of the Gunnison River (Ugland and others, 1991). At USGS streamflow-gaging station 09112500, the basin has a drainage area of 748 km². The basin has been simulated with the PRMS model by Hay and others (1993) and McCabe and Hay (1995). This application contains 71 HRUs with elevations that range from 2,583 to 3,796 meters (table 1).

Mean monthly baseline conditions for the East River Basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from just below -18 °C during December and January to just below 3.5 °C during July and August. Mean monthly maximum temperatures range from approximately -2 °C during December and January to above 20 °C during July. On a mean monthly basis, precipitation ranges from 1.3 mm/day (during June) to 4.4 mm/day (during February). Because of freezing temperatures during the winter months, precipitation falls as snow and accumulates throughout the winter. The majority of annual mean streamflow occurs during May through July, with the peak monthly streamflow occurring during June.

Current (2002) and projected total water demand in the Gunnison River Basin is about equal to the native supply (Colorado Water Conservation Board, 2002a). The East River Basin is representative of many snowmelt dominated, high-elevation basins in Colorado that supply much of the water to downstream users. Because of its importance as a source of water to the Colorado River, the U.S. Geological Survey conducted studies (Leavesley and others, 1992; Hay and others, 1993; McCabe and Hay, 1994; and McCabe and Hay, 1995) of the effects of potential climate change on the water resources of the East River Basin.

Tourism that results from abundant local recreational options is the largest source of revenue in the region (Gunnison Country Chamber of Commerce, 2009). Many of these recreational activities, such as fishing, whitewater boating, snowmobiling, and skiing, are directly dependent on the basin's water resources. The Crested Butte ski area falls within the East River Basin and has a base elevation of 2,856 meters and top elevation of 3,707 meters. The ski area typically receives more than 7 meters of snowfall, and operates from late November to early April.

East River Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the East River Basin (fig. 10D). The large range in the precipitation projections indicates a high level

of uncertainty with a slight increase in the central tendencies shown for the significant A1B scenario (table 5A). Seasonal variability also is apparent in the projections with increasing precipitation in December through February and decreasing precipitation in May and November (fig. 29D).

All PRMS projections indicate steady increases in temperature (figs. 11D, 12D, 30D, 31D, and tables 5B and 5C), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5B and 5C).

All projections indicate an overall increase in annual mean evapotranspiration (fig. 13D and table 5D). High variability in seasonal evapotranspiration is projected, with large increases from October through May followed by greater uncertainty from June through September (fig. 32D).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14D and table 5E). This increase is because of the projected increase of annual temperature (figs. 11D and 12D and tables 5B and 5C). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5E).

East River Streamflow Variables

The PRMS projections indicate a decrease in annual mean streamflow (fig. 15D and table 5F), though the uncertainties associated with these streamflow projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, particularly during the snowmelt months (fig. 33D). Streamflow is projected to increase from March through May and decrease in June through August (fig. 33D). By 2060, peak timing of streamflow is projected to shift from June to May.

The PRMS projections indicate an overall decrease in surface runoff and subsurface flow (tables 5G and 5H) with a high degree of uncertainty (figs. 16D and 17D) for most scenarios. Groundwater flow is projected to decrease (for the A2 emission scenario only; fig. 18D and table 5I). By the end of the 21st century, the projected month of peak surface runoff and subsurface flow shifts from June to May, while the groundwater peak remains in June (figs. 34D–36D).

East River Snow Variables

Because of increases in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19D), percentage of area covered by snow (fig. 20D), snowpack water equivalent (fig. 21D), and snowmelt (fig. 22D) will decrease (tables 5J–5M). The large range in the seasonal projections (figs. 37D–40D) indicates a high degree of uncertainty associated with the projected changes during the snow accumulation and melt season. The largest projected decreases in the percent of precipitation that falls as snow and percentage of area covered by snow are seen in the transition spring and fall months. The PRMS simulation of

baseline conditions (red line, fig. 39) indicates that the timing of maximum snowpack water equivalent occurs in March and April. The PRMS projections indicate by the end of the 21st century, this will shift to March. Also projected are substantial increases in snowmelt during March and April, followed by large decreases in May, and especially June (fig. 40).

East River Soil Moisture Variables

The PRMS projections of infiltration (figs. 23D and 41D) are similar to those seen in the precipitation plots (fig. 10D and 29D). The large range in the infiltration projections indicates a high level of uncertainty with no significant overall trend (table 5N). High variability in seasonal infiltration also is projected (fig. 41D), with large increases during the snow accumulation months followed by large decreases during the snowmelt months.

Projected changes in the inflow to the subsurface reservoir (fig. 24D) and recharge to the groundwater reservoir (fig. 25D) are similar to the changes seen in the subsurface (fig. 17D) and groundwater (fig. 18D) flow components, respectively. The projections show an overall negative annual mean trend in the inflows to the subsurface reservoirs for all scenarios (tables 5O). The large range in the projections of annual mean recharge indicates a high level of uncertainty, with decreases projected for the A2 scenario (table 5P). High variability in seasonal recharge is projected (figs. 42D and 43D).

There is a slight decrease in annual mean soil moisture for A1B and A2 scenarios (fig. 26D and table 5Q). High variability in seasonal timing of soil moisture is projected (fig. 44A), with decreases projected from June through September and substantial increases projected in March and April.

The PRMS projections indicate that the storage in the subsurface (fig. 27D) and the groundwater reservoirs (fig. 28D) reached equilibrium quickly, resulting in their annual mean storage values following the same trends as their inflows (tables 5O and 5P; and 5R, and 5S, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45D and 46D).

Discussion of East River Projections

Streamflow in the East River Basin is under increasing demand from water users and recreationalists within the basin. Decreases in streamflow resulting from future changes in climate may add to the stress that this basin will experience as a result of projected increases in domestic and industrial water use (Colorado Water Conservation Board, 2006a). The effects of climate change in the East River Basin may also result in earlier snowmelt resulting in earlier spring runoff and affecting the conditions that support recreational activities such as skiing.

Precipitation Runoff Modeling System Projections for Feather River, California

The Feather River Basin is located in northern California. The basin has a drainage area of 9,324 km² at the basin outlet reservoir, Lake Oroville. The basin was simulated with the PRMS model by Kocot and others (2005). The PRMS application, which consists of 8 subbasins, contains 324 HRUs with elevations that range from 325 to 2,212 meters (table 1).

The Feather River Basin is a valuable hydrologic resource for California (Kocot and others, 2005). The basin is a major contributor to the California State Water Project, which distributes water throughout California for domestic use, irrigation, and hydropower production. The outlet reservoir, Lake Oroville, holds 8 percent of the state's reservoir capacity and plays an important role in flood management, water quality, and the health of fisheries, affecting areas downstream at least as far south as the Sacramento/San Joaquin River Delta. The climate is Mediterranean, with warm dry summers and cool wet winters. Spring snowmelt from the basin is relied upon to meet the State Water Project's summer water demands.

The Feather River Basin is sensitive to slight changes in temperature which affect the formation and melting of snow. The basin includes large areas that are at or near the seasonal snow line (1,676 meters elevation), and winter rain or rain-on-snow occurrences are common. This basin has been recognized as one of the first in California anticipated to be affected by climate-induced change to snowpack, which will have relatively large effects on the timing and quantity of streamflow (Freeman, 2008).

Mean monthly baseline conditions for the Feather River Basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from -6.5 °C during December and January to 6.3 °C during July. Mean monthly maximum temperatures range from above 4 °C during December and January to above 28 °C during July and August. On a mean monthly basis, precipitation ranges from less than 0.25 mm/day (during July) to above 7 mm/day (during January). Because of winter freezing temperatures in the highest elevations of the basin and near freezing temperatures in mid-elevations of the basin, precipitation falls as snow on higher elevations and as a mixture of rain and snow on mid-elevations; precipitation falls as rain below 1,000 meters. Snow accumulates and melts in this basin throughout the winter, with maximum snowpack accumulation historically occurring by April 1 (Kocot and others, 2005). Corresponding to winter precipitation, the majority of mean annual streamflow occurs during the winter and spring months, with the peak monthly streamflow occurring from March to as late as May (Kocot and others, 2005).

Existing PRMS forecast applications focus on seasonal (3 to 9 months) and medium-range (one week to one month) streamflows (Kocot and others, 2005). Currently (2005), in the Feather River Basin, the California Department of Water Resources makes seasonal and medium-range forecasts of total streamflow into Lake Oroville, and hydroelectric power

operators use their own suite of statistical models to manage power generation within the basin. The National Oceanic and Atmospheric Administration California-Nevada River Forecasting Center also employs the National Weather Service River Forecasting System for flood and water-supply forecasting for the Feather River Basin. Agricultural, fishery, logging, and local user groups may benefit from improved forecasts of long-term climatological trends. Improved understanding of how the Feather River Basin responds to changing climatic conditions will help water managers safeguard this resource.

Feather River Input Variables

Changes in annual mean precipitation in the PRMS projections of the Feather River Basin are not significant, but highly variable (fig. 10E). The large range in the precipitation projections indicates a large amount of uncertainty in the GCM scenarios used in this study with no significant overall trend (table 5A). Seasonal variability also is apparent in these precipitation projections with the largest amount in the winter months (fig. 29E).

All PRMS projections indicate steady increases in temperature (figs. 11E, 12E, 30E, 31E, and tables 5B and 5C), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5B and 5C).

All projections indicate steady increases in annual mean evapotranspiration (fig. 13E table 5D). High variability in seasonal evapotranspiration is projected, with increases projected from October through May, followed by large projected decreases in the central tendencies in June and July (fig. 32E).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14E and table 5E). This increase is because of the projected increase of annual temperature (figs. 11E and 12E and tables 5B and 5C). The projected changes in the length of the growing season are smallest for the B1 scenario and largest for the A2 scenario (table 5E).

Feather River Streamflow Variables

The PRMS projections indicate both increases and decreases in annual mean streamflow with no significant trend (fig. 15E and table 5F), though the uncertainties associated with these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, especially in the winter months, with minimal changes shown for the summer months (fig. 33E). Streamflow is projected to increase from November through March and decrease in May and June. By the end of the 21st century peak monthly streamflow is projected to shift from March to possibly as early as January.

The PRMS projections indicate variable changes in annual mean surface runoff and subsurface flow (figs. 16A and 17A) with no significant trends (table 5G and 5H), while

the annual mean groundwater flow is projected to decrease for the A1B scenario (fig. 18E and table 5I). High variability (indicating uncertainty) in the seasonal timing of these variables is projected (figs. 34E, 35E, and 36E). Surface runoff is projected to increase in December and January by the end of the 21st century. Subsurface flow is projected to increase December through March and decrease April through June, with the timing of the timing shifting from March to possibly January (fig. 35E). Groundwater flow is projected to increase December through April, followed by a decrease May through October (fig. 36E).

Feather River Snow Variables

Because of increases in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19E), percentage of area covered by snow (fig. 20E), snowpack water equivalent (fig. 21E), and snowmelt (fig. 22E) will decrease (tables 5J–5M). On a monthly basis, the central tendencies of the five GCMs and three emission scenarios indicate decreases in all months, with the exception of snowmelt (figs. 37E–40E). Because of projected warmer temperatures, snowmelt is projected to increase December and January and decrease March through June.

Feather River Soil Moisture Variables

The PRMS projections of infiltration (figs. 23E and 41E) are similar to those seen in the precipitation plots (figs. 10E and 29E). The large range in the infiltration projections indicates a high level of uncertainty with no significant overall trend, with the exception of a slight increase in the A1B scenario (table 5N). High variability in seasonal infiltration also is projected (fig. 41E), with increases during snow accumulation season followed by decreases in snowmelt season.

Changes in the inflow to the subsurface reservoir (fig. 24E) and recharge to the groundwater reservoir (fig. 25E) are similar to the changes seen in the subsurface (fig. 17E) and groundwater (fig. 18E) flow components, respectively. The projections show decreases in the recharge to the groundwater reservoirs for the A1B scenario only (table 5P) with a high level of uncertainty.

There is a decrease in the projected annual mean soil moisture (fig. 26E and table 5Q). High variability in seasonal timing of soil moisture is projected (fig. 44E), with increases projected during the winter months followed by decreases from April through September.

The PRMS projections indicate that the storage in the subsurface (fig. 27E) and the groundwater reservoirs (fig. 28E) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5O and 5P; 5R and 5S, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45E and 46E).

Discussion of Feather River Projections

For the Feather River Basin, projections indicate that basin temperatures likely are to increase over the next century, and precipitation likely is to fall on the basin in cycles of wetter and drier decadal oscillations. Because large areas of this basin straddle the seasonal snowline, slight changes in temperature will affect precipitation form. The projections also indicate that the cumulative climatic effects on the basin will likely produce more winter runoff, earlier peak streamflow, and less spring snowmelt. The broader-scale effects of climate change on the flow regime of the Feather River project no significant overall trend in annual mean streamflow in the basin, but the uncertainty associated with this is large. These results did not consider many of the important feedback mechanisms which act between the land surface and the atmosphere, such as the relation between evaporation and relative humidity.

The effects of climate change in Feather River Basin may alter both the quantity and timing of streamflow as has been discussed by Dettinger and Cayan (1995). This could potentially change the conditions that affect: water availability in summer for the California State Water Project; winter high flows; and water quality and quantity for fisheries.

Precipitation Runoff Modeling System Projections for South Fork Flathead River, Montana

The South Fork of the Flathead River Basin is located on the west side of the Continental Divide in northwestern Montana. The South Fork Flathead River flows into the Flathead River and ultimately into the Clark Fork of the Columbia River. At USGS streamflow-gaging station 12362500, the basin has a drainage area of 4,307 km². The basin was simulated with the PRMS model by Chase (2011). The PRMS application for this basin contains 106 HRUs with elevations that range from 1,045 to 2,078 meters (table 1).

Mean monthly baseline conditions for the South Fork Flathead River Basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from approximately -11 °C during December, January, and February to just below 6.0 °C during July and August. Mean monthly maximum temperatures range from -3.6 °C during December to 21 °C during August. On a mean monthly basis, precipitation ranges from 1.9 mm/day (during August) to 5.5 mm/day (during November). Because of freezing temperatures during the winter months, most precipitation falls as snow and accumulates throughout the winter. The majority of annual mean streamflow occurs during April through July, with the peak monthly streamflow occurring during May or June.

The South Fork Flathead River Basin mostly is undeveloped and covered with forests, which have been affected by wildfires and mountain pine beetles during the 1998–2008 drought. Hungry Horse Reservoir, on the lower end of the

South Fork Flathead River, stores water behind Hungry Horse Dam. Hungry Horse Dam is operated by the Bureau of Reclamation for power generation, flood control, recreation, and flow augmentation for endangered species. Releases from Hungry Horse Dam must meet minimum flow requirements below Hungry Horse Dam and on the main stem Flathead River at Columbia Falls for bull trout, and must provide spring and summer flow augmentation for salmon and steelhead in the Columbia River (Mary Mellema, Bureau of Reclamation, written commun., 2010). These uses together with potential future uses of Hungry Horse storage water pose many challenges to water resources planners and managers.

This PRMS application is used to simulate runoff in the South Fork Flathead basin for forecasting inflow to Hungry Horse Reservoir.

Flathead River Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the South Fork Flathead River (fig. 10F), with an overall increase in the central tendencies projected for the A1B and A2 scenarios (table 5A). The large range in the precipitation projections indicates a large amount of uncertainty. The high variability in the precipitation projections are demonstrated in the monthly plot, with a tendency towards increasing precipitation by the end of the 21st century in all but the summer months (fig. 29F).

All PRMS projections indicate steady increases in temperature (figs. 11F, 12F, 30F, 31F, and tables 5B and 5C), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5B and 5C).

All projections indicate steady increases in annual mean evapotranspiration (fig. 13F and table 5D). High variability in seasonal evapotranspiration is projected, with projected increases in all months except July through September. In July through September evapotranspiration is limited by the decreases in precipitation (figs. 29F and 32F).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14F and table 5E). This increase is because of the projected increase of annual temperature (figs. 11F and 12F and tables 5B and 5C). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5E).

Flathead River Streamflow Variables

Though the USGS streamflow-gaging station 12362500 is downstream from Hungry Horse Reservoir, natural streamflows (without the effect of storage in the reservoir) were simulated in PRMS. The PRMS projections indicate an increase in annual mean streamflow for the A1B (fig. 15F and table 5F), though the uncertainties associated with these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, particularly during the months of

April, May, and June (fig. 33*F*). By the end of the 21st century, streamflow is projected to increase from November through April and decrease in May through July. The PRMS simulation of baseline conditions (red line, fig. 33*F*) indicate that peak streamflow occurs in May and June, whereas the PRMS projections indicate that, by the year 2030, the projected month of peak streamflow will shift to May.

In general, the PRMS projections indicate increases in surface runoff and groundwater flow (figs. 16*F* and 18*F* and tables 5*G* and 5*I*). Streamflow is dominated by the subsurface flow component which projects no significant trend (table 5*H*). High variability (indicating uncertainty) in the seasonal timing of these variables is projected (figs. 34*F* and 35*F*), with large increases in surface runoff during November, February, April and May by the end of the 21st century. Projections of subsurface flow indicate increases during November through April followed by large decreases during May through July (fig. 35*F*), with groundwater flow increasing in all months (fig. 36*F*).

Flathead River Snow Variables

Because of the increase in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19*F*), percentage of area covered by snow (fig. 20*F*), and snowpack water equivalent (fig. 21*F*) will decrease (tables 5*J*–5*L*). The projections show no significant trend for annual mean snowmelt (fig. 22*F* and table 5*M*). Warmer temperatures will result in less snowfall and increased snowmelt November through April, so that less snow is available to melt in May, June, and July (fig. 40*F*).

Flathead River Soil Moisture Variables

The PRMS projections of infiltration (figs. 23*F* and 41*F*) are similar to those of precipitation (figs. 10*F* and 29*F*). The large range in the infiltration projections indicates a high level of uncertainty with a positive trend in the A1B scenario (table 5*N*). High variability in seasonal infiltration also is projected (fig. 41*F*), with large increases projected from October through April and large decreases projected for June and July.

Projected changes in the inflow to the subsurface reservoir shows no significant trend (fig. 24*F* and table 5*O*), though projected mean monthly inflow to the subsurface reservoir is projected to increase November through April and decrease May through July (fig. 42*F*). A significant positive trend in groundwater reservoir recharge is projected for the A1B and A2 scenarios (fig. 25*F* and table 5*P*). High variability in seasonal recharge is projected (fig. 43*F*).

There is a decrease in the projected annual mean soil moisture for the A1B and A2 scenarios (fig. 26*F* and table 5*Q*). High variability in seasonal timing of soil moisture is projected (fig. 44*A*), with substantial decreases projected from May through October and substantial increases projected in February and March.

The PRMS projections indicate that storage in the sub-surface reservoirs (fig. 27*F* and table 5*R*) and the groundwater reservoirs (fig. 28*F* and table 5*S*) both significantly increase. High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45*F* and 46*F*).

Discussion of Flathead River Results

Hungry Horse Reservoir, on the South Fork Flathead River, is an important component in the Columbia River system of reservoirs operated by the Bureau of Reclamation. Even as demand for the water stored in the reservoir is increasing, climate projections indicate that the quantity and timing of streamflow entering the reservoir could be changing. PRMS projections indicate that flows could increase in the winter and early spring and decrease in late spring and through the summer. Uncertainty associated with the magnitude of these changes in streamflow is large. These results did not consider potential future land-cover dynamics, such as forest fire or pine beetle damage. The combined effects of climate change and land-cover changes in the South Fork Flathead River Basin could affect water available for use in Montana and in downstream states.

Precipitation Runoff Modeling System Projections for Flint River, Georgia

The Flint River Basin is located in Macon County Georgia, above Montezuma. Downstream from Montezuma, the Flint River joins the Chattahoochee River and ultimately drains into the Gulf of Mexico. At USGS streamflow-gaging station 02349500, the basin has a drainage area of 7511 km². The basin was simulated with the PRMS model by Viger and others (2010). The PRMS application contains 128 HRUs with elevations that range from 90 to 286 meters (table 1).

Mean monthly baseline conditions in the Flint River Basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from just below 2.0 °C during December and January to almost 21 °C during July. Mean monthly maximum temperatures range from approximately 14.0 °C during December and January to over 32 °C during July. On a mean monthly basis precipitation is highly variable, with ranges from below 2.0 mm/day (during May and August) to more than 4.0 mm/day (during February, March, and July). The majority of annual mean streamflow occurs during February and March.

The upper portion of the Flint River flows unimpeded by major impoundments for about 320 river kilometers and provides habitat for biologically diverse communities that have been lost in the impounded reaches of many other rivers in the eastern United States. River shoal habitats in the upper Flint support a variety of native fishes, mussels, and aquatic plants. The upper Flint River Basin supplies water to a growing population in the Atlanta metropolitan area and is a major recreational resource for the region (Gregory and others,

2006). Urban growth in the headwaters, which encompass a portion of the Atlanta metropolitan area, and increasing demands for offstream water use will affect future flow regimes in the upper Flint River.

The USGS selected “Water availability for ecological needs in the Upper Flint River Basin, Georgia” as a national priority science thrust program starting in 2006 (Gregory and others, 2006). The Flint River Science project is part of a federally funded program to address key national science priorities including landslides/debris flows, fire science, integrated landscape monitoring, and water availability. The fundamental purpose of this project is to advance the science needed to specify the hydrologic conditions necessary to support flowing-water ecosystems; information that is critical for management of this water resource.

Flint River Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the upper Flint River Basin (fig. 10*G*), with two of three scenarios having significant decreases in mean annual precipitation (table 5*A*). Seasonal variability also is apparent in these precipitation projections (fig. 29*G*).

All PRMS projections indicate steady increases in temperature (figs. 11*G*, 12*G*, 30*G*, 31*G*, and tables 5*B* and 5*C*), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5*B* and 5*C*).

All projections indicate steady increases in annual mean evapotranspiration (table 5*D*), with a large amount of uncertainty, especially in the A2 and A1B scenarios by the end of the 21st century (fig. 13*G*). High variability in seasonal evapotranspiration is projected, with increases in evapotranspiration during October through April and July (fig. 32*G*).

All PRMS projections indicate steady increases in the length of the growing season (fig 14*G* and table 5*E*). This increase is because of the projected increase of annual temperature (figs. 11*G* and 12*G* and tables 5*B* and 5*C*). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5*E*).

Flint River Streamflow Variables

The PRMS projections indicate a decrease in annual mean streamflow and corresponding flow components for all but the B1 scenarios (figs. 15*G*–18*G* and tables 5*F*–5*I*), though the uncertainties associated with these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected. Increased flows in July are because of increased precipitation in both the baseline conditions and the projections (fig. 33*G*).

Flint River Snow Variables

The Flint River Basin receives a small amount of snow. Projections of the snow related variables (figs. 19*G*–22*G* and 37*G*–40*G*) are difficult to interpret because of the small amounts of snow present (simulations indicate that less than 1 percent of the basin is snow covered in January).

Flint River Soil Moisture Variables

The PRMS projections of infiltration (figs 23*G* and 41*G*) are similar to those of precipitation (figs. 10*G* and 29*G*). The large range in the infiltration projections indicates a high level of uncertainty with a significant overall negative trend for the A2 scenario (table 5*N*). High variability in seasonal infiltration also is projected (fig. 41*G*).

Projected changes in the inflow to the subsurface reservoir (fig. 24*G*) and recharge to the groundwater reservoir (fig. 25*G*) are similar to the changes seen in the subsurface (fig. 17*G*) and groundwater (fig. 18*G*) flow components, respectively. The projections show decreases in annual mean inflows to the subsurface reservoirs and recharge for all but the B1 scenario (tables 5*O* and 5*P*). High variability in seasonal recharge is projected (figs. 42*G* and 43*G*).

There is a significant decrease in the trends in annual mean soil moisture (fig. 26*G* and table 5*Q*) for the A1B and A2 scenarios, though the uncertainties are large and increase with time (fig. 26*G*). High variability in seasonal timing of soil moisture is projected (fig. 44*G*), with decreases in every month.

The PRMS projections indicate that the storage in the subsurface (fig. 27*G*) and the groundwater reservoirs (fig. 28*G*) reached equilibrium quickly, resulting in simulated values similar to the inflows (tables 5*O* and 5*P* compared to tables 5*R* and 5*S*). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45*G* and 46*G*).

Discussion of Flint River Results

In the upper Flint River watershed, increasing demand for water from the steady growth of the Atlanta metropolitan area has the potential to alter streamflow throughout the watershed. Increased development, combined with severe droughts in the area, has resulted in shortages and restrictions on the limited surface-water supplies.

The broader-scale effect of climate change on the flow regime of the Flint River indicates a slight drying of the watershed, but the variability associated with the magnitude of this projected drying is large. These results did not consider potential future land-cover dynamics. They also do not answer the question of whether the potentially adverse effects because of climate change, combined with the added urbanization stresses, can be mitigated with careful land-use planning in the Flint River Basin.

The combined effects of decreased precipitation (particularly during the summer months) and urbanization in the Flint River may alter both the quantity and timing of streamflow. This has the potential to change the conditions that support biological diversity in aquatic communities.

Precipitation Runoff Modeling System Projections for Naches River, Washington

The Naches River Basin is a tributary to the Yakima River Basin in Washington. Much of the Naches River Basin lies on the east-side slopes of the Cascade Mountains. At USGS streamflow-gaging station 12494000, the basin has a drainage area of 2,437 km². The basin was simulated with the PRMS model by Mastin and Vaccaro (2002). The PRMS application contains 363 HRUs with elevations that range from 562 to 1,650 meters (table 1).

Mean monthly baseline conditions in the Naches River Basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from below -7.0 °C during December to above 7.5 °C during July and August. Mean monthly maximum temperatures range from below -1.0 during December to 22 °C during July and August. On a mean monthly basis, precipitation ranges from less than 0.4 mm/day (during August) to more than 4.0 mm/day (during December). Because of freezing temperatures during the winter months, precipitation falls as snow and accumulates throughout the winter. The majority of annual mean streamflow occurs during April through June, with the peak monthly streamflow occurring during May.

Agriculture is the principal economic activity in the Yakima River Basin with about 200 km² in the low-lying semiarid-to-arid parts of the basin. Agriculture in the basin depends on irrigation with a demand of about 3,200 million cubic-meters of water that is mostly supplied with surface water. Five principle reservoirs in the upper portions of the basin with a capacity of 1,300 million cubic-meters supplement flows in the rivers to supply irrigators and provide adequate in streamflows for anadromous fish habitat (Mastin and Vaccaro, 2002). Sufficient water to meet the demands of the basin in any one year is dependent highly on an adequate spring snowpack.

Naches River Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Naches River Basin (fig. 10H). The large range in the precipitation projections indicates a high level of uncertainty with a trend towards increasing precipitation by the latter half of the 21st century for the A2 and A1B scenarios and (table 5A). Seasonal variability also is apparent in these precipitation projections with increases during the winter months and the largest decreases during July (fig. 29H).

All PRMS projections indicate steady increases in temperature (figs. 11H, 12H, 30H, 31H, and tables 5B and 5C), with associated uncertainties increasing with time. The

projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5B and 5C).

All projections indicate steady increases in annual mean evapotranspiration for the A2 and B1 scenarios (fig. 13H and table 5D). High variability in seasonal evapotranspiration is projected, with the increases in March through April, and the decreases in July and August (fig. 32A).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14H and table 5E). This increase is because of the projected increase of annual temperature (figs. 11H and 12H and tables 5B and 5C). The projected changes in the length of the growing season are smallest for the B1 scenario and largest for the A2 scenario (table 5E).

Naches River Streamflow Variables

The PRMS projections indicate a slight increase in annual mean streamflow for the A1B and A2 scenarios (fig. 15H and table 5F), though the uncertainties associated with all of these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, with flows increasing in November through March and decreasing in May through July (fig. 33H). The PRMS simulation of baseline conditions (red line, fig. 33H) indicate that peak streamflow occurs in May. The PRMS projections indicate that, by the end of the 21st century, the projected month of peak streamflow occurring as early as February.

The PRMS projections indicate an increase in annual mean surface runoff for all scenarios (fig. 16G and table 5G), while subsurface flow will increase for the A2 scenario and a decrease for the B1 scenario (fig. 17G and table 5H). Groundwater flow will increase for the A1B and A2 scenarios (fig. 18H and table 5I). High variability (indicating uncertainty) in the seasonal timing of these variables is projected with surface runoff projected to increase from October through June (fig. 24H). Subsurface flow will increase during November through March and decrease during April through July. The PRMS simulation of baseline conditions (red line, fig. 33H) indicate that peak streamflow occurs in May, while the PRMS projections indicate that this may shift to February. Groundwater flow will increase from December through March, followed by a decrease from May through August (fig. 36H).

Naches River Snow Variables

Because of increases in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19H), percentage of area covered by snow (fig. 20H), snowpack water equivalent (fig. 21H), and snowmelt (fig. 22H) will decrease (tables 5J–5M). Seasonal projections indicate decreases in the percent of precipitation that falls as snow, percentage of area covered by snow, and snowpack water equivalent during all snow accumulation and melt months (figs. 37H–39H). Snowmelt is projected to

increase during December through March, followed by large decreases during April and May and smaller decreases in June (fig. 40H).

Naches River Soil Moisture Variables

The PRMS projections of infiltration (figs. 23H and 41H) are similar to those of precipitation (figs. 10H and 29H). The large range in the infiltration projections indicates a high level of uncertainty with a positive trend (table 5N). High variability in seasonal infiltration also is projected (fig. H), with large increases projected November through February and decreases in July.

Projected changes in the inflow to the subsurface reservoir (fig. 24H) and recharge to the groundwater reservoir (fig. 25H) are similar to the changes seen in the subsurface (fig. 17H) and groundwater (fig. 18H) flow components, respectively. The projections show increases for inflow to the subsurface reservoir for the A1B and A2 emissions (table 5O and 5P). High variability in seasonal recharge is projected (figs. 42H and 43H).

There is a decrease in the projected annual mean soil moisture (fig. 26H and table 5Q). High variability in seasonal timing of soil moisture is projected (fig. 44A), with increases in winter months, because of melting snow, and decreases during April through October.

The PRMS projections indicate that the storage in the subsurface (fig. 27H) and the groundwater reservoirs (fig. 28H) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5O and 5P; 5R and 5S, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45H and 46H).

Discussion of Naches River Results

The PRMS projections show seasonal changes in snow and rain dynamics for the Naches River Basin. Projections indicate that increases in air temperature will result in increased winter runoff because the form of the precipitation will change from snow to rain. When precipitation falls as rain, it becomes streamflow more quickly than if it had fallen as snow and accumulated in the snowpack. Also, more rain and less snow means less runoff in the late spring and summer. This has implications for water managers who rely on the snowmelt from the spring snowpack to replenish reservoirs and provide early season irrigation. Whereas projections indicate an increase in total annual runoff, runoff during the spring is less, and therefore, water managers will need to store more winter runoff to meet the irrigation and in-stream flow demands. The effects of climate change in the Naches River basin may alter the timing of streamflow and has the potential

to change the conditions that support biological diversity in aquatic communities.

Precipitation Runoff Modeling System Projections for Pomperaug River, Connecticut

The Pomperaug River Basin is located in western Connecticut. The river flows into the Housatonic River which is a tributary to the Long Island Sound. At USGS streamflow-gaging station 01204000, the basin has a drainage area of 194 km². The basin was simulated by Bjerklie and others (2009). The PRMS application contains 55 HRUs with elevations that range from 63 to 332 meters (table 1).

Mean monthly baseline conditions in the Pomperaug River Basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from approximately -7.5 °C during February to just above 16.0 °C during July. Mean monthly maximum temperatures range from just below 3.0 °C during January to above 28.0 °C during July. On a mean monthly basis precipitation is highly variable, with ranges from less than 3.4 mm/day (during February, June, July, and December) to more than 4.25 mm/day (during March, August, and November). Because of freezing temperatures during the winter months, precipitation falls as snow and accumulates throughout the winter. The majority of annual mean streamflow occurs during the winter and spring, with the peak monthly streamflow occurring during March.

The basin supports significant areas of groundwater withdrawal from wells completed in stratified glacial deposits, and numerous individual domestic wells completed in bedrock. Water withdrawals for commercial and industrial use are almost entirely from groundwater; however, withdrawals for agricultural use come mostly from surface water. Portions of eight towns are included within the basin boundaries. Land use and land cover in the basin has changed over the last century from primarily agriculture to primarily forest with an increasing, but still relatively small, percentage of urban and residential use (Bjerklie and others, 2009).

In recent years, increased use of surface-water and groundwater supplies in the Pomperaug River Basin has created concern that insufficient flows remain for future development, aquatic habitat, and recreational use. In response, the USGS, in cooperation with the Pomperaug River Watershed Coalition developed a PRMS application for the Pomperaug River as part of a study to evaluate the potential effects of land-use change and water management strategies on streamflow, in-stream habitat, and groundwater availability in the Pomperaug River Basin. Although the basin is rural, the population increased by 15 percent during the last decade as a result of suburban sprawl (Bjerklie and others, 2009). Increases in population and development can lead to changes in land use and land cover that causes change in the distribution of runoff

and groundwater recharge. Unless these cumulative effects are considered when allocating the water resources, in-stream flow losses and degradation of water quality and ecosystems can occur. Because of uncertainty about the quantity of water available for future use, the effects of new applications for water diversion are poorly understood.

Pomperaug River Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Pomperaug River Basin (fig. 10I). The large range in the precipitation projections indicates a high level of uncertainty with a significant positive trend for the B1 scenario (table 5A). Seasonal variability also is apparent in these precipitation projections (figure 29I).

All PRMS projections indicate steady increases in temperature (figs. 11I, 12I, 30I, and 31I and tables 5B and 5C), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5B and 5C).

All projections indicate steady increases in annual mean evapotranspiration (fig. 13I and table 5D). High variability in seasonal evapotranspiration is projected with increases in all months (figure 32I).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14I and table 5E). This increase is because of the projected increase of annual temperature (figs. 11I and 12I and tables 5B and 5C). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5E).

Pomperaug River Streamflow Variables

The PRMS projections indicate a decrease in annual mean streamflow for the A2 scenario (fig. 15I and table 5F), though the uncertainties associated with these streamflow projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, with a decrease in all months with the exception of January (fig. 33I).

The PRMS projections of surface runoff vary (fig. 16I) with no significant trend (table 5G). With the exception of increases in January, the changes are minimal (fig. 34I). Subsurface and groundwater flow decrease over time (tables 5F, 5H, and 5I). High variability (indicating uncertainty) in the seasonal timing of these variables is projected with decreases in all months for the groundwater flow (figure 36I).

Pomperaug River Snow Variables

Because of the increase in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19I), percentage of area covered by snow (fig. 20I), snowpack water equivalent (fig. 21I), and snowmelt (fig. 22I) will decrease (tables 5J–5M). Projected warmer temperatures will: (1) result in less snowfall,

which will result in less snow available to melt in the spring; (2) decrease the amount of precipitation that falls as snow, percentage of area covered by snow, and snowpack water equivalent in every month (figs. 37I–39I); and (3) increase the snowmelt in January (fig. 40A), which is consistent with the projections of monthly streamflow (fig. 33I) and surface runoff (fig. 34I).

Pomperaug River Soil Moisture Variables

The PRMS projections of infiltration (figs. 23I and 41I) are similar to those of precipitation (figs. 10I and 29I). The large range in the infiltration projections indicates a high level of uncertainty with a positive trend for the B1 scenario (table 5N). High variability in seasonal infiltration also is projected (fig. 41I).

Projected changes in the inflow to the subsurface reservoir (fig. 24I) and recharge to the groundwater reservoir (fig. 25I) are similar to the changes seen in the subsurface (fig. 17I) and groundwater (fig. 18I) flow components, respectively. The projections show no significant trends in the inflows to the subsurface reservoirs (table 5O). The large range in the projections of annual mean recharge indicates a high level of uncertainty, with decreases projected for the A1B and A2 scenarios (table 5P). High variability in seasonal recharge is projected (figs. 42A and 43A).

There is a decrease in the projected annual mean soil moisture (fig. 26I and table 5Q). High variability in seasonal timing of soil moisture is projected (fig. 44I), with decreases in nearly every month. The largest decreases occur during the summer months when potential evapotranspiration demand is highest.

The PRMS projections indicate that the storage in the subsurface (fig. 27I) and the groundwater reservoirs (fig. 28I) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5O and 5P; 5R and 5S, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45I and 46I).

Discussion of Pomperaug River Results

In the Pomperaug River Basin, increasing demand for groundwater from steady population growth has the potential to reduce groundwater base flow to streams throughout the watershed. Increasing stress on groundwater and base flow directly may affect the capacity to supply high quality groundwater to residents and at the same time not be able to preserve adequate base flow to streams for aquatic habitat. The broader-scale effects of climate change on the flow regime of the Pomperaug River indicate an overall slight drying of the basin with a somewhat larger reduction in groundwater relative to streamflow, however the uncertainty associated with the magnitude of this drying is large. These results did not consider potential future population and land-use changes. The results also do not answer the question of whether the potentially

adverse effects because of climate change can be mitigated with careful basinwide land-use planning. The combined effects of climate change and urbanization in the watershed may alter both the quantity and timing of streamflow and has the potential to change the conditions that support biological diversity in aquatic communities.

Precipitation Runoff Modeling System Projections for Sagehen Creek, California

The Sagehen Creek Basin is located on the east slope of the northern Sierra Nevada and is a tributary to the Truckee River in California. At USGS streamflow-gaging station 10343500, the basin has a drainage area of 27 km². The basin was simulated with a fully coupled groundwater/surface-water model (Markstrom and others, 2008). The PRMS portion of the GSFLOW application contains 128 HRUs with elevations that range from 1,941 to 2,589 meters (table 1).

Mean monthly baseline conditions in the Sagehen Creek basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from less than -9.5 °C during December and February to just above 5.0 °C during July and August. Mean monthly maximum temperatures range from 2.0 °C during December to over 24.0 °C during July and August. On a mean monthly basis, precipitation ranges from less than 0.25 mm/day (during July) to almost 8.0 mm/day (during January). Because of freezing temperatures during the winter months, precipitation falls as snow and accumulates throughout the winter. The majority of annual mean streamflow occurs from April through June, with the peak monthly streamflow occurring during May.

Sagehen Creek Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Sagehen Creek Basin (fig. 10J). The large range in the precipitation projections indicates a high level of uncertainty with a significant positive trend for the A1B scenario (table 5A). Seasonal variability also is apparent in these precipitation projections (fig. 29J), with the largest potential increases and uncertainties in precipitation in the winter months.

All PRMS projections indicate steady increases in temperature (figs. 11J, 12J, 30J, 31J, and tables 5B and 5C), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5B and 5C).

All projections indicate steady increases in annual mean evapotranspiration (fig. 13J and table 5D). High variability in seasonal evapotranspiration is projected, with increases during December through May and decreases during July through September (fig. 32J).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14J and table 5E). This increase is because of the projected increase of annual

temperature (figs. 11J and 12J and tables 5B and 5C). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5E).

Sagehen Creek Streamflow Variables

The PRMS projections indicate no significant overall trend in annual mean streamflow (fig. 15J and table 5F), though the uncertainties associated with these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, with increases during January through April, and decreases during May and June. The PRMS simulation of baseline conditions (red line, fig. 33A) indicate that peak streamflow occurs in May, while the PRMS projections indicate that, by the end of the 21st century, peak streamflow will occur in April.

The PRMS projections indicate a positive trend in annual mean surface runoff for the A1B and B1 scenarios (fig. 16J and table 5G), whereas the subsurface and groundwater flow (figs. 17J and 18J) have no trend (tables 5H–5J).

High variability (indicating uncertainty) in the seasonal timing of these variables is projected (figs. 34A and 35A), with increases in January surface runoff (fig. 34J). Subsurface flow is the dominant component of streamflow and has a similar seasonal pattern (figs. 15J and 18J, table 5F and 5H). The month of peak subsurface flow is projected to shift from May to April. High variability in the seasonal timing of groundwater flow also is projected (fig. 36J).

Sagehen Creek Snow Variables

Because of the increase in temperatures the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19J), percentage of area covered by snow (fig. 20J), and snowpack water equivalent (fig. 21J) will decrease (tables 5J–5L) with the exception of the A2 scenario for snowpack water equivalent (no significant trend). Projected warmer temperatures will: (1) result in less snowfall, which will result in less snow available to melt in the spring; (2) decrease the amount of precipitation that falls as snow, percentage of area covered by snow, and snowpack water equivalent in every month (figs. 37J–39J); and (3) increase the snowmelt during November through March, leaving less snow to melt during May and June (fig. 40J).

Sagehen Creek Soil Moisture Variables

The PRMS projections of infiltration (figs. 23J and 41J) are similar to those of precipitation (figs. 10J and 29J). The large range in the infiltration projections indicates a high level of uncertainty with a positive trend for the A1B scenario (table 5N). High variability in seasonal infiltration also is projected (fig. 41J), with projected increases during the snow accumulation season.

Projected changes in the inflow to the subsurface reservoir (fig. 24*J*) and recharge to the groundwater reservoir (fig. 25*J*) are similar to the changes seen in the subsurface (fig. 17*J*) and groundwater (fig. 18*J*) flow components, respectively. The projections show no significant trends in inflow to these reservoirs (tables 5*O* and 5*P*). High variability in seasonal recharge is projected (figs. 42*J* and 43*J*).

There is a decrease in the projected annual mean soil moisture for the A1B scenario (fig. 26*J* and table 5*Q*). High variability in seasonal timing of soil moisture is projected (fig. 44*J*), with increases during December through March and decreases during May through October.

The PRMS projections indicate that the storage in the subsurface (fig. 27*J*) and the groundwater reservoirs (fig. 28*J*) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5*O* and 5*P*; 5*R* and 5*S*, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45*J* and 46*J*).

Discussion of Sagehen Creek Results

In northern Nevada and central California increased development, combined with severe droughts, have resulted in shortages and restrictions on limited surface-water supplies. The broader-scale effects of climate change on the flow regime and water-balance components of Sagehen Creek indicate a shift in timing of streamflow from May to April with no overall discernible trend in annual flows. The uncertainties associated with the projections are large. These results did not consider many of the important feedback mechanisms which act between the land surface and the atmosphere.

Precipitation Runoff Modeling System Projections for Sprague River, Oregon

The Sprague River Basin is located in the northeastern headwaters of the Klamath River in Oregon. At USGS streamflow-gaging station 11501000, the basin has a drainage area of 4,053 km². The basin was simulated with the PRMS model by Hay and others (2009). The PRMS application contains 102 HRUs with elevations that range from 1,316 to 2,203 meters (table 1).

Mean monthly baseline conditions in the Sprague River basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from less than -7.0 °C during December to above 6.5 °C during July and August. Mean monthly maximum temperatures range from 3.2 °C during December to over 26.0 °C during July and August. On a mean monthly basis, precipitation ranges from less than 0.4 mm/day (during July) to almost 2.6 mm/day (during January). Because of freezing temperatures during the winter months, precipitation falls as snow and accumulates throughout the winter. The majority of annual mean streamflow occurs in spring, with the peak monthly streamflow occurring during April.

The Sprague River flows to the west draining inactive volcanoes, rims, scarps, buttes, and fault-block mountains along the northern, eastern, and southern basin boundaries. Lower elevation benches and tablelands are used as rangeland and bottomlands along the main stream and major tributaries are used for irrigated agriculture. Conifer forests, which cover approximately 80 percent of the basin area, are predominant on the slopes of buttes and mountains. The Sprague River Basin supplies approximately 25 percent of total inflow to the Upper Klamath Lake. Demand for water from the lake has increased in recent years in response to needs such as irrigated agriculture, minimum downstream flow for threatened Coho salmon, water supply for nearby wildlife refuges, hydropower, and a minimum lake level for habitat and water-quality protection of two endangered sucker species (Risley and others, 2005).

Sprague River Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Sprague River Basin (fig. 10*K*). The large range in the precipitation projections indicates a high level of uncertainty with a significant positive trend for the A1B scenario (table 5*A*). Seasonal variability also is apparent in these precipitation projections with increases in precipitation during the winter months (fig. 29*K*).

All PRMS projections indicate steady increases in temperature (figs. 11*K*, 12*K*, 30*K*, 31*K*, and tables 5*B* and 5*C*), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5*B* and 5*C*). Projections for the A1B scenario indicate steady increases in annual mean evapotranspiration (fig. 13*K* and table 5*D*). High variability in seasonal evapotranspiration is projected, with increases projected from November through April, followed by decreases in June through September (fig. 32*K*).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14*K* and table 5*E*). This increase is because of the projected increase of annual temperature (figs. 11*K* and 12*K* and tables 5*B* and 5*C*). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5*E*).

Sprague River Streamflow Variables

The PRMS projections indicate no significant trend in annual mean streamflow (fig. 15*K* and table 5*F*), though the uncertainties associated with these projections are large. The large range in the streamflow projections indicates a large amount of uncertainty. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, with increases during December through March and decreases from April through June (fig. 33*K*). The PRMS simulation of baseline conditions (red line, fig. 33*K*) indicates that peak streamflow occurs in April, whereas the PRMS projections indicate

that, by the end of the 21st century, the projected month of peak streamflow will be March.

The PRMS projections indicate increases in the surface runoff for the A1B and A2 scenarios (fig. 16K and table 5G), whereas there is no trend in the subsurface and groundwater flow (figs. 17K and 18K and tables 5H and 5I). High variability (indicating uncertainty) in the seasonal timing of these variables is projected (figs. 34K–36K), with increases in surface runoff and subsurface flow during December through March, decreases in subsurface flow during April through June, and increases in groundwater flow during December through May.

Sprague River Snow Variables

Because of increases in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19K), percentage of area covered by snow (fig. 20K), snowpack water equivalent (fig. 21K), and snowmelt (fig. 22K) will decrease (tables 5J–5M). Seasonally, precipitation that falls as snow (fig. 37K), percentage of area covered by snow (fig. 38K), and snowpack water equivalent (fig. 39K) are all projected to decrease during the snow accumulation and melt seasons. The projections also indicate that snowmelt (fig. 40K) will increase in February and decrease during November, January, and March through June by the end of the 21st century. Warmer temperatures will result in less snowfall resulting in less snow available to melt in the spring.

Sprague River Soil Moisture Variables

The PRMS projections of infiltration (figs. 23K and 41K) are similar to those of precipitation (figs. 10A and 29A). The large range in the infiltration projections indicates a high level of uncertainty with a significant increase in infiltration for the A1B scenario (table 5N). High variability in seasonal infiltration also is projected (fig. 41K), with increases in infiltration during snow accumulation season followed by decreases in the snowmelt season.

Projected changes in the inflow to the subsurface reservoir (fig. 24K) and recharge to the groundwater reservoir (fig. 25K) are similar to the changes seen in the subsurface (fig. 17K) and groundwater (fig. 18K) flow components, respectively. The projections show no significant trends in inflow to these reservoirs (tables 5O and 5P). High variability in seasonal recharge is projected (figs. 42K and 43K).

There is a decrease in the projected annual mean soil moisture for the A1B scenario (fig. 26K and table 5Q) though the uncertainties associated with these projections are large. High variability in seasonal timing of soil moisture is projected (fig. 44K), with increases during December through March, followed by decreases during April through September.

The PRMS projections indicate that the storage in the subsurface (fig. 27K) and the groundwater reservoirs (fig. 28K) reach equilibrium quickly, resulting in simulated

values similar to the inflows (see tables 5O and 5P; 5R and 5S, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45K and 46K).

Discussion of Sprague River Results

Supplying approximately 25 percent of inflow to the Upper Klamath Lake, the Sprague River Basin is vital to environmental and human water needs within the Klamath River Basin. As water demands increase, the reliability of flows from headwater basins like the Sprague becomes increasingly critical in water-management decisions. The broader scale effects of climate change on the flow regime of the Sprague River indicate increased annual high flows earlier in the spring as overall basin storage decreases. However, the uncertainty associated with these changes is large. It is apparent that with significant changes in seasonal patterns of inflow to the Upper Klamath Lake, one or a combination of the following measures may be necessary: (1) modification of the operation of the lake as a storage reservoir, (2) creation of additional storage capacity to meet water demands, and (3) reprioritization of water deliveries for environmental and human needs.

Precipitation Runoff Modeling System Projections for Starkweather Coulee, North Dakota

The Starkweather Coulee Basin is a major basin within the Devils Lake Basin in northeastern North Dakota. At USGS streamflow-gaging station 05056239, the basin has a drainage area of 543 km² of which about 259 km² probably does not contribute to streamflow (Vining, 2002). The basin was simulated with the PRMS model by Vining (2002). The PRMS application contains 50 HRUs with elevations that range from 446 to 491 meters (table 1).

Mean monthly baseline conditions in the Starkweather Coulee basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from less than –20.0 °C during December, January, and February to above 10.0 °C during July. Mean monthly maximum temperatures range from just above –10.0 °C during January to almost 27.0 °C during August. On a mean monthly basis, precipitation ranges from approximately 0.6 mm/day (December, January, and February) to over 3.0 mm/day (during July). Because of freezing temperatures during the fall and winter months, precipitation falls as snow and accumulates throughout the winter. The majority of annual mean streamflow occurs during April and May, with the peak monthly streamflow occurring during April.

The topography of the basin is mostly level to slightly rolling in the south, somewhat more rolling in the central, and mostly level again in the north, with predominately loam to silty clay soils throughout the basin. Thousands of small depressions and wetlands exist on the surface. Many of the original depressions and wetlands were drained years ago and currently are being farmed. The majority of the land in

the basin is used in production agriculture, although some land areas are used as pasture or are enrolled in conservation programs (Vining, 2002).

The region around the Devils Lake Basin has experienced intermittent flooding conditions since the summer of 1993 during which copious amounts of precipitation filled many lakes, wetlands, and depressions. The low-relief topography of the region and the lack of a prominent outlet from Devils Lake means that precipitation and runoff water remains within the basin. Millions of dollars have been spent protecting infrastructure around Devils Lake from rising waters (Vining, 2002). Alternatives have been proposed for stabilizing basin water levels including the re-establishment and expansion of wetlands and depressions in the upper basin, including Starkweather Coulee basin, to store runoff water. Increases in the amount of surface-water storage would likely decrease runoff to Devils Lake, but if the surface-water storage capacity is exceeded, then the wetlands and depressions likely will discharge most of the runoff water received.

Starkweather Coulee Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Starkweather Coulee Basin (fig. 10*L*). The large range in the precipitation projections indicates a high level of uncertainty with no significant overall trend (table 5*A*). Seasonal variability also is apparent in these precipitation projections (fig. 29*L*).

All PRMS projections indicate steady increases in temperature (figs. 11*L*, 12*L*, 30*L*, 31*L*, and tables 5*B* and 5*C*), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5*B* and 5*C*).

All projections indicate steady increases in annual mean evapotranspiration (fig. 13*L* and table 5*D*). High variability in seasonal evapotranspiration is projected, with large increases during April and May (fig. 32*L*).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14*L* and table 5*E*). This increase is because of the projected increase of annual temperature (figs. 11*L* and 12*L* and tables 5*B* and 5*C*). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5*E*).

Starkweather Coulee Streamflow Variables

The PRMS projections indicate a decrease in annual mean streamflow for scenarios A1B and A2 (fig. 15*L* and table 5*F*), though the uncertainties associated with these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, particularly in April (fig. 33*L*). Projections indicate little or no streamflow during September through February, because of subfreezing temperatures in the basin, with decreases in flow during April and May.

The PRMS projections indicate decreases in surface runoff, subsurface and groundwater flow for the A1B and A2 scenarios (figs. 16*L*–18*L* and tables 5*G*–5*I*). High variability (indicating uncertainty) in the seasonal timing of these variables is projected (figs. 34*L*–36*L*), with the largest decreases in surface flow during April through August. Subsurface and groundwater flow are the dominant components of streamflow and have a similar seasonal pattern (figs. 15*L*, 17*L* and 18*L*, tables 5*F*, 5*G* and 5*I*) with the largest decreases in April.

Starkweather Coulee Snow Variables

The PRMS projections indicate no significant trend in the annual mean percentage of precipitation that falls as snow (fig. 19*L* and table 5*J*). Seasonal projections indicate little change in precipitation form (fig. 37*L*) is expected during December and January because of the cold temperatures (figs. 11*L* and 12*L*), with decreases in the percent of precipitation that falls as snow in all other snowfall months. Because of increases in temperatures, the PRMS projections indicate that the percentage of area covered by snow (fig. 20*L*) and snowpack water equivalent (fig. 21*L*) will decrease (tables 5*J* and 5*L*), with high seasonal variability in these values (figs. 38*L* and 39*L*) and decreases during November through March. PRMS projections also indicate no significant trend in annual mean snowmelt (fig. 22*L* and table 5*M*), with increases during November through February and large decreases during March and April.

Starkweather Coulee Soil Moisture Variables

The PRMS projections of infiltration indicate an increase for the A2 emission scenario (fig. 23*L* and table 5*N*). High variability in seasonal infiltration also is projected (fig. 41*L*), with increases during October through February and May and decreases during April and August.

Projected changes in the inflow to the subsurface reservoir (fig. 24*L*) and recharge to the groundwater reservoir (fig. 25*L*) are similar to the changes seen in the subsurface (fig. 17*L*) and groundwater (fig. 18*L*) flow components, respectively. The large range in the projections indicates a high level of uncertainty, with decreases projected in the inflows to the subsurface reservoirs and recharge for the A1B and A2 scenarios (tables 5*O* and 5*P*). High variability in seasonal recharge is projected (figs. 42*L* and 43*L*).

There is a decrease in the projected annual mean soil moisture for the A1B and A2 scenarios (fig. 26*L* and table 5*Q*), with the large uncertainties associated with these projections. High variability in seasonal timing of soil moisture is projected (fig. 44*L*), with decreases projected for all months except March.

The PRMS projections indicate that the storage in the subsurface (fig. 27*L*) and the groundwater reservoirs (fig. 28*L*) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5*O* and 5*P*; 5*R* and 5*S*,

respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45L and 46L).

Discussion of Starkweather Coulee Results

In the Starkweather Coulee Basin, continued above-normal precipitation along with drainage of wetlands and depressions could allow intermittent flooding conditions to occur. Additional drainage downstream could increase Devils Lake water levels causing additional damage to regional infrastructures and requiring additional spending on protective measures.

The broader-scale effects of climate change on streamflow and water storage in Starkweather Coulee Basin indicate the possibility of an overall slight drying of the basin, but the uncertainty associated with this drying is large. Alternating flood and drought conditions, which occurred in the past, are also possible. The results did not consider potential future land-cover dynamics. However, land-cover dynamics in the basin would likely be limited to the type of agriculture and drainage practiced on the landscape.

Precipitation Runoff Modeling System Projections for Trout Lake, Wisconsin

The Trout Lake Basin is located in the Northern Highlands district in north-central Wisconsin, in an area with many lakes. At USGS streamflow-gaging station 05357245, the basin has a drainage area of 120 km². The basin was simulated with a fully coupled groundwater/surface-water model (GSFLOW; Markstrom and others, 2008). The PRMS portion of the application contains 146 HRUs with elevations that range from 491 to 522 meters (table 1).

Mean monthly baseline conditions in the Trout Lake Basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from approximately -15.0 °C during January to more than 14.5 °C during July. Mean monthly maximum temperatures range from just below -6.0 °C during January to almost 26.0 °C during July and August. On a mean monthly basis, precipitation ranges from approximately 0.6 mm/day (during February) to more than 3.6 mm/day (during June and July). Because streamflow is dominated by groundwater, there is not a large amount of variability in the mean monthly distribution of streamflow.

The aquifer consists of 40–60 m of unconsolidated Pleistocene glacial sediments mostly comprised of glacial outwash sands and gravel. In this area, streamflow is dominated by groundwater contributions; however, surface runoff can occur during intense rainfall periods and spring snowmelt. Surface runoff also occurs locally in low-lying areas near streams and lakes where the unsaturated zone is thin. The Trout Lake basin (which includes Trout Lake and all four of the basins that flow into the lake) has been the focus of previous regional groundwater modeling studies including a two-dimensional analytic element screening model and three-dimensional,

finite-difference models. See Walker and Bullen (2000) for additional descriptions of the setting and Pint (2002) and Hunt and others (2005; 2008) for more description of previous modeling efforts.

The Trout Lake Basin is one of five research basins in the U.S. Geological Survey's Water, Energy and Biogeochemical Budgets program, and is collocated with the National Science Foundation funded North Temperate Lakes Long Term Ecological Research site (Magnuson and others, 1984). A major focus of the Trout Lake program is the development of a hydrologic model to allow predictions of hydrologic and biogeochemical response to future conditions including land use and climate change.

Trout Lake Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Trout Lake Basin (fig. 10M). The large range in the precipitation projections indicates a high level of uncertainty with no significant overall trend (table 5A). Seasonal variability also is apparent in these precipitation projections, especially from May through October (fig. 29M).

All PRMS projections indicate steady increases in temperature (figs. 11M, 12M, 30M, 31M, and tables 5B and 5C), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario and largest for the A2 (tables 5B and 5C).

All projections indicate steady increases in annual mean evapotranspiration (fig. 13M and table 5D). High variability in seasonal evapotranspiration is projected, with increases in all months, especially April (fig. 32M).

All PRMS projections indicate steady increases in the length of the growing season (fig. 14M and table 5E). This increase is because of the projected increase of annual temperature (figs. 11M and 12M and tables 5B and 5C). The projected changes in the length of the growing season are smallest for the B1 scenario and largest for the A2 scenario (table 5E).

Trout Lake Streamflow Variables

The PRMS projections indicate a decrease in annual mean streamflow for the A1B and A2 scenarios (fig. 15M and table 5F), though the uncertainties associated with these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, particularly during June through December (fig. 33M).

The PRMS projections indicate increases in surface runoff on an annual mean basis for the A2 scenario (fig. 16M and table 5G). High variability (indicating uncertainty) in the seasonal timing of these variables is projected (fig. 34M). The projections indicate minimal changes in subsurface flow on an annual mean basis (fig. 17M and table 5H), the largest seasonal decreases occurring in April and May (fig. 35M). The PRMS projections indicate no significant trend in groundwater

flow (fig. 18*M* and table 5*I*). High variability (indicating uncertainty) in the seasonal timing of groundwater flow is projected (fig. 36*L*).

Trout Lake Snow Variables

Because of increases in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19*M*), percentage of area covered by snow (fig. 20*M*), snowpack water equivalent (fig. 21*M*), and snowmelt (fig. 22*M*) will decrease, with the exception of the B1 emission scenario for snowmelt (tables 5*J*–5*M*). Projected warmer temperatures will: (1) result in less snowfall, which will result in less snow available to melt in the spring; (2) decrease the amount of precipitation that falls as snow, percentage of area covered by snow, and snowpack water equivalent in every month (figs. 37*A*–39*A*); and (3) increase the snowmelt in November through March, leaving less snow to melt in April and May (fig. 40*M*).

Trout Lake Soil Moisture Variables

The PRMS projections of infiltration (figs. 23*M* and 41*M*) are similar to those of precipitation (figs. 10*M* and 29*M*). The large range in the infiltration projections indicates a high level of uncertainty with no significant overall trend (table 5*N*). High variability in seasonal infiltration also is projected (fig. 41*M*).

Projected changes in the inflow to the subsurface reservoir (fig. 24*M*) and recharge to the groundwater reservoir (fig. 25*M*) are similar to the changes seen in the subsurface (fig. 17*M*) and groundwater (fig. 18*M*) flow components, respectively. The projections show slight decreases in inflows to the subsurface reservoirs and recharge on an annual mean basis (tables 5*O* and 5*P*). High variability in seasonal recharge is projected (figs. 42*M* and 43*M*), with large decreases in inflows to the subsurface reservoirs and groundwater recharge during April through May.

There is a decrease in the projected annual mean soil moisture (fig. 26*M* and table 5*Q*). High variability in seasonal timing of soil moisture is projected (fig. 44*A*), with decreases in all but the coldest months.

The PRMS projections indicate that the storage in the subsurface (fig. 27*M*) and the groundwater reservoirs (fig. 28*M*) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5*O* and 5*P*; 5*R* and 5*S*, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45*M* and 46*M*).

Discussion of Trout Lake Results

The snow-component results suggest a shorter snow-covered season, with smaller snowpack volumes and a tendency for mid-winter melts. The snow-covered season will begin later and end earlier. This has implications for winter recreation, and could potentially affect phenological

responses, with corresponding changes in the ecosystem. The recharge results suggest a shift in the timing of the recharge. The current conditions (water years 1988 through 1999) result in a hydrologic regime characterized by a buildup of water in the snowpack, and a large spring recharge occurrence before losses through transpiration begin. The projections indicate a flattening of the seasonal recharge response, with more prominent recharge pulses during the growing season and late fall. The flattening of the recharge dynamic of this northern temperate hydrologic system could result in a more southern lake-stream watershed. This change also potentially could affect the seasonal nature of water budgets to lakes and the seasonal distribution of streamflow, which are both dependent on groundwater levels. As with the snow-covered response, a shift in the timing of recharge could alter phenological responses with associated ecosystem changes.

The soil moisture results indicate a reduction of soil moisture, which could potentially change the overall vegetation in the system. This has obvious ecosystem implications, and potentially could result in a changed and less diverse plant assemblage. Further, the system likely would be more prone to fires, which dramatically could alter the hydrologic response after a fire.

One of the principle concerns in the Trout Lake area is the fate of hydrologic budgets of area lakes and resulting lake levels. This has wide ranging implications for property values, recreational use of the lakes, the hard-water/soft-water status and trophic state of the lakes, and the biotic response within the lakes. Because of the precipitation rates being relatively higher than basinwide evapotranspiration rates, flat terrain, and coarse aquifer sediments, groundwater has a strong effect on the hydrologic system in the watershed. Therefore, a more sophisticated representation of the groundwater system would likely provide a more representative view of the response of the basin hydrology to change. A coupled groundwater-surface water model (GSFLOW; Markstrom and others, 2008) can be used to predict the response of the full hydrologic system, including groundwater and lake levels, which in turn will allow a more complete assessment of the response of the system to climate and land-use change.

Precipitation Runoff Modeling System Projections for Yampa River, Colorado

The Yampa River Basin at Steamboat Springs, Colorado is part of the Green River Basin, and an important tributary of the Colorado River. At USGS streamflow-gaging station 09239500, the basin has a drainage area of 1,471 km². The basin was simulated with PRMS by Hay and others (2006b; 2006c). The PRMS model contains 68 HRUs with elevations that range from 2,124 to 3,504 meters (table 1).

Mean monthly baseline conditions in the Yampa River basin are shown in figures 5 and 6. Mean monthly minimum temperatures range from below –15.0 °C during December and January to above 5.5 °C during July. Mean monthly

maximum temperatures range from below -2.0°C during January to almost 23.0°C during July. On a mean monthly basis, precipitation ranges from less than 1.25 mm/day (during August) to 3.0 mm/day (during April). Because of freezing temperatures during the winter months, precipitation falls as snow and accumulates throughout the winter. The majority of annual mean streamflow occurs during May and June, with the peak monthly streamflow occurring during June.

The Yampa River at Steamboat Springs, Colorado was included in the Hydro-Climatic Data Network, indicating that the streamflow records prior to 1987 (when the Stagecoach Reservoir was completed) for this streamflow-gaging station relatively are “unaffected by artificial diversions, storage, or other works of man in or on the natural stream channels or in the watershed” (Slack and Landwehr, 1992). The Yampa River Basin is one of only a few river basins in Colorado where current and projected total water demand is less than the native supply (Colorado Water Conservation Board, 2002). Because of its importance as a source of water to the Colorado River, the U.S. Geological Survey conducted studies (Hay and others, 2006b; Hay and others, 2006c) of the effects of potential climate change on the water resources of the Yampa River Basin.

Tourism that results from abundant local recreational options is the largest source of revenue in the region (Steamboat on the Move, 2009), but agriculture also is important. Many of the recreational activities, such as fishing, whitewater boating, snowmobiling, and skiing, are dependent directly on the basins water resources. The Steamboat Ski area falls within the Yampa River Basin and has a base elevation of 2,203 meters and top elevation of 2,768 meters. The ski area typically receives more than 8 meters of snowfall, and operates from late November to early April.

Yampa River Input Variables

Changes in annual mean precipitation vary in the PRMS projections of the Yampa River Basin (fig. 10*N*). The large range in the precipitation projections indicates a high level of uncertainty with a positive significant trend for the A1B scenario (table 5*A*). Seasonal variability also is apparent in these precipitation projections, with increasing precipitation during December through February and decreasing precipitation in May (fig. 29*N*).

All PRMS projections indicate steady increases in temperature (figs. 11*N*, 12*N*, 30*N*, 31*N*, and tables 5*B* and 5*C*), with associated uncertainties increasing with time. The projected changes of both maximum and minimum daily temperatures are smallest for the B1 scenario (tables 5*B* and 5*C*).

All projections indicate steady increases in annual mean evapotranspiration (fig. 13*N* and table 5*D*). High variability in seasonal evapotranspiration is projected, with increases during October through May and decreases during June through August.

All PRMS projections indicate steady increases in the length of the growing season (fig. 14*N* and table 5*E*). This

increase is because of the projected increase of annual temperature (figs. 11*N* and 12*N* and tables 5*B* and 5*C*). The projected changes in the length of the growing season are smallest for the B1 scenario (table 5*E*).

Yampa River Streamflow Variables

The PRMS projections indicate a decrease in annual mean streamflow (fig. 15*N* and table 5*F*), though the uncertainties associated with these projections are large. High variability (indicating uncertainty) in the seasonal timing of streamflow is projected, with increases during March through April, decreases during June through July, and minimal change during August through February. The PRMS simulation of baseline conditions (red line, fig. 33*N*) indicate that peak streamflow occurs in June, whereas the PRMS projections indicate that, by as early as 2030, the projected month of peak streamflow will be May.

The PRMS projections indicate decreases in surface runoff for the B1 scenario (fig. 16*N* and table 5*G*), in subsurface flow for all three scenarios (fig. 17*N* and table 5*H*), and in groundwater flow for the A1B and A2 scenarios (fig. 18*N* and table 5*I*). High variability (indicating uncertainty) in the seasonal timing of these variables is projected. By the end of the 21st century, peak timing is projected to shift from June to April for surface runoff, from June to May for subsurface flow, and from July to possibly June for groundwater flow.

Yampa River Snow Variables

Because of increases in temperatures, the PRMS projections indicate that the annual mean percentage of precipitation that falls as snow (fig. 19*N*), percentage of area covered by snow (fig. 20*N*), snowpack water equivalent (fig. 21*N*), and snowmelt (fig. 22*N*) will decrease (tables 5*J*–5*M*). Seasonally, these variables have a high degree of uncertainty during the snow accumulation and melt months (figs. 37*N*–40*N*). Projected warmer temperatures result in a decrease in the percent of precipitation that falls as snow, percentage of area covered by snow, and snowpack water equivalent. The largest projected decreases in the percent of precipitation that falls as snow and percentage of area covered by snow are seen in the spring and fall transition months. The PRMS simulation shows maximum snowpack water equivalent for baseline conditions (red line, fig. 33*N*) occurring in March and April, whereas the PRMS projections indicate that the snowpack water equivalent maximums will shift to March by the end of the 21st century. Projections also indicate increases in snowmelt during March and April and decreases in June. Despite this, there is no change in the peak timing of snowmelt (May) projected (fig. 40*N*). Warmer temperatures also will result in less snowfall so that less snow is available to melt during May through July.

Yampa River Soil Moisture Variables

The PRMS projections of infiltration (figs. 23*N* and 41*N*) are similar to those of precipitation (figs. 10*N* and 29*N*). The large range in the infiltration projections indicates a high level of uncertainty with no significant overall trend (table 5*N*). High variability in seasonal infiltration also is projected (fig. 41*N*), with large increases during March and April followed by large decrease during June.

Projected changes in the inflow to the subsurface reservoir (fig. 24*N*) and recharge to the groundwater reservoir (fig. 25*N*) are similar to the changes seen in the subsurface (fig. 17*N*) and groundwater (fig. 18*N*) flow components, respectively. The projections show significant negative trends in inflow to the subsurface reservoir for scenarios A1B and A2 (table 5*O*). The large range in the projections of annual mean recharge indicates a high level of uncertainty, especially during the snow accumulation and melt seasons.

There is a decrease in the projected annual mean soil moisture (fig. 26*A* and table 5*Q*). High variability in seasonal timing of soil moisture is projected (fig. 44*N*), with decreases during June through September and increases during March and April.

The PRMS projections indicate that the storage in the subsurface (fig. 27*N*) and the groundwater reservoirs (fig. 28*N*) reach equilibrium quickly, resulting in simulated values similar to the inflows (see tables 5*O* and 5*P*; 5*R* and 5*S*, respectively). High variability in seasonal timing of subsurface and groundwater storage is projected (figs. 45*A* and 46*A*).

Discussion of Yampa River Results

Streamflow in the Yampa River Basin is under increasing demand from water users and energy companies in the southwestern United States, and recreationalists within the basin. Potential decreases in streamflow resulting from future changes in climate may add to the stress that this basin will experience as a result of projected increases in domestic and industrial water use (Colorado Water Conservation Board, 2006b). The effects of climate change in the Yampa River Basin may also result in earlier snowmelt resulting in earlier spring runoff and affecting the conditions that support recreational activities such as skiing.

Integrated Watershed-Scale Response to Climate Change for Selected Basins Across the United States

Intercomparisons between the PRMS projections made for the 14 basins were difficult to make because the PRMS applications were set up in different ways for different purposes. For example, how the HRUs were delineated, selection of input parameters related to routing and soil moisture computations, and differences in how the measured climate data

was distributed to the HRUs resulted in simulated results that were difficult to compare.

In addition to this report, the results of this study have been summarized and evaluated in 10 different publications. Results produced by this study have been evaluated in the following publications (table 1):

1. Battaglin and others (2011) identifies the potential effects of future climate on hydrologic conditions for two Colorado basins, the East River at Almont and the Yampa River at Steamboat Springs, and for two ski areas within those basins. All Precipitation Runoff Modeling System projections of the scenarios indicate decreases in basin mean snow-covered area and snowpack water equivalent, with uncertainty around the projected decreases increasing with time. However, when examined on a monthly basis, the changes are most dramatic during fall and spring and less important during winter. Presumably, ski area locations are picked because of a tendency to receive snow and retain snowpack relative to the surrounding area. This effect of ski area location within the basin was examined by comparing simulations of March snow-covered area and snowpack water equivalent for the entire basin with simulations for the portion of the basin that represents the ski area in the PRMS applications. These projections indicate only small changes in March snow-covered area at both ski areas for the three scenarios until around 2050. After 2050, larger decreases are projected, but also there is large uncertainty. Results from this modeling effort show that there is a wide range of possible outcomes for projected snowpack conditions in Colorado. The results also highlight the differences between projections for entire basins and projections for particular locations within those basins.
2. Bjerklie and others (2011) developed a regional watershed model using geographic information system-based parameters for watersheds contributing to Long Island Sound, including the Connecticut River Basin. The study region covers approximately 15,800 square miles, extending from a moderate coastal climate zone in the south to a mountainous northern New England climate zone dominated by snowmelt in the north. The input data, including daily precipitation and maximum and minimum daily temperature for 47 years (1961–2007), indicate that precipitation and temperature have been increasing across the modeled area. Although both maximum and minimum daily temperature have been increasing, minimum temperature has increased more than maximum temperature. The Precipitation Runoff Modeling System simulation indicates that there is an upward trend in groundwater recharge across most of the modeled region. The trend in simulated snowfall generally is not significant across the region, except in the extreme

southeast part of Connecticut and in parts of central Massachusetts where it has been downward.

3. Christiansen and others (2011) examine trends in the total length (annual), as well as changes in the timing of onset (spring) and the end (fall) of the growing season. The results showed an increase in the annual growing season length in all 14 basins of this study, averaging 27 to 47 days for the 3 scenarios. The change in the spring and fall growing season onset and end varied across the 14 basins, with larger increases in seasonal change occurring in the mountainous regions, and smaller increases occurring in the midwest, northeast, and southeast regions.
4. Hay and others (2011) present an overview of this study, details of the methodology, results from the 14 basin simulations, and interpretation of these results. A key finding is that the hydrological response of the different geographical regions of the United States to potential climate change may be different, depending on the dominant physical processes of that particular region. Also considered is the tremendous amount of uncertainty present in the climate emission scenarios and how this uncertainty propagates through the hydrologic simulations. This report concludes with a discussion of the lessons learned and potential for future work.
5. Koczo and others (2011) evaluate the projected changes in temperature and precipitation from several climate change emission scenarios are applied using three different baseline conditions, each corresponding to measured temperature and precipitation data from specified historical periods during the 20th century. Differences and similarities between projections of hydrologic components (that is, snowpack formation and melt, evapotranspiration, and streamflow) are examined and results indicate that the selection of a specific time period used for baseline conditions has a substantial effect on some, but not all, hydrologic variables. This effect seems to be amplified in hydrologic variables which accumulate over time, for instance actual evapotranspiration and soil-moisture content. Projections also indicate that to account for uncertainty related to the selection of baseline conditions, a range of different baseline conditions need to be considered.
6. Mastin and others (2011) find a future of declining spring snowpack in the 8 snowmelt dominated basins of the 14. Snow-covered area and snow-water equivalent were used to compare the spring snowpack for current conditions (2006) with three time periods in the future (2030, 2060, and 2090) using the three scenarios. Distributed snow-water equivalent and snow-covered area values were evaluated by elevation zones. The change in the spring snowpack over time was greater than the change among scenarios suggesting that even for a globally reduced carbon emission scenario, large decreases in snow-water equivalent likely are to occur.
7. Milly and Dunne (2011) analyze the effects of hydrologic adjustment on the projections of runoff change associated with projected 21st-century climate change. Results in a case study including 3 climate models and 10 river basins in the contiguous United States found that relative (that is, fractional or percentage) runoff change computed with hydrologic adjustment more often than not was less positive (or, equivalently, more negative) than what was projected by the climate models. The dominant contributor to this decrease in runoff was a ubiquitous change in runoff (median -11 percent) caused by the hydrologic model's apparent amplification of the climate-model-implied growth in potential evapotranspiration.
8. Risley and others (2011b) used projected daily streamflow and energy output from the Precipitation Runoff Modeling System applications to compute a range of statistics. With a side-by-side comparison of the statistical analyses for the 14 basins, regional climatic and hydrologic trends over the 21st century could be identified. Low-flow statistics (95-percent exceedance, 7-day mean annual minimum, and summer mean monthly streamflow) decreased for almost all basins. Annual maximum daily streamflow also decreased in all the basins except for all four basins in California and the Pacific Northwest. An analysis of the supply of available energy and water for the basins indicated that ratios of evaporation to precipitation and potential evapotranspiration to precipitation for most of the basins will increase. Probability density functions were developed to assess the uncertainty and multimodality in the effect of climate change on mean annual streamflow variability. Probability density functions for two basins in the upper Midwest were more dispersed and skewed than for all the other basins. For most of the 14 basins, the probability density shape did not change over the 21st century.
9. Viger and others (2011) examine the projected effects of long-term urbanization and climate change on the freshwater resources of the Flint River basin. Projections of urbanization through 2050 derived for the Flint River basin by the FOREcasting SCEnarios (FORE-SCE, Sohl and Saylor, 2008) land-cover change model also were used as input to the PRMS application. Comparison of the central tendency of streamflow simulated based on the three climate-change scenarios showed a slight decrease in overall streamflow relative to simulations under

current conditions (water years 1988 through 1999), mostly caused by decreases in the surface-runoff and groundwater components. The addition of information about forecasted urbanization of land surfaces to the hydrologic simulation mitigated the decreases in streamflow, mainly by increasing surface runoff.

10. Walker and others (2011a) fitted empirical orthogonal functions to the Precipitation Runoff Modeling System model output driven by the ensemble of climate projections, and provided a basis for randomly (but representatively) generating realizations of hydrologic response to future climates. For each realization, the 1.5-year flood was calculated to represent a flow important for sediment transport and channel geomorphology. The empirical probability density function of the 1.5-year flood was estimated using the results across the realizations, for each basin. Nine of the 14 basins studied showed clear temporal shifts in the probability density functions of the 1.5-year flood projected into the 21st century. In the Western United States, where the annual peak discharges are heavily influenced by snowmelt, three basins show at least a 10 percent increase in the 1.5-year flood in the 21st century; the remaining two basins demonstrate increases in the 1.5-year flood, but the temporal shifts in the probability density functions and the percent changes are not as distinct. Four basins in the Eastern Rockies/Central United States show at least a 10 percent decrease in the 1.5-year flood; the remaining two basins demonstrate decreases in the 1.5-year flood, but the temporal shifts in the probability density function and the percent changes are not as distinct. Two basins in the eastern United States show at least a 10 percent decrease in the 1.5-year flood; the remaining basin shows little or no change in the 1.5-year flood.

Summary

The simulated hydrologic response of different climate change emission scenarios for the 21st century were evaluated in 14 basins from different hydroclimatic regions across the United States using the Precipitation Runoff Modeling System, a process-based, distributed-parameter watershed model. This study involved five major steps: (1) setup, calibrate, and evaluate the Precipitation Runoff Modeling System (PRMS) model in 14 basins across the United States by local USGS personnel; (2) acquire selected simulated carbon emission scenarios from the World Climate Research Programme's Coupled Model Intercomparison Project; (3) statistical downscaling of these scenarios to create PRMS input files which reflect the future climatic conditions of these scenarios; (4) generate PRMS projections for the carbon emission scenarios for the 14 basins; and (5) analyze the simulated hydrologic response.

All climate change emission scenarios evaluated with PRMS show a steady increase in maximum and minimum temperatures progressing through the 21st century. This generally results in projections of earlier spring snowmelt and increases in evapotranspiration, if soil moisture is available. If these projected increases in temperature hold true, then the largest changes in the timing of streamflow will occur in the snowmelt-dominated basins with winter temperatures close to the freezing point. This has long been recognized that an increase in winter temperature initially resulting in less snow (change of form from snow to rain) and eventually earlier snowmelt. The PRMS projections for the three emission scenarios indicate that streamflow could increase in the winter and early spring and decrease in late spring and through the summer for many of the basins; however, uncertainty associated with the magnitude of these changes in streamflow is large. Basins in which precipitation form (rain or snow) is sensitive to the air temperature are the most sensitive to changes in temperature and precipitation. This is consistent with the findings of previous work: warmer temperatures at mid-elevation mountainous regions show decreasing snowpack and earlier melt despite the increase in precipitation, whereas high-elevation regions that remain well below freezing during winter show little effects.

A change in timing and volume of snowmelt has large implications for water managers who rely on the melt from the spring snowpack to replenish reservoirs and provide early season irrigation; earlier snowmelt may lead to increased water scarcity. Even when the annual streamflow shows a positive trend into the 21st Century, spring streamflow from snowmelt is projected to decrease, and therefore, water managers may need to store more winter streamflow to meet the irrigation and instream flow demands. Increasing temperatures and population will invariably lead to increasing water demands, making water-management decisions increasingly difficult due to the uncertainty in the quantity and timing of the spring snowmelt.

The ecological implications of changes in the timing of and volume of snowmelt are not well understood. The PRMS projections indicate changes in the timing of peak flows due to snowmelt in all but the coldest basins, which may affect fresh-water mixing in estuarine ecosystems and increase the length of the summer drought that characterizes much of western North America. A shorter snow-covered season, with smaller snowpack volumes, and a tendency for mid-winter melts has implications for winter recreation, but could also affect phenological responses, with corresponding changes in the ecosystem. Earlier snowmelt and increases in evapotranspiration rates may result in drier forests and more wildfires and threaten fish and amphibian habitat. Projected reductions in soil moisture could lead to changes in overall vegetation. This has ecological implications and could result in a changed and less diverse plant assemblage. Further, the basins could be more prone to wildfires.

This study did not consider future population growth and land-use changes. It is possible that changes in land cover can

affect climate at the regional and local scales. The methodologies for adequate simulation of changes in urban systems, agricultural systems, ecosystem disturbance regimes, and soil effects are not yet well developed. The combined effects of climate change and urbanization may alter both the quantity and timing of streamflow and can change the conditions of water quality that support biological diversity in aquatic communities. Also not addressed, is whether the adverse effects of climate change can be mitigated with careful land-use planning.

In many of the basins, projected decreases in soil moisture may increase agricultural water consumption, resulting in an increase in overall water demand. A reduction in soil moisture will cause drier conditions in the root zone likely causing increased irrigation groundwater withdrawals to maintain current agricultural production levels. Such simulations are beyond the scope of the methods used in this study, as simulating effects of increased irrigation requires a coupled groundwater-surface water model.

When evaluating the projections developed for the 14 basins, it was apparent that the consistent application of the hydrology model in each basin would lead to more comparable results in inter-regional studies of this type. Of particular benefit, would be a hydrologic modeling structure which would help address the issues of (1) differing temporal and spatial scales and resolutions, (2) evolving data availability and needs, (3) differing calibration methods and purpose, and (4) internal and external institutional constraints. Future research, including the development of a national hydrologic modeling structure with consistent formulation, would enhance this type of climate change study.

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Appendix

Appendix. Documentation of New PRMS Modules

Introduction

This appendix describes the inputs and outputs, equations, computations, references, assumptions, and limitations of the new PRMS modules that were developed for this study. Each described module is encoded as a Fortran 90 source code file.

Styles and Formats

The following font styles and formats are used in this Appendix:

- Module variables are identified 12-point, italic, Times New Roman font.
- Module parameters and dimensions are identified 12-point, bold, Times New Roman font.
- Constants are identified using upper case 12-point, bold, Times New Roman font.
- Module names are identified in 12-point, Courier font.

Transpiration Period Modules

The period of active transpiration is determined for each HRU by one of two user-specified options. These options are described below.

Table A1–1. Input parameters to module `transp_frost_prms`.

[HRU: hydrologic response unit; nhru: number of HRUs; nmonths: number of months in a year]

Parameter name	Description	Dimension variable	Units	Type	Range	Default value
spring_frost	The solar date (number of days after winter solstice) of the last killing frost of the spring	nhru	solar day of year	integer	1 to 366	111
fall_frost	The solar date (number of days after winter solstice) of the first killing frost of the fall	nhru	solar day of year	integer	1 to 366	264

Table A1–2. Variables used in module `transp_frost_prms`.

[HRU: hydrologic response unit; one: a constant; nhru: number of HRUs]

Variable name	Description	Dimension	Units	Type
Input Variables				
<i>active_hrus</i>	Number of active HRUs	ONE	none	long
<i>hru_route_order</i>	Routing order for HRUs	nhru	none	long
Output Variables				
<i>basin_transp_on</i>	Switch indicating whether transpiration is occurring anywhere in the basin (0=no; 1=yes)	ONE	none	long
<i>transp_on</i>	Switch indicating whether transpiration is occurring (0=no; 1=yes)	nhru	none	long

transp_frost_prms

The first option for determination of the period of active transpiration, or growing season length, (module `transp_frost_prms`) uses the killing frost approach (Christiansen and others, 2011). The growing season length can be defined as the time between the last and first freezing air temperatures (U.S. Department of Agriculture, National Agricultural Statistics Service, 1999) and has been used in numerous studies (Wang 1963; Brinkman, 1979; Cooter and LeDuc, 1995; Kunkel and others, 2004). This module defines the growing season length as the period between the last killing frost in the spring and the first killing frost in the fall (U.S. Army Corps of Engineers, 1987).

The user specifies the solar dates of the last killing frost of the spring and the first killing frost of the fall for each HRU. The period of active transpiration are the time steps between these dates. This can be determined by using the `frost_date_prms` module (described below in section “Pre-process Module: `frost_date_prms`”), saving HRU temperatures in the Statvar Output File, and computing the average solar date for the last killing frost of the spring and the first of the fall. The input parameters used to determine period of active transpiration are defined in table A1–1. The computed variables are defined in table A1–2.

transp_tindex_prms

The second option for determination of the period of active transpiration (module `transp_tindex_prms`) uses a temperature index approach. This module computes a temperature index, which is the sum of the daily maximum temperature for each HRU once the model reaches the transpiration starting trigger date (parameter **transp_beg**).

The period of active transpiration for each HRU actually begins once the temperature index reaches the temperature index threshold (parameter **transp_tmax**). The period of transpiration for each HRU is terminated when the simulation time step reaches the date specified by the **transp_end** parameter. The input parameters used to determine period of active transpiration are defined in table A1–3. The computed variables are defined in table A1–4.

Table A1–3. Input parameters to module `transp_tindex_prms`.

[HRU: hydrologic response unit; nhru: number of HRUs; nmonths: number of months in a year; one: a constant]

Parameter name	Description	Dimension variable	Units	Type	Range	Default value
hru_area	Area of each HRU	nhru	acres	float	0.01 to 1.0e9	1
temp_units	Units for measured temperature (0=Fahrenheit; 1=Celsius)	ONE	none	long	0 to 1	0
transp_beg	Month to begin summing <i>tmaxf</i> or <i>tmaxc</i> (table A1–4) for each HRU; when sum is >= to transp_tmax , transpiration begins	nhru	month	long	1 to 12	4
transp_end	Month to stop transpiration computations; transpiration is computed thru end of previous month	nhru	month	long	1 to 12	10
transp_tmax	Temperature index to determine the specific date of the start of the transpiration period. Subroutine sums <i>tmaxf</i> or <i>tmaxc</i> (table A1–4) for each HRU starting with the first day of month transp_beg . When the sum exceeds this index, transpiration begins	nhru	degrees	float	0 to 1000	500

Table A1–4. Variables used in module `transp_tindex_prms`.

[HRU: hydrologic response unit; one: a constant; nhru: number of HRUs; F, Fahrenheit]

Variable name	Description	Dimension	Units	Type
Input variables				
<i>active_hrus</i>	Number of active HRUs	ONE	none	long
<i>hru_route_order</i>	Routing order for HRUs	nhru	none	long
<i>tmaxc</i>	HRU adjusted daily maximum temperature	nhru	degrees Celsius	float
<i>tmaxf</i>	HRU adjusted daily maximum temperature	nhru	degrees F	float
Output variables				
<i>basin_transp_on</i>	Switch indicating whether transpiration is occurring anywhere in the basin (0=no; 1=yes)	ONE	none	long
<i>transp_on</i>	Switch indicating whether transpiration is occurring (0=no; 1=yes)	nhru	none	long

Potential Evapotranspiration Modules

potet_jh_prms

Potential evapotranspiration (PET) is computed for each HRU by one of three user-specified options. The first option (module `potet_jh_prms`) is the modified Jensen-Haise formulation (Jensen and others, 1970), in which PET is computed as a function of air temperature, solar radiation, and two coefficients that can be estimated using regional air temperature, elevation, vapor pressure, and plant cover. The input parameters used to compute the potential evapotranspiration variables and are defined in table A1-5. The computed

variables are defined in table A1-6. Equation symbols, used in this section of the report, which correspond directly to PRMS input or output, are defined in these two tables.

PET is calculated as:

$$potet_{HRU}^m = jh_coef_{month} (tavgf_{HRU} - jh_coef_hru) \frac{swrad_{HRU}^m}{2.54\lambda_{HRU}} \quad (1)$$

$$\lambda_{HRU} = 597.3 - (0.5653tavgf_{HRU}) \quad (2)$$

where

λ_{HRU} is the latent heat of vaporization on the HRU for time step m , calories/gram

Table A1-5. Input parameters to module `potet_jh_prms`.

[HRU: hydrologic response unit; nhru: number of HRUs; nmonths: number of months in a year; one: a constant]

Parameter name	Description	Dimension variable	Units	Type	Range	Default value
hru_area	Area of each HRU	nhru	acres	float	0.01 to 1.0e9	1
jh_coef	Monthly air temperature coefficient used in Jensen-Haise potential evapotranspiration computations.	nmonths	per degrees	float	0.005 to 0.06	0.014
jh_coef_hru	Air temperature coefficient used in Jensen-Haise potential evapotranspiration computations for each HRU.	nhru	per degrees	float	5 to 20	13
temp_units	Units for measured temperature (0=Fahrenheit; 1=Celsius)	ONE	none	long	0 to 1	0

Table A1-6. Variables used in module `potet_jh_prms`.

[HRU: hydrologic response unit; one: a constant; nhru: number of HRUs; F, Fahrenheit]

Variable name	Description	Dimension	Units	Type
Input variables				
<i>active_hrus</i>	Number of active HRUs	ONE	none	long
<i>basin_area_inv</i>	Inverse of total basin area as sum of HRU areas	ONE	1/acres	float
<i>hru_route_order</i>	Routing order for HRUs	nhru	none	long
<i>route_on</i>	Kinematic routing switch (0=daily; 1=storm period)	ONE	none	long
<i>swrad</i>	Computed shortwave radiation for each HRU	nhru	langleys	float
<i>tavgc</i>	HRU adjusted daily average temperature	nhru	degrees Celsius	float
<i>tavgf</i>	HRU adjusted daily average temperature	nhru	degrees F	float
<i>tmaxc</i>	HRU adjusted daily maximum temperature	nhru	degrees Celsius	float
<i>tmaxf</i>	HRU adjusted daily maximum temperature	nhru	degrees F	float
<i>newday</i>	Switch signifying the start of a new day (0=no; 1=yes)	ONE	none	long

Table A1–6. Variables used in module `potet_jh_prms`.—Continued

[HRU: hydrologic response unit; one: a constant; nhru: number of HRUs; F, Fahrenheit]

Variable name	Description	Dimension	Units	Type
Output variables				
<i>basin_potet</i>	Basin area-weighted average of potential evapotranspiration	ONE	inches	float
<i>potet</i>	Potential evapotranspiration on an HRU	nhru	inches	float

potet_hamon_hru_prms

The second option (module `potet_hamon_hru_prms`) for computing PET is the empirical Hamon formula-tion, in which PET is computed as a function of daily mean air temperature and possible hours of sunshine according to (Hamon, 1961; Murray, 1967; and Federer and Lash, 1978). The input parameters used to compute the PET variables are defined in table A1–7. The computed variables are defined in table A1–8. Equation symbols, used in this section of the report, which correspond directly to PRMS input or output, are defined in these two tables. PET is calculated as:

$$potet_{HRU}^m = hamon_coef_{HRU} (sh_{HRU}^m)^2 \rho_{HRU} \quad (3)$$

$$\rho_{HRU} = 216.7 \frac{6.108e^{\left(\frac{17.26939 \cdot avgc_{HRU}}{avgc_{HRU} + 273.3}\right)}}{avgc_{HRU} + 273.3} \quad (4)$$

where

sh_{HR}^m is length of the daylight period, for time step m , radians per 12-hour period, and

ρ_{HRU} is the saturated water-vapor density (absolute humidity), grams per cubic meter.

potet_pan_prms

The third option (module `potet_pan_prms`) is used when pan evaporation data from one or more measurement stations are available. The station associated with each HRU is specified by parameter **hru_pansta** (table A1–9). PET is computed from the measured pan evaporation and a monthly coefficient. The input parameters used to compute the potential evapotranspiration variables and are defined in table A1–9. The computed variables are defined in table A1–10. Equation symbols, used in this section of the report, which correspond directly to PRMS input or output, are defined in these two tables. PET is calculated as:

$$potet_{HRU}^m = epan_coef_{month} pan_evap_{sta}^m \quad (5)$$

Table A1–7. Input parameters to module `potet_hamon_hru_prms`.

[HRU: hydrologic response unit; nmonths: number of months in a year; nhru: number of HRUs; one: a constant; F, Fahrenheit]

Parameter name	Description	Dimension variable	Units	Type	Range	Default value
hamon_coef	Monthly air temperature coefficient used in Hamon potential evapotranspiration computations, see PRMS manual	nmonths	inch-cubic meter per gram	float	0.004 to 0.008	0.0055
hru_area	Area of each HRU	nhru	acres	float	0.01 to 1.0e9	1
temp_units	Units for measured temperature (0=Fahrenheit; 1=Celsius)	ONE	none	long	0 to 1	0

Table A1–8. Variables used in module `potet_hamon_hru_prms`.

[HRU: hydrologic response unit; one: a constant; nhru: number of HRUs; ndays: number of days in a year]

Variable name	Description	Dimension	Units	Type
<i>active_hrus</i>	Number of active HRUs	ONE	none	long
<i>basin_area_inv</i>	Inverse of total basin area as sum of HRU areas	ONE	1/acres	float
<i>hru_route_order</i>	Routing order for HRUs	nhru	none	long
<i>route_on</i>	Kinematic routing switch (0=daily; 1=storm period)	ONE	none	long
<i>soltab_sunhrs</i>	Hours between sunrise and sunset for each HRU	ndays, nhru	hours	float
<i>tavgc</i>	HRU adjusted daily average temperature	nhru	degrees Celsius	float
<i>tmaxc</i>	HRU adjusted daily maximum temperature	nhru	degrees Celsius	float
<i>tmaxf</i>	HRU adjusted daily maximum temperature	nhru	degrees Fahrenheit	float
<i>newday</i>	Switch signifying the start of a new day (0=no; 1=yes)	ONE	none	long
Output variables				
<i>basin_potet</i>	Basin area-weighted average of potential et	ONE	inches	float
<i>potet</i>	Potential evapotranspiration on an HRU	nhru	inches	float

Table A1–9. Input parameters to module `potet_pan_prms`.

[HRU: hydrologic response unit; nmonths: number of months in a year; nhru: number of HRUs]

Parameter name	Description	Dimension variable	Units	Type	Range	Default value
epan_coef	Evaporation pan coefficient	nmonths	none	float	0.2 to 3	1
hru_area	Area of each HRU	nhru	acres	float	0.01 to 1.0x10 ⁹	1
hru_pansta	Index of pan evaporation station used to compute HRU potential evapotranspiration	nhru	none	long	0 to 0	0
transp_beg	Month to begin summing <i>tmaxf</i> or <i>tmaxc</i> (table A1–4) for each HRU; when sum is >= to transp_tmax (table A1–3), transpiration begins	nhru	month	long	1 to 12	4
transp_end	Month to stop transpiration computations; transpiration is computed thru end of previous month	nhru	month	long	1 to 12	10
transp_beg	Month to begin summing <i>tmaxf</i> or <i>tmaxc</i> (table A1–4) for each HRU; when sum is >= to transp_tmax (table A1–3), transpiration begins	nhru	month	long	1 to 12	4
transp_end	Month to stop transpiration computations; transpiration is computed thru end of previous month	nhru	month	long	1 to 12	10

Table A1–10. Variables used in module `potet_pan_prms`.

[HRU: hydrologic response unit; one: a constant; nhru: number of HRUs; nevap: number of climate stations that measure pan evaporation]

Variable name	Description	Dimension	Units	Type
Input variables				
<i>active_hrus</i>	Number of active HRUs	ONE	none	long
<i>basin_area_inv</i>	Inverse of total basin area as sum of HRU areas	ONE	1/acres	float
<i>hru_route_order</i>	Routing order for HRUs	nhru	none	long
<i>pan_evap</i>	Measured pan evaporation at each measurement station	nevap	inches	float
<i>route_on</i>	Kinematic routing switch (0=daily; 1=storm period)	ONE	none	long
<i>newday</i>	Switch signifying the start of a new day (0=no; 1=yes)	ONE	none	long
Output variables				
<i>basin_potet</i>	Basin area-weighted average of potential et	ONE	inches	float
<i>potet</i>	Potential evapotranspiration on an HRU	nhru	inches	float

Preprocess Module: frost_date_prms

Determination of the mean last killing frost of the spring and the first killing frost of the fall can be done with the PRMS preprocessing module `frost_date_prms`. This module is activated by running the PRMS executable with the `-preprocess` argument on the command line. For each

HRU, this module will determine the solar year dates defining the active period of transpiration for each year in the specified period of simulation. The mean values of these dates are written out to a PRMS Parameter File at both the HRU and basin resolution. The input parameters used to compute the frost date outputs and are defined in table A1–11. The computed variables are defined in table A1–12.

Table A1–11. Input parameters to module `frost_date_prms`.

[HRU: hydrologic response unit; nhru: number of HRUs; one: a constant]

Parameter name	Description	Dimension variable	Units	Type	Range	Default value
frost_temp	Temperature of the killing frost	ONE	temperature degrees	float	-10.0 to 32.0	28.0
temp_units	Units for measured temperature (0=Fahrenheit; 1=Celsius)	ONE	none	long	0 or 1	0
hru_area	Area of each HRU	nhru	acres	float	0.01 to 1.0x10 ⁹	1.0

Table A1–12. Variables used in module `frost_date_prms`.

[HRU: hydrologic response unit; one: a constant; nhru: number of HRUs]

Variable name	Description	Dimension	Units	Type
Input variables				
<i>active_hrus</i>	Number of active HRUs	ONE	none	long
<i>basin_area_inv</i>	Inverse of total basin area as sum of HRU areas	ONE	1/acres	float
<i>hru_route_order</i>	Routing order for HRUs	nhru	none	long
<i>route_on</i>	Kinematic routing switch (0=daily; 1=storm period)	ONE	none	long
<i>soltab_sunhrs</i>	Hours between sunrise and sunset for each HRU	ndays, nhru	hours	float

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